

**Fluid Physics**  
**12.330J/8.292J**  
Spring 2004

**COURSE OUTLINE**

**Introduction: Outstanding problems in fluid physics**

**Definition of fluid; statistical mechanics**

**Hydrostatics, mass continuity**

**The Bernoulli Equation**

**The shallow-water equations**

- Gravity waves
- Hydraulic jumps, bores
- Surface waves

**Potential flow**

- Kelvin's circulation theorem
- Forces on objects
- Some real-world problems

**Compressible flow**

- Thermodynamics
- Sound waves
- Shocks
- Bernoulli Equation for compressible flows

**Stratified flows**

- Internal waves
- Flow over topography

**Rotating flows**

- Barotropic flows, point vortices
- Rossby waves
- Baroclinic flows

**Instabilities**

- Convection
- Rayleigh shear flow instability
- Kelvin-Helmholtz instability

**Viscous flows**

- Stress
- The Navier-Stokes equations
- Viscous dissipation
- Ekman layers
- Turbulence

**Physics of fluid phenomena**

- Stellar dynamics
- Hurricanes
- Superfluids
- Other

Room

8-302

# Fluid Physics

12.330 J / 8.292 J

Spring 2000

Lecture 1: Go over structure of course  
Introduce Greg Lawson (Rob Newb, 2002)

Many outstanding enigmas in physics involve fluids  
but physics of fluids rarely taught. (Engineering courses)

One emphasis will be on physical (geo-astr) systems

Course:  $\gamma_2$  Theory  
 $\gamma_2$  applications to physical systems

Reading Assignment: Faber (note page xix)  
Sections 1.1-1.3

## Some outstanding problems

→ Photo of Sun

### 1. Stellar physics

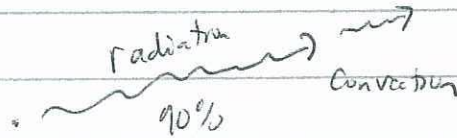
Ordinary star: Self-gravitating, rotating gaseous sphere

Radius changes in  $10^{10}$  years (long)

Sun would collapse in  $\sim 10^3$  s were it not for

gas pressure  
Heat transfer:

Radiative in innermost 90%



But solar granulation: 5 minute time scale

No generally accepted theory for convective heat transport, modified by rotation, magnetic fields

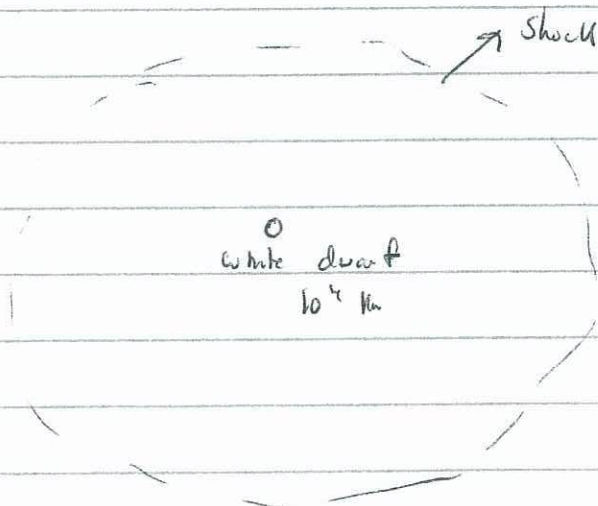
Sunspots : Streams of ionized gas

Solar wind : What drives it?

No good theory for observed variation in solar output

\* Graph of solar output

Supernovae :



Neutrons absorbed in inner regions, produces shock wave that blows off envelope of gas.

But shock wave stops at finite radius, envelope should collapse! But it does not?!

### 3. Climate Dynamics

Picture of Earth

Heat Flux by Oceans and Atmospheres

Paradox of Thermohaline circulation

Importance of mixing

### 4. Spinal Arms in Tropical Cyclones

Pictures

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Fluids:

Definition: No rigidity

Solid has definite shape that changes only in response to changing external conditions. Fluid has no preferred shape

Simple solid: material whose change in shape is small in response to a small force.

But: Jelly, paint, pitch, ice

Fluids ~~do~~ not in motion exert no stress

Fluid characteristics:

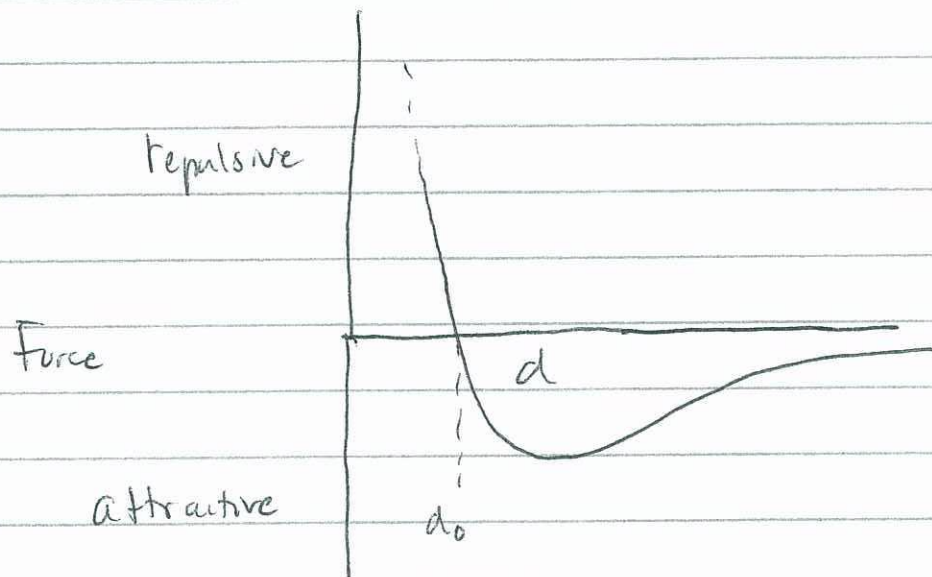
- Density
- Viscosity
- Compressibility



What determines whether a fluid is a liquid or gas?

Molecular structure and intermolecular forces

At  $d \lesssim 10^{-8}$  cm, "strong force" operates:



Gas: average spacing  $> 10 d_0$

Liquid or solid:  $d \approx d_0$

Perfect Gas: Potential energy of each molecule within the force field of its neighbor is negligible compared to its kinetic energy.

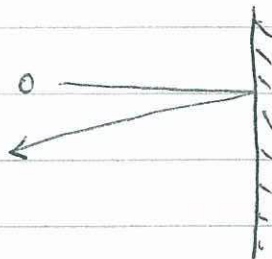
Each molecule moves independently, except for collisions.

Liquid: each molecule well within ~~KE > PE~~ force field of neighbors.

Derivation of ideal gas law:

Each molecule moves independently

Collisions are perfectly elastic



$$\text{Force / unit area} = \frac{\text{number of molecules} \times \text{mass of each molecule} \times \frac{du}{dt}}{\text{unit area}}$$

But characteristic time between collisions is  $\Delta t = \frac{l}{u}$ , where  $l$  is intermolecular spacing

$$P = \frac{\text{number of molecules}}{\text{unit area}} \times m \times \frac{u du}{\Delta t}$$

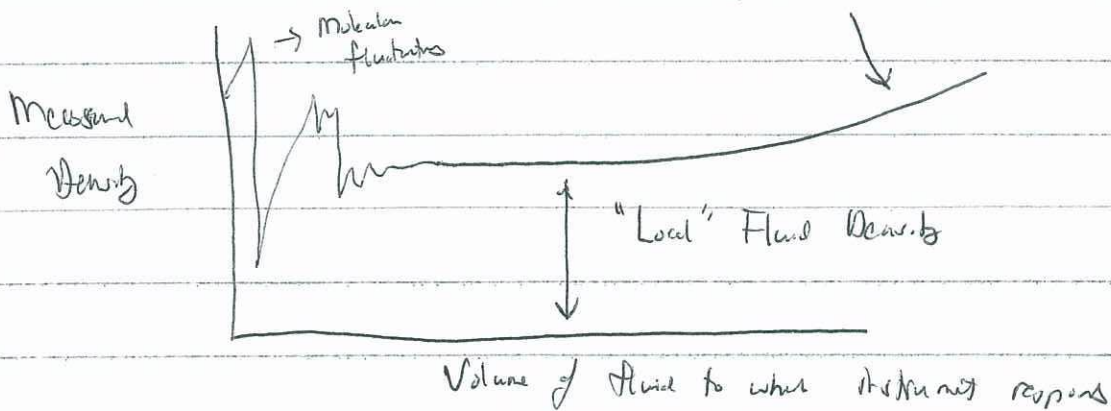
$$= \frac{\text{number of molecules}}{\text{unit volume}} \times m \times u du$$

$$= \rho \int u du = \rho \frac{1}{2} u^2 = \rho RT$$

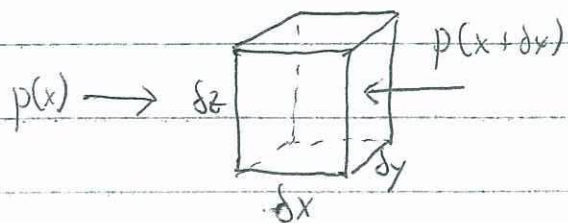
Phase	Intermolecular Forces	Amplitude of random thermal movement to do	Molecular Arrangement	Type of statistics needed
Solid	strong	$\ll 1$	ordered	quantum
Liquid	medium	$\sim 1$	partially ordered	quantum + classical
Gas	weak	$\gg 1$	disordered	classical

## Continuum hypothesis:

Macroscopic behavior of fluids is the same as if they were continuous in structure



**Pressure:** The force per unit area exerted on an infinitesimal area in a direction normal to the area. A result of molecular collisions.



Force in x direction on element:  $p(x) \delta y \delta z - p(x+\delta x) \delta y \delta z$

Hydrostatics:

Ideal gas:  $p = \rho \frac{R^*}{m} T \equiv \rho R T$

$$\frac{\partial p}{\partial z} = -\rho g = -\frac{\rho g}{RT}$$

Special case of isothermal gas:  $T = \text{constant}$

$$\frac{\partial \ln p}{\partial z} = -\frac{1}{H} \quad H \equiv \frac{RT}{g}$$

$$\left[ p = p_0 e^{-z/H} \right] \quad \text{Exponential}$$

Earth:  $H \sim 8 \text{ km.}$

Star: Self-gravitating sphere  
neglect rotation  
Hydrostatic equilibrium

Mass in spherical shell of thickness  $\delta r$ :

$$\delta m = 4\pi r^2 \rho \delta r$$

Hydrostatic:

$$\frac{\partial p}{\partial r} = -\rho \frac{G m}{r^2}$$



~~2~~  
Insert, page 2

$$\Rightarrow \frac{\partial}{\partial r} \left( \frac{r^2}{\rho} \frac{\partial \rho}{\partial r} \right) + G \frac{dm}{dr} = 0$$

$$\left[ \frac{\partial}{\partial r} \left( \frac{r^2}{\rho} \frac{\partial \rho}{\partial r} \right) + 4\pi r^2 \rho G = 0 \right]$$

" Polytropic sphere:

$$\rho = K p^{1/\gamma}$$

Substitution:  $z = p^{1-1/\gamma}$

$$\left[ \frac{1}{r^2} \frac{\partial}{\partial r} r^2 \frac{\partial z}{\partial r} + 4\pi K^2 G z^{\frac{1}{\gamma-1}} = 0 \right]$$

Lane-Emden equation

B.C.s: Solution finite at  $r=0$

$$\frac{\partial z}{\partial r} = 0 \text{ at } r=0$$

$$4\pi \int_0^\infty \rho r^2 dr = \text{Mass of star}$$

Analytic solutions for  $\frac{1}{\gamma-1} = n = \text{integer } 0, 1, 5$

More later in course

$$\frac{du}{dt} \equiv \frac{d^2x}{dt^2} = F/m$$

$$m = \rho \delta x \delta y \delta z$$

$$\frac{du}{dt} = \frac{\delta y \delta z}{\rho \delta x \delta y \delta z} (p(x) - p(x + \delta x))$$

$$\lim_{\delta x \rightarrow 0} \lim_{\delta y \rightarrow 0, \delta z \rightarrow 0} \frac{du}{dt} = - \frac{1}{\rho} \frac{\partial p}{\partial x}$$

Generalize:  $\underline{V} \equiv (u, v, w)$

$$\left[ \frac{d\underline{V}}{dt} = - \frac{1}{\rho} \nabla p \right]$$

Suppose Fluid element Subject to gravity:

$$\left[ \frac{d\underline{V}}{dt} = - \frac{1}{\rho} \nabla p - \nabla g z \right] \quad \begin{array}{l} \text{Euler's equation} \\ \text{No viscosity} \end{array}$$

Example: Incompressible Liquid at rest:

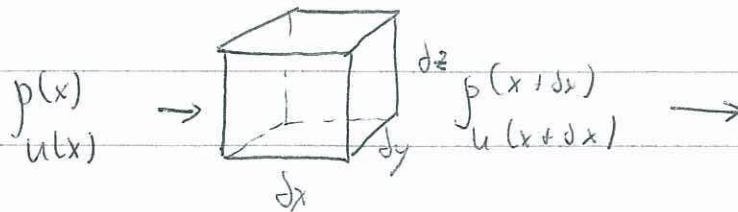
↑  
~~~~~  $z = 0$

Hydrostatic:  $-\frac{1}{\rho} \nabla p - \nabla g z = 0$       ↓  $\frac{\partial p}{\partial x} = \frac{\partial p}{\partial y} = 0$

$$\frac{dp}{dz} = -\rho g$$

$$\left[ p = p_a + \rho g d \right]$$

# Mass Continuity



$$\frac{dm}{dt} = \frac{d}{dt} \int \rho \, dx \, dy \, dz = \int dy \, dz \left( \rho(x) u(x) - \rho(x+dx) u(x+dx) \right)$$

$$\frac{\partial \rho}{\partial t} = - \frac{d}{dx} (\rho u)$$

$$\lim_{dx \rightarrow 0} : \frac{\partial \rho}{\partial t} = - \frac{\partial}{\partial x} (\rho u)$$

$$\text{Generalize: } \left[ \frac{\partial \rho}{\partial t} = - \nabla \cdot \rho \underline{\underline{v}}} \right] \quad \text{Mass Continuity}$$

$$\text{Incompressible Flow: } \left[ \nabla \cdot \underline{\underline{v}} = 0 \right]$$

Distinction between total and partial derivatives:  $A = A(x, y, z, t)$

Following a fluid element:  $\frac{d}{dt} \rightarrow d(A) = \left( \frac{\partial A}{\partial t} \right)_{x,y,z} dt$

$$+ \left( \frac{\partial A}{\partial x} \right)_{t,y,z} dx + \left( \frac{\partial A}{\partial y} \right)_{x,z,t} dy + \left( \frac{\partial A}{\partial z} \right)_{x,y,t} dz$$

$$\left[ \begin{aligned} \frac{dA}{dt} &= \frac{\partial A}{\partial t} + u \frac{\partial A}{\partial x} + v \frac{\partial A}{\partial y} + w \frac{\partial A}{\partial z} \\ &= \frac{\partial A}{\partial t} + \underline{\underline{v}} \cdot \nabla A \end{aligned} \right]$$

Euler's equation:

$$\left[ \frac{\partial \underline{V}}{\partial t} = - \underline{V} \cdot \nabla \underline{V} - \frac{1}{\rho} \nabla p - \nabla g z \right]$$

Continuity:

$$\frac{\partial \rho}{\partial t} = - \nabla \cdot \rho \underline{V}$$

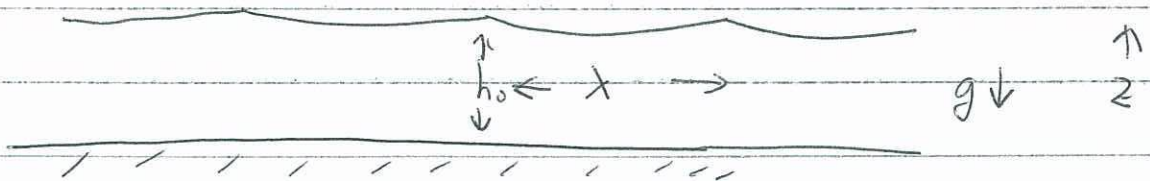
or  $\frac{d\rho}{dt} = \underline{V} \cdot \nabla \rho - \nabla \cdot \rho \underline{V}$

$$\left[ \frac{d\rho}{dt} = - \rho \nabla \cdot \underline{V} \right]$$

[Homework: Show that Euler:  $\left[ \frac{\partial \rho \underline{V}}{\partial t} + \nabla \cdot \rho \underline{V} \underline{V} = - \nabla p - \rho \nabla g z \right]$ ]

Examples of Behaviors of Euler Fluids:

1. Shallow water waves:



Mean depth  $h_0$

Local depth  $h$

Uniform gravity  $g$

Assume  $\lambda \gg h_0$



No breaking waves

Neglect surface tension



Assume hydrostatic balance (evaluate a posteriori)

Euler equations:

$$\left\{ \begin{array}{l} \frac{\partial u}{\partial t} + \underline{V} \cdot \nabla u = -\alpha_0 \frac{\partial p}{\partial x} \\ \frac{\partial v}{\partial t} + \underline{V} \cdot \nabla v = -\alpha_0 \frac{\partial p}{\partial y} \\ \alpha_0 \frac{\partial p}{\partial z} = -g \end{array} \right. \quad \alpha_0 \equiv \frac{1}{\rho_0} \text{ specific volume}$$

Note:  $\frac{\partial}{\partial x} \left( \alpha_0 \frac{\partial p}{\partial z} \right) = \frac{\partial}{\partial x} (-g) = 0 = \frac{\partial}{\partial z} \left( \alpha_0 \frac{\partial p}{\partial x} \right)$

$$\frac{\partial}{\partial z} \left( \alpha_0 \frac{\partial p}{\partial y} \right) = 0$$

$\frac{du}{dt}, \frac{dv}{dt}$  independent of  $z$  within fluid

If  $u, v$  independent of  $z$  initially, will always be independent of  $z$ .

Mass continuity: Define  $\vec{V} \equiv (u, v)$

$$\frac{\partial u}{\partial t} + \vec{V} \cdot \nabla u = -\alpha_0 \frac{\partial p}{\partial x}$$

$$\frac{\partial v}{\partial t} + \vec{V} \cdot \nabla v = -\alpha_0 \frac{\partial p}{\partial y}$$

Mass Continuity:  $\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} + \frac{\partial w}{\partial z} = 0$

$$\frac{\partial w}{\partial z} = -\frac{\partial u}{\partial x} - \frac{\partial v}{\partial y} = -\nabla \cdot \vec{V}$$

$$\int_0^h \frac{\partial w}{\partial z} = w_s = -\int_0^h \nabla \cdot \vec{V} = -h(\nabla \cdot \vec{V})$$

But  $w_s = \frac{dh}{dt}$

$$\left[ \frac{dh}{dt} = \frac{\partial h}{\partial t} + \vec{V} \cdot \nabla h = -h(\nabla \cdot \vec{V}) \right]$$

$$\left[ \frac{\partial h}{\partial t} = -\nabla \cdot h\vec{V} \right]$$

Pressure gradient terms

$$\alpha_0 \frac{\partial p}{\partial z} = -g$$

$$p = p_a + \frac{\rho}{\alpha_0} (h-z)$$

$$\alpha_0 \frac{\partial p}{\partial x} = \cancel{\alpha_0 \frac{\partial p_a}{\partial x}} + g \frac{\partial h}{\partial x}$$

$$\alpha_0 \frac{\partial p}{\partial y} = g \frac{\partial h}{\partial y}$$

Shallow water equations:

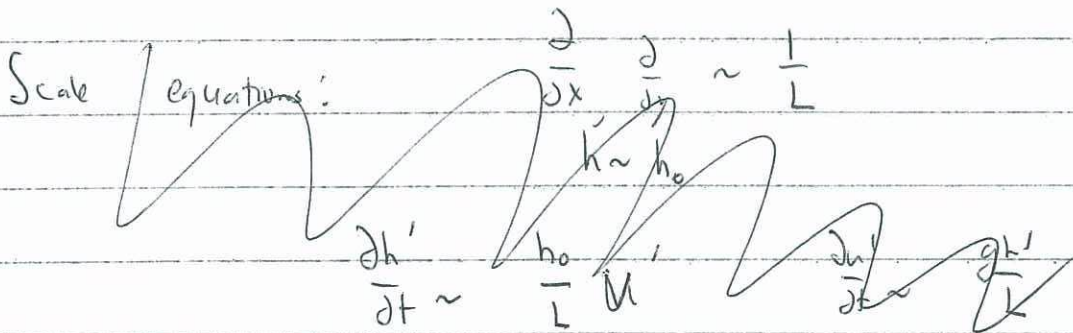
$$\left[ \begin{array}{l} \frac{\partial u}{\partial t} + \vec{V} \cdot \nabla u = -g \frac{\partial h}{\partial x} \\ \frac{\partial v}{\partial t} + \vec{V} \cdot \nabla v = -g \frac{\partial h}{\partial y} \\ \frac{\partial h}{\partial t} = -\nabla \cdot h \vec{V} \end{array} \right]$$

Linearize about resting state:

$$\vec{V} = 0 \quad h = h_0 \quad (\text{equilibrium solution})$$

$$\begin{aligned} \vec{V} &= 0 + \vec{V}' \\ h &= h_0 + h' \end{aligned}$$

$$\text{Assume } h' \ll h_0 \quad V' \ll ?$$



$$\frac{\partial u'}{\partial t} + \vec{V}' \cdot \nabla u' = -g \frac{\partial h'}{\partial x}$$

$$\frac{\partial v'}{\partial t} + \vec{V}' \cdot \nabla v' = -g \frac{\partial h'}{\partial y}$$

$$\left[ \frac{\partial h'}{\partial t} = -\vec{V}' \cdot \nabla h' - h_0 \nabla \cdot \vec{V}' - h' \nabla \cdot \vec{V}' \right]$$

Neglect quadratic terms, check after the fact:

$$\frac{\partial u'}{\partial t} = -g \frac{\partial h'}{\partial x}$$

$$\frac{\partial v'}{\partial t} = -g \frac{\partial h'}{\partial y}$$

$$\frac{\partial h'}{\partial t} = -h_0 \nabla \cdot \vec{V}' = -h_0 \left( \frac{\partial u'}{\partial x} + \frac{\partial v'}{\partial y} \right)$$

$$\Rightarrow \frac{\partial^2 h'}{\partial t^2} = -h_0 \left( \frac{\partial}{\partial x} \frac{\partial u'}{\partial t} + \frac{\partial}{\partial y} \frac{\partial v'}{\partial t} \right) = gh_0 \nabla^2 h'$$

$$\boxed{\frac{\partial^2 h'}{\partial t^2} - gh_0 \nabla^2 h' = 0}$$

One dimension:

$$\frac{\partial^2 h'}{\partial t^2} - gh_0 \frac{\partial^2 h'}{\partial x^2} = 0$$

$$h' = F(x - ct)$$

$$\left[ c^2 \equiv gh_0 \right]$$

Shallow water wave speed

Example:  $h_0 = 10 \text{ m}$   $c = 10 \text{ m s}^{-1}$



Two-dimensional water waves:

$$\frac{\partial^2 h'}{\partial t^2} - gh_0 \nabla^2 h' = 0$$

$$h' = e^{ik_x x + i k_y y + i \omega t}$$

$$c \equiv \sqrt{gh_0}$$

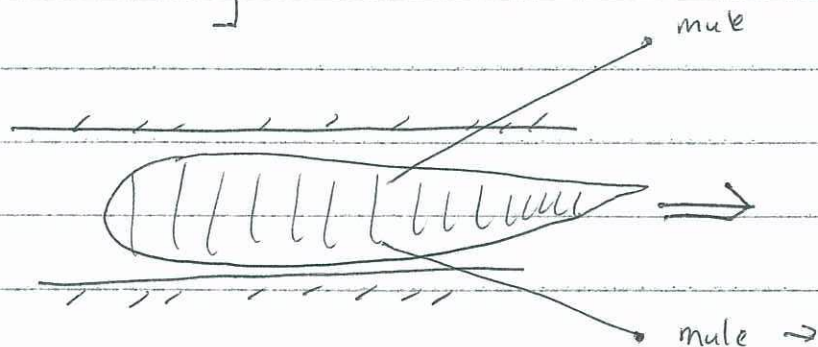
$$\omega^2 = gh_0 (k^2 + l^2)$$

$$C_{gx} = \frac{\partial \omega}{\partial k} = \frac{k}{\sqrt{k^2 + l^2}} c$$

$$C_{gy} = \frac{\partial \omega}{\partial l} = \frac{l}{\sqrt{k^2 + l^2}} c$$

$$\left[ |C_g| = \sqrt{C_{gx}^2 + C_{gy}^2} = c \right]$$

Ship in Canal:



If boat speed =  $\sqrt{gh_0}$ , waves cannot take energy away from ship  
no wave drag!

Is the solution of shallow water waves consistent with the hydrostatic approximation?

$$\left| \frac{dw}{dt} \right| \sim O\left(w \frac{\partial w}{\partial z}\right) \sim \frac{1}{h_0} \left(\frac{dh}{dt}\right)^2 \sim \frac{1}{h_0} \left(\frac{\partial h'}{\partial t}\right)^2 = h_0 \omega^2$$

$$= g h_0^2 (u^2 + v^2)$$

$$\left| \frac{dw}{dt} \right| \ll g$$

$$\frac{2}{h_0} \ll \frac{1}{u^2 + v^2} = \frac{\lambda^2}{2\pi}$$

$$\ll \left[ \lambda \gg \sqrt{2\pi} h_0 \right]$$

[See Insert!]

Steady Flow of an Euler Fluid:

$$\frac{d\vec{V}}{dt} = -\alpha_0 \nabla p - \nabla g z$$

Dot product with  $\vec{V}$ :

$$\frac{d}{dt} \left[ \frac{1}{2} \vec{V} \cdot \vec{V} \right] = -\alpha_0 \vec{V} \cdot \nabla p - \vec{V} \cdot \nabla g z$$

$$\frac{d}{dt} \left( \frac{1}{2} |\vec{V}|^2 \right) = -\alpha_0 \vec{V} \cdot \nabla p - g w$$

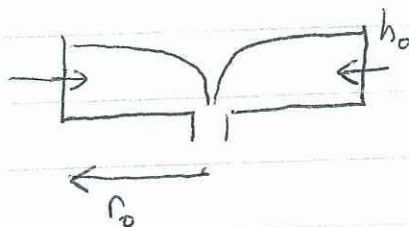
$$\left[ \frac{d}{dt} \left[ \frac{1}{2} |\vec{V}|^2 + \alpha_0 p + g z \right] = 0 \right]$$

Bernoulli:

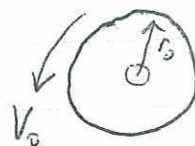
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## Bath tub Vortex

Idealized Set-up:



Rotating Screen



Assume flow is axisymmetric and steady

Euler's equations:

Tangential motion:

$$\frac{dm}{dt} = 0, \quad m \equiv rV = r_0 V_0$$

Radial motion:

$$\frac{du}{dt} = -g \frac{\partial h}{\partial r} + \frac{V^2}{r}$$

Approximation:  $\left| \frac{du}{dt} \right| \ll \frac{V^2}{r}, \quad g \frac{\partial h}{\partial r} \sim \frac{V^2}{r}$

$$g \frac{dh}{dr} = \frac{V^2}{r} = \frac{m^2}{r^3}$$

$$h = h_0 + \frac{1}{2} \frac{m^2}{g r_0^2} \left( 1 - r_0^2/r^2 \right)$$

or  $\frac{h}{h_0} = 1 + \chi (1 - 1/\chi^2)$

$$\chi \equiv r/r_0$$

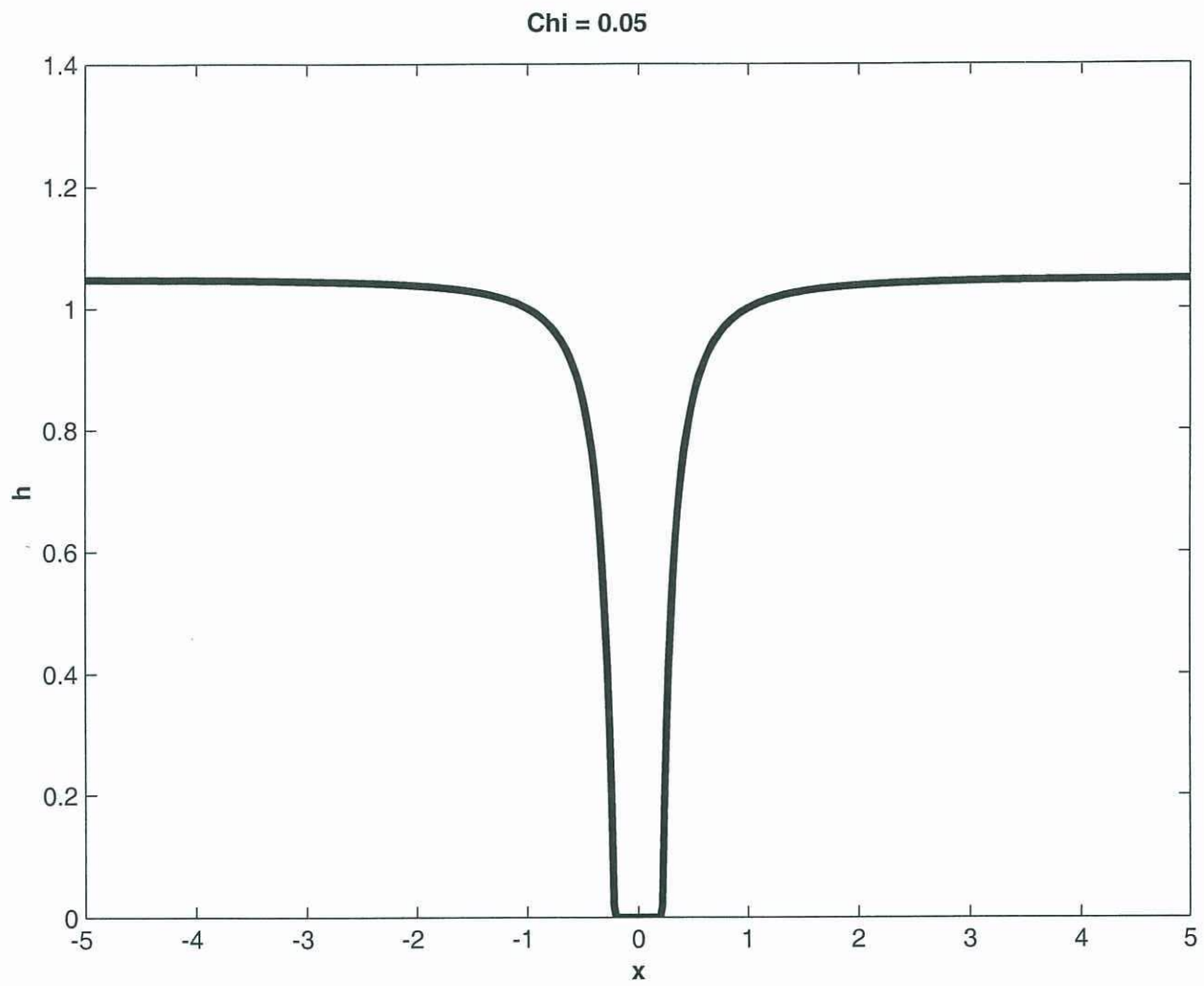
$$\chi \equiv \frac{1}{2} \frac{m^2}{g r_0^2 h_0}$$

$h=0$  when

$$\chi = \sqrt{\frac{\chi}{1+\chi}}$$

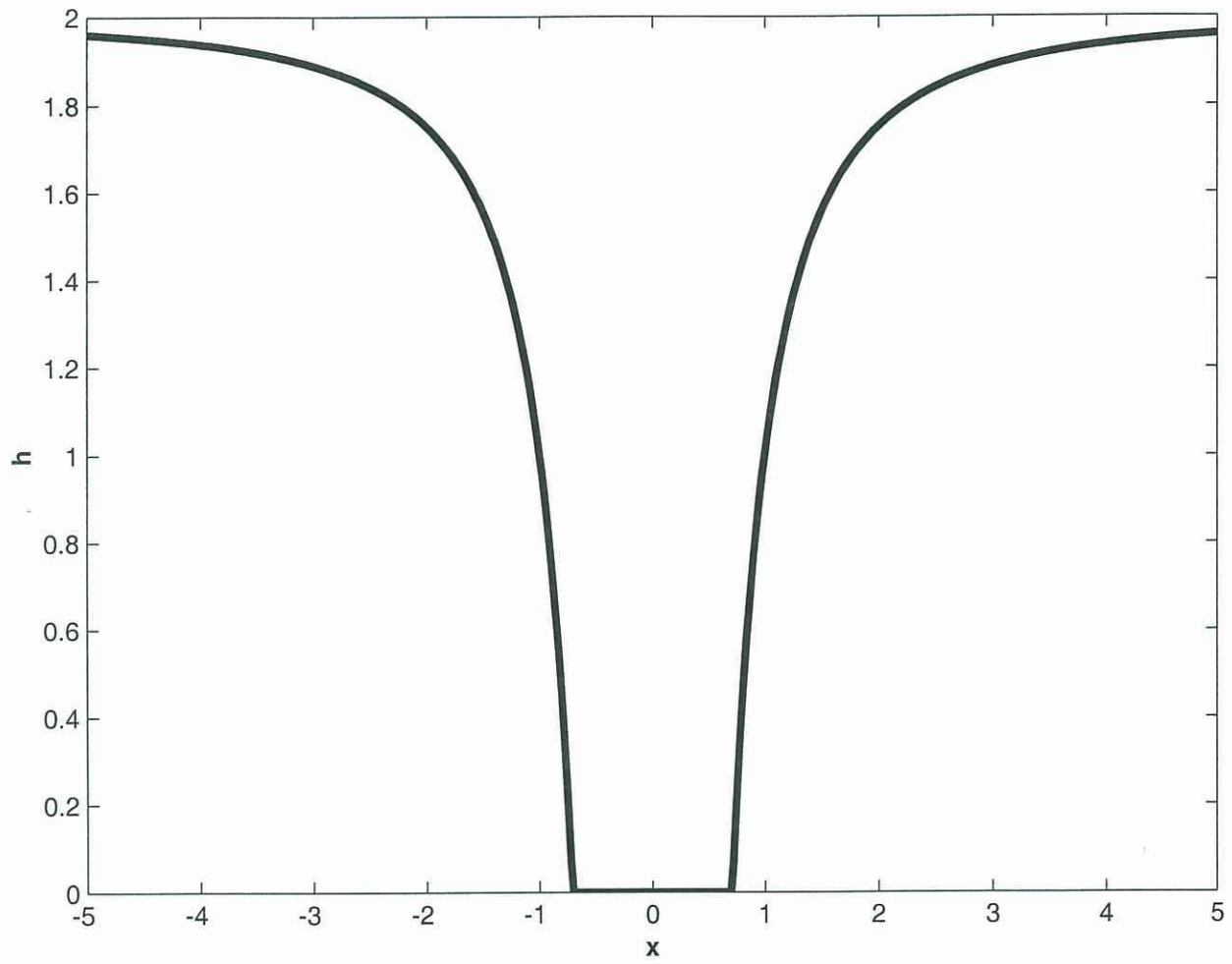
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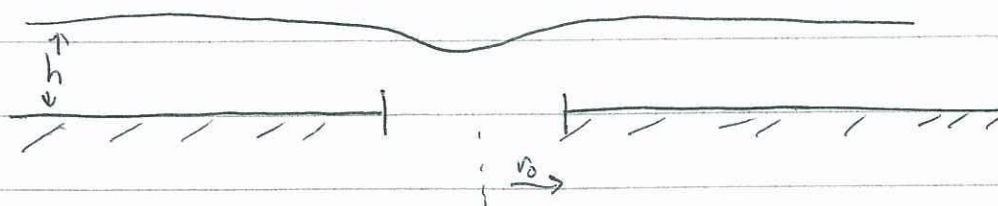


Chi = 1



Application:

Draining shallow water:



Steady: Mass conservation:  $r > r_0$ :  $\left[ 2\pi r h \rho u = Q_0 \right] = \text{mass flow rate}$

Bernoulli, along streamlines:

$$\frac{1}{2} u^2 + \alpha_0 p + g z = \text{Constant}$$

$$\alpha_0 p = \alpha_0 p_a + g(h - z)$$

$$\Rightarrow \frac{1}{2} u^2 + \alpha_0 p_a + g h = \text{Constant} = g h_0 + \alpha_0 p_a$$

$$\left[ \frac{1}{2} u^2 = g(h_0 - h) \right]$$

Eliminate  $u$ :

$$\frac{Q_0^2}{8\pi^2 r^2 \rho^2 g} = h^2(h_0 - h)$$

Cubic, in  $h$ .

Note:  $h < h_0$ , everywhere

Also note that right-hand side has a maximum value:

$$2h h_0 - 3h^2 = 0$$

$$\underline{h = 0 \text{ or } h = \frac{2}{3} h_0}$$

Maximum possible flow rate is

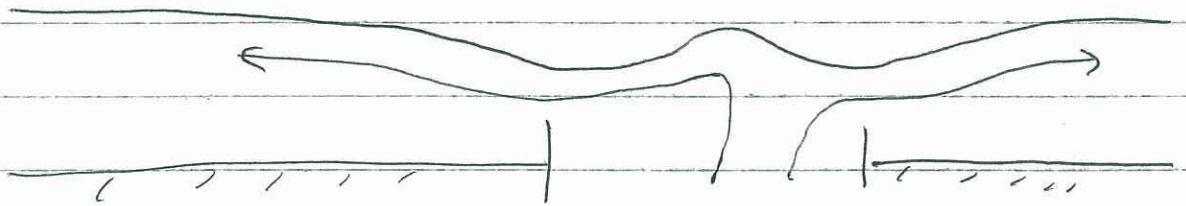
$$\left[ \frac{Q_0^2}{8\pi^2 r_0^3 \rho^2 g} = \frac{4}{27} h_0^3 \right] = \frac{1}{2} h^3 \Big|_{r_0}$$

At critical value of  $Q$

$$u = \frac{Q_0}{2\pi r_0 h \rho} = \frac{2\pi \rho r_0 \sqrt{gh^3}}{2\pi \rho r_0 h} = \sqrt{gh}$$

At large values of  $u$ , group velocity of waves cannot travel outward.  
What happens to this flow??

What happens when  $Q$  is negative (source)?



-16 1/2-

Re-visit bathtub vortex

This time relax cyclostrophic approximation

Along streamline:

$$\frac{1}{2} u^2 + \frac{1}{2} v^2 + \alpha_0 p_a + gh \approx \frac{1}{2} V_0^2 + \alpha_0 p_a + gh_0$$

$$\frac{1}{2} u^2 + \frac{1}{2} \frac{m^2}{r^2} + g(h-h_0) = \frac{1}{2} \frac{m^2}{r_0^2}$$

$$u = \frac{Q}{2\pi r h \rho}$$

Define  $Q' = \frac{Q_0}{\pi \rho r_0 \sqrt{gh_0^3 g}}$

$$\chi \equiv \frac{1}{2} \frac{m^2}{gh r_0^2}$$

$$\chi \equiv r/r_0$$

$$h' \equiv h/h_0$$

$$\left[ h'^3 + h'^2 \left( \chi \left( \frac{1}{\chi^2} - 1 \right) - 1 \right) + \frac{Q'^2}{\chi^2} = 0 \right]$$

Edge of hole at  $\chi_m (\chi \geq 1)$

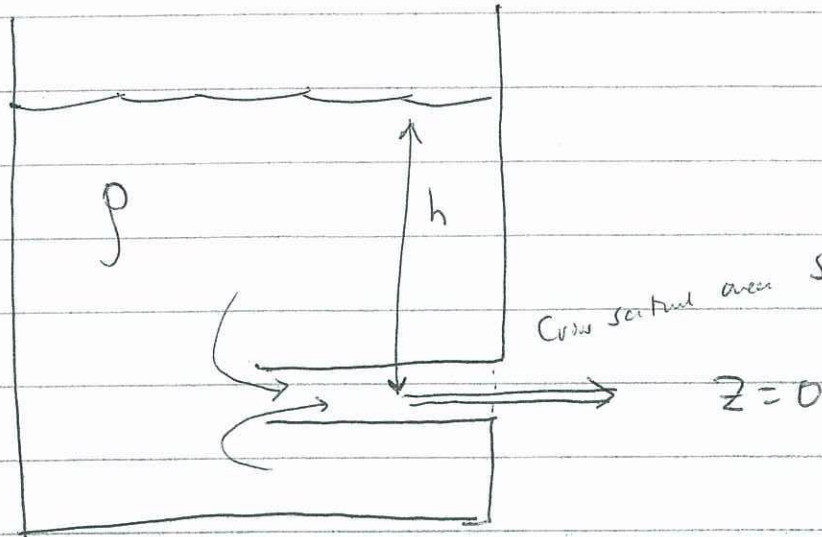
$$Q'^2_{max} = \frac{4}{27} \chi_m^2 \left( 1 - \chi \left( \frac{1}{\chi_m^2} - 1 \right) \right)^3$$

Decreases with  $\chi$ .  $Q' = 0$  when  $\chi = \frac{\chi_m^2}{1 - \chi_m^2}$

Small  $h'$ :  $h'^2 \approx \frac{Q'^2}{\chi^2 - \chi(1 - \chi_m^2)}$



Drawing Container of Euler Fluid:



Bernoulli:  $\frac{1}{2} u^2 + gz + p/\rho = p_a/\rho + gh$

At jet exit,  $p = p_a$ ,  $z = 0$ :  $\left[ \frac{1}{2} u^2 = gh \right]$  Same as free fall!

Let's calculate reaction force on vessel:

Export of horizontal momentum:  $u \cdot [\rho u S] = F = \rho S u^2 = \underline{\underline{2\rho S gh}}$

Let's calculate it from a different approach: pressure on walls of vessel:

~~Pressure on right and left sides cancels~~

Assume  $\underline{V} = 0$  everywhere on walls. Pressure same on right and left

side except opposite on face:

$$p = p_a + \cancel{\rho g (z-h)} + \rho g (h-z)$$

Uncompensated pressure force =  $\rho g h$

$\frac{1}{2}$  of other estimate: what gives??

Answer: jet emerges with smaller diameter than orifice  
 $\left[ S_j = \frac{1}{2} S \right]$

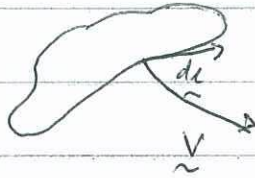
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## Irrotational (Potential) Flow

An irrotational flow is one for which  $\vec{\omega} \equiv \nabla \times \vec{V} = 0$

Under what circumstances will an irrotational flow remain irrotational?

Consider a material curve entirely within a fluid (in 3-D)



Each point on curve moves with flow at that point.

Define a Circulation,  $C$ :

$$C \equiv \oint \vec{V} \cdot d\vec{l}$$

$$\frac{d\vec{V}}{dt} = -\alpha \nabla p - \frac{g}{\rho} \nabla \rho$$

$$\oint \frac{d\vec{V}}{dt} \cdot d\vec{l} = \oint \frac{d}{dt} (\vec{V} \cdot d\vec{l}) - \oint \vec{V} \cdot \frac{d\vec{l}}{dt}$$

$$= \oint \frac{d}{dt} (\vec{V} \cdot d\vec{l}) = - \oint \vec{V} \cdot d\vec{V}$$

$$= \oint \frac{d}{dt} (\vec{V} \cdot d\vec{l}) - \oint \frac{1}{2} \frac{d}{dt} |\vec{V}|^2 > 0$$

$$= \frac{dC}{dt} = \oint -\alpha \nabla p \cdot d\vec{l} - \oint \nabla \phi \cdot d\vec{l} \Rightarrow 0$$

$$\frac{dc}{dt} = - \oint \alpha \nabla_p \cdot d\vec{l}$$

Special cases:

$$\alpha = \alpha_0$$

$$\alpha = \alpha(p) \text{ o.g., Constant entropy surface}$$

Then  $\frac{dc}{dt} = 0$  (Kelvin's Circulation Theorem)

$$C \equiv \oint \vec{V} \cdot d\vec{l} = (\text{by Stokes theorem}) \iint (\nabla \times \vec{V}) \cdot \hat{n} dA$$

If  $\vec{\eta} \equiv \nabla \times \vec{V} = 0$  initially everywhere, then for all time  $\vec{\eta} = 0$

On entropy surfaces or in incompressible flow. (Except at boundaries)

Since  $\nabla \times \vec{V} = 0$ ,  $\vec{V} = \nabla \Phi$   $\Phi = \text{"velocity potential"}$

$$\nabla \cdot \vec{V} = \nabla^2 \Phi = 0 \quad (\text{Euler fluid})$$

Since  $\vec{V} = \nabla \Phi$  is linear, velocity potentials may be superposed.  
(pressure, energy do not superpose linearly, in general)

Example: Flow owing to point source (or sink) of mass:

$$\nabla^2 \Phi = 0 \text{ except at point}$$



Green's Function solution:

$$3-D: \quad \Phi = -\frac{Q}{4\pi R}$$

 $Q =$  Source strength

$$V_r = \frac{\partial \Phi}{\partial R} = \frac{Q}{4\pi R^2}$$

Volume flow through sphere

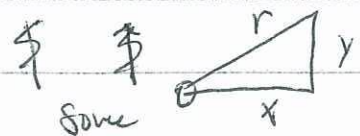
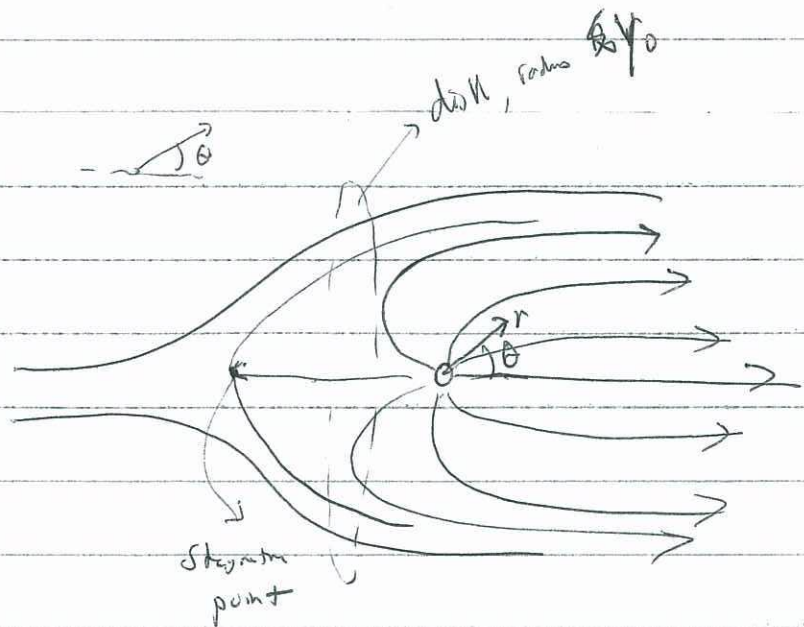
$$\text{Constant flow:} \quad \Phi = i\pi x$$

$$u = \frac{\partial \Phi}{\partial x} = U$$

Point source in constant flow:

$$V_r = U \cos \theta + \frac{Q}{4\pi r^2}$$

$$V_\theta = -U \sin \theta$$



Momentum flux change through disk:

$$u = U + \frac{Q}{4\pi R^2} \cos \theta = U + \frac{Qx}{4\pi R^3} = U + \frac{Qx}{4\pi (x^2 + y^2)^{3/2}}$$

~~Revised~~

~~Flux of disk~~

Flux through disk of radius  $y_0 = 2\pi \int_0^{y_0} ( ) y dy$

Momentum flux =  $2\pi \rho \int_0^{y_0} \left( u + \frac{Qx}{4\pi (x^2 + y^2)^{3/2}} \right)^2 y dy$

Difference between flux at  $\pm L =$

$$2\pi \rho \int_0^{y_0} \left[ \left( u + \frac{QL}{4\pi (L^2 + y^2)^{3/2}} \right)^2 - \left( u - \frac{QL}{4\pi (L^2 + y^2)^{3/2}} \right)^2 \right] y dy$$

$$= \rho u Q L \int_0^{y_0} \frac{y dy}{(L^2 + y^2)^{3/2}}$$

Let  $v = L^2 + y^2$

$dv = 2y dy$

~~2\pi~~  $\rho u Q L \frac{1}{2} \int_{L^2}^{L^2 + y_0^2} \frac{dv}{v^{3/2}} = -\rho u Q L \left( \frac{1}{\sqrt{L^2 + y_0^2}} - \frac{1}{L} \right)$

$y_0 \rightarrow \infty:$

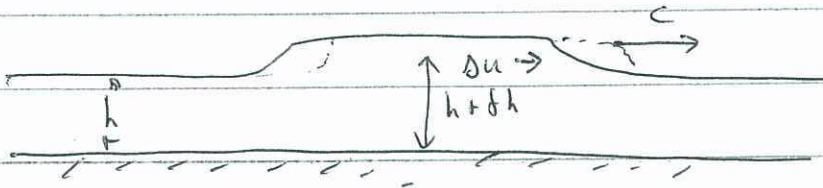
$$= \boxed{\rho u Q}$$


## Bore and Hydraulic Jumps

We consider shallow water waves of small amplitude at  $\lambda \gg h$ , waves are non-dispersive:  $c = \sqrt{gh}$  regardless of  $\lambda$ . Any arbitrary shape will propagate at this speed. But what happens in nonlinear regime?

This is complicated, but general tendency toward steepening.

Consider:



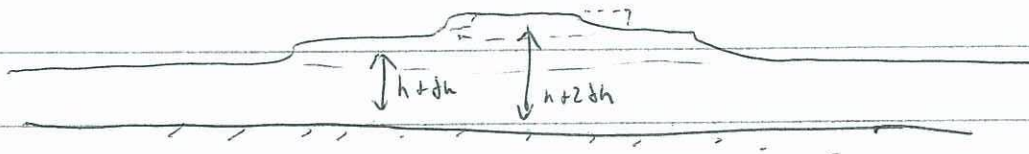
Rate of mass flow from left to right is  $c \delta h$

In between steps, fluid must be moving, so let

Mass conservation gives

$$\left[ \delta u (h + \delta h) = c \delta h \right]$$

Now consider two steps:

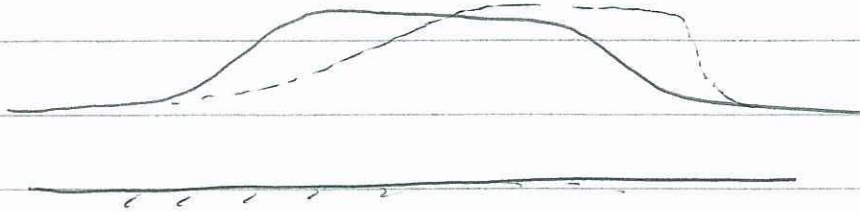


phase speed of upper wave is  $\delta u + \sqrt{g(h + \delta h)}$

Faster, because a.) fluid deeper, and b.) traveling on top of moving fluid



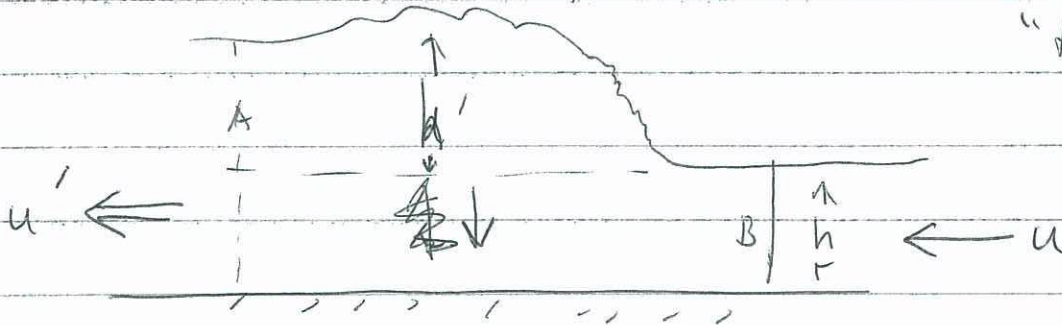
Any finite amplitude lump will steepen:



As wave front steepens, it develops Fourier Components with wavelengths  $\lambda$  not necessarily  $\gg h$ . It turns out that the phase speeds of these waves is  $\propto \sqrt{gh}$ , so they begin to propagate backward relative to the edge, showing up as undulations.

$$(K \propto \frac{1}{h}) \quad h' < 3h$$

Detailed analysis shows that if slope is  $\sim \frac{1}{3}$ , edge (with undulations) is stable. But if  $(h'/h) > \frac{1}{3}$ , edge continues to steepen until it breaks into a "foaming front".



Lots of dissipation; Bernoulli's theorem not applicable.

Mass continuity:  $[u'h' = uh]$



Momentum flux through B =  $\rho h u^2$   
 " " " A =  $\rho h' u'^2$

$$\frac{d}{dx} \left( \frac{\rho}{2} u^2 \right) = \frac{\rho}{2} \frac{du^2}{dx}$$

~~Leftward~~ <sup>Rightward</sup> Force on A:  $p = p_a + \rho g (h' - z)$

$$\text{Force} = \int_0^{h'} [p_a + \rho g (h' - z)] dz = h' p_a + \frac{1}{2} \rho g h'^2$$

$$\text{Force on B} = - \int_0^h [p_a + \rho g (h - z)] dz = - \left[ h p_a + \frac{1}{2} \rho g h^2 \right]$$

$$\text{Atmosphere force on wave front} = -p_a (h' - h)$$

$$\text{Sum} = \frac{1}{2} \rho g (h'^2 - h^2)$$

$$\text{Equate to momentum change: } \frac{1}{2} \rho g (h'^2 - h^2) = \rho h' u'^2 - \rho h u^2$$

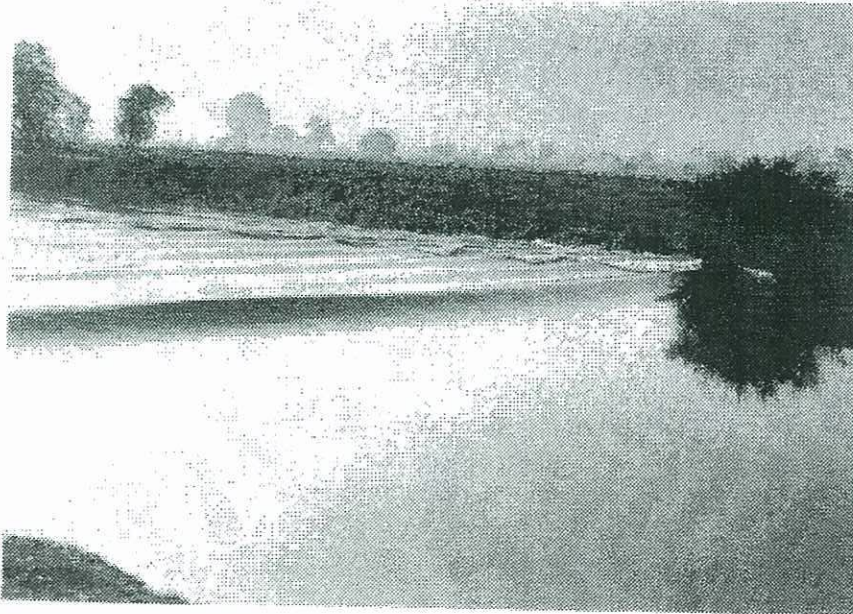
$$\text{Use } u' h' = u h$$

$$\left[ u^2 (= c^2) = \frac{g h'}{2 h} (h' + h) \right]$$

$$c^2 = g h \text{ when } h' = h$$

Curiously  $c^2$  ~~can be~~ <sup>is</sup> greater than  $g h'$ , if  $h' + h > 2h$   
 $\underline{h' > h}$

(see photos, pgs 72-73)





## Fluid Physics

Compressible Flows

Read 3.1-3.3

Ideal Gas Law:

(perfect gas)

$$p = \rho T \frac{k}{\bar{m}}$$

$k$  = Boltzmann constant,  $\bar{m}$  = mean molecular weight  
 $R \equiv k/\bar{m}$

First Law

$$Q = \frac{du}{dt} + \frac{dw}{dt}$$

 $du$  = change in internal energy $dw$  = work $Q$  = heating (per unit mass)

For reversible expansion in which pressure <sup>against</sup> ~~against~~ which expansion is  
 doing work = gas pressure

$$\frac{dw}{dt} = p \frac{d\alpha}{dt}$$

 $\alpha$  = specific volume

$$\frac{du}{dt} = C_v \frac{dT}{dt}$$

$$Q = C_v \frac{dT}{dt} + p \frac{d\alpha}{dt}$$

$$= C_v \frac{dT}{dt} + \frac{d}{dt}(p\alpha) - \alpha \frac{dp}{dt}$$

$$= (C_v + R) \frac{dT}{dt} - \alpha \frac{dp}{dt} \equiv C_p \frac{dT}{dt} - \alpha \frac{dp}{dt}$$

Sound

Consider first a 1-D sound wave:

Momentum:  $\frac{du}{dt} = -\alpha \frac{dp}{dx}$

Mass:  $\frac{1}{\alpha} \frac{d\alpha}{dt} = \nabla \cdot \underline{v} = \frac{du}{dx}$

Key assumption for sound waves: Expansion is reversible and adiabatic:

Let  $\alpha = \alpha(S, p)$   $S = \text{entropy}$ , invariant in rev. adiabatic expansion

$$d\alpha = \left( \frac{\partial \alpha}{\partial p} \right)_S dp, \quad \frac{1}{\alpha} \frac{d\alpha}{dt} = \frac{1}{\alpha} \left( \frac{\partial \alpha}{\partial p} \right)_S \frac{dp}{dt}$$

Let  $\beta \equiv -\frac{1}{\alpha} \left( \frac{\partial \alpha}{\partial p} \right)_S = \frac{1}{p} \left( \frac{\partial p}{\partial p} \right)_S$

Momentum:  $\frac{du}{dt} = -\alpha \frac{dp}{dx}$

Mass:  $\beta \frac{dp}{dt} = -\frac{du}{dx}$

$$\frac{d}{dt} = \frac{\partial}{\partial t} + u \frac{\partial}{\partial x}$$

Small amplitude waves: neglect nonlinear advection terms:

$$\frac{\partial u}{\partial t} = -\alpha \frac{\partial p}{\partial x}$$

$$\beta \frac{\partial p}{\partial t} = -\frac{\partial u}{\partial x}$$



Eliminate  $p$ :

$$\frac{\partial^2 u}{\partial t^2} - \frac{\bar{\alpha}}{\bar{\beta}} \frac{\partial^2 u}{\partial x^2} = 0$$

$$\left[ c^2 = \bar{\alpha} / \bar{\beta} \right]$$

Speed of sound

$$= \frac{\bar{\alpha}^2}{-(\partial \alpha / \partial p)_s} = \frac{1}{(\partial \rho / \partial p)_s} = \left( \frac{\partial p}{\partial \rho} \right)_s$$

Ideal gas:

$$\alpha = \frac{RT}{p} \quad \left( \frac{\partial \alpha}{\partial p} \right)_s = -\frac{\alpha}{p} + \frac{R}{p} \left( \frac{\partial T}{\partial p} \right)_s$$

$$C_p \frac{dT}{dt} - \alpha \frac{dp}{dt} = 0$$

$$\left( \frac{\partial T}{\partial p} \right)_s = \frac{\alpha}{C_p}$$

$$T ds = C_p dT - \alpha dp$$

$$= C_p \left[ \frac{dp}{pR} - \frac{p}{pR^2} dp \right] - \frac{dp}{p}$$

$$\Rightarrow \left( \frac{dp}{d\rho} \right)_s = \frac{C_p}{C_v} RT = c^2$$

$$-(\partial \alpha / \partial p)_s = \frac{\alpha}{p} \left[ 1 - \frac{R}{C_p} \right] = \frac{C_v}{C_p} \frac{\alpha}{p}$$

$$c^2 = \frac{\bar{\alpha}^2}{-(\partial \alpha / \partial p)_s} = \frac{C_p}{C_v} \frac{RT}{p} = \frac{C_p}{C_v} RT$$

$$c^2 = \frac{C_p}{C_v} RT$$

non-dispersive

Atmosphere:

$$C_p \approx 1000 \text{ J/Kg}$$

$$C_v \approx 713 \text{ J/Kg}^{-1}$$

$$R = 287 \text{ J/Kg}$$

$$c^2 \approx 400 \text{ J m}^2 \text{ s}^{-2}$$

$$c = 20 \sqrt{\text{J}} \text{ m s}^{-1}$$

$$T = 280 \text{ K}$$

$$c = 330 \text{ m s}^{-1}$$

Doppler shift:

$$c' = \bar{u} + c$$

$$= \omega + k\bar{u}$$

$$\omega' = c'k = ck + \bar{u}k$$

$$= k(c + \bar{u})$$

Bernoulli's equation for compressible Flow;

Momentum Equation:  $\frac{d\vec{V}}{dt} = -\alpha \nabla p - \nabla g z$

Steady

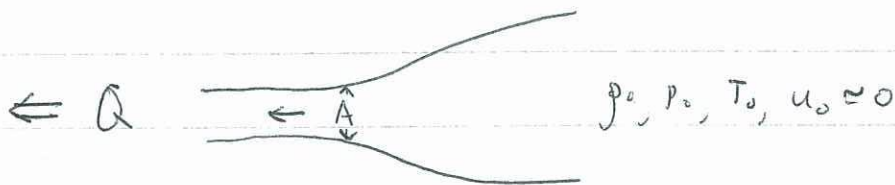
$\vec{V} \cdot \frac{d}{dt}$ :  $\frac{d}{dt} \left[ \frac{1}{2} |\vec{V}|^2 \right] = -\alpha \vec{V} \cdot \nabla p - g w$

$$= -\alpha \frac{dp}{dt} - g \frac{dz}{dt}$$

First Law, adiabatic ideal gas:  $\alpha \frac{dp}{dt} = c_p \frac{dT}{dt}$

$$\Rightarrow \frac{d}{dt} \left[ \frac{1}{2} |\vec{V}|^2 + \underbrace{c_p T}_{\text{enthalpy}} + g z \right] = 0$$

Forced flow through a Constriction:



$$Q = \rho u A$$

$$\frac{1}{2} u^2 + c_p T = c_p T_0 \quad \text{Bernoulli}$$

First law:  $c_p \frac{dT}{dt} = \alpha \frac{dp}{dt} = \frac{RT}{p} \frac{dp}{dt}$

$$c_p \frac{d \ln T}{dt} = R \frac{d \ln p}{dt}$$

$$\ln T/T_0 = R/c_p \ln p/p_0 \Rightarrow T = T_0 \left( p/p_0 \right)^{R/c_p}$$

$$\frac{1}{2} u^2 = c_p T_0 \left[ 1 - (P/P_0)^{R/c_p} \right] \quad (1)$$

$$Q = \dot{p} u A = \frac{p}{RT} u A = \frac{p u A}{R T_0 (P/P_0)^{R/c_p}}$$

$$= \frac{u A}{R T_0} p_0^{R/c_p} p^{1-R/c_p} = \frac{u A}{R T_0} p_0^{R/c_p} p^{c_v/c_p} \quad (2)$$

Eliminate  $p$  between (1) and (2):

$$Q = \frac{p_0}{R T_0} A u \left[ 1 - \frac{1}{2} \frac{u^2}{c_p T_0} \right]^{c_v/R}$$

Note that  $Q$  has a maximum value:

$$\frac{\partial Q}{\partial u} = 0 : \quad \text{when} \quad u^2 = \frac{2 c_p T_0}{1 + 2 c_v/R}$$

$$= \frac{2 c_v/R}{1 + 2 c_v/R} c_0^2$$

$$\left[ \text{Show that } u_{\max} = c_{\text{local}} \right]$$

$$T = T_0 - \frac{1}{2} u^2 / c_p$$

$$= T_0 \left[ 1 - \frac{1}{1 + 2 c_v/R} \right]$$

$$= 2 T_0 \frac{c_v/R}{1 + 2 c_v/R}$$

$$c^2 = \frac{c_p}{c_v} R T$$

$$= \frac{2 c_p T_0}{1 + 2 c_v/R} = u^2 \checkmark$$

At which  $Q = \dot{p}_0 A \left( \frac{2 c_p T_0}{1 + 2 c_v/R} \right)^{1/2} \left[ \frac{2 c_v/R}{1 + 2 c_v/R} \right]^{c_v/R}$

$$= \dot{p}_0 A c_0 \left[ \frac{2 c_v/R}{1 + 2 c_v/R} \right]^{\frac{1}{2} + c_v/R}$$

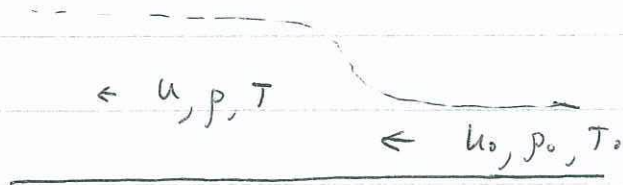
$$\text{Maximum Thrust} = \dot{p} A u^2 = Q u = \dot{p}_0 A c_0^2 \left[ \frac{2 c_v/R}{1 + 2 c_v/R} \right]^{1 + c_v/R}$$

$$= \frac{c_p}{c_v} \dot{p}_0 A \left[ \frac{2 c_v/R}{1 + 2 c_v/R} \right]^{c_p/R}$$

Limits speed of jet airplanes

## One-Dimensional Shocks:

Coordinate system moving with shock:



Energy Conservation (no dependence on viscosity):

$$\frac{1}{2} u^2 + c_p T = \frac{1}{2} u_0^2 + c_p T_0$$

Mass Conservation:  $\rho_0 u_0 = \rho u$ Overall momentum Conservation:  $\rho u^2 - \rho_0 u_0^2 = p_0 - p$ Ideal gas law:  $p = \rho R T$ Eliminate  $u, p, T$ :

$$u_0^2 = c_0^2 \left[ 1 + \frac{p - p_0}{p_0} \frac{c_p + c_v}{2c_p} \right]$$

gives propagation speed,  $u_0$ , as function of  $(p - p_0)/p_0$ 

$$u^2 = u_0^2 \left[ 1 - \frac{c_v}{c_p} \frac{c_0^2}{u_0^2} \frac{p - p_0}{p_0} \right]$$

$$= u_0 \left[ 1 - \frac{c_v}{c_p} \frac{\sigma}{1 + \sigma \frac{c_p + c_v}{2c_p}} \right] \quad \sigma \equiv \frac{p - p_0}{p_0}$$

$$\int_{\text{From } \frac{d}{dx} \frac{1}{\rho u}} = \frac{1}{\rho_0} \frac{dp}{dx}$$

$$\downarrow \quad \rho u \frac{du}{dx} = \frac{dp}{dx}$$

$$\Rightarrow \frac{d}{dx} (\rho u^2) = \frac{dp}{dx}$$

Since  $\rho u = \text{constant}$



# Some applications of Compressible Flow theory:

## 1. Hydrostatic equilibrium:

$$\frac{\partial p}{\partial z} = -\rho g = -\frac{p}{RT} g$$

### a. Isothermal atmosphere: $T = \text{constant} = T_0$

$$p = p_0 e^{-gz/RT_0} \equiv p_0 e^{-z/H}$$

$$H = \text{temperature scale height} \equiv RT_0/g$$

### b. Isentropic atmosphere:

$$T = T_0 \left( \frac{p}{p_0} \right)^{R/c_p}$$

$$\frac{\partial p}{\partial z} = -\frac{g}{RT_0} p_0^{R/c_p} p^{c_v/c_p}$$

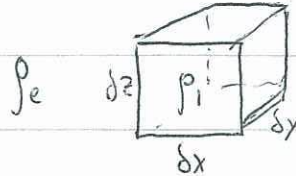
$$p = p_0 - \frac{c_p}{c_v c_p} \frac{g}{RT_0} p_0^{R/c_p} \left[ \frac{c_v + c_p}{c_p} p \right]$$

$$p = p_0 \left[ 1 - \frac{gz}{c_p T_0} \right]^{R/c_p}$$

## Fluid Physics

Convection Stability

Basic idea of stability:



$$\text{Pressure force on box} = p_0 \delta x \delta y - (p_0 + \delta p) \delta x \delta y$$

$$\frac{\partial p}{\partial z} = -\rho_e g \quad \delta p = -\rho_e g \delta z$$

$$\text{Pressure force} = \rho_e g \delta z \delta x \delta y$$

$$\text{Weight} = -\rho_i g \delta z \delta x \delta y$$

$$\text{Sum} = (\rho_e - \rho_i) g \delta x \delta y \delta z$$

= Difference between weight of box & weight of fluid it displaces

Archimedes

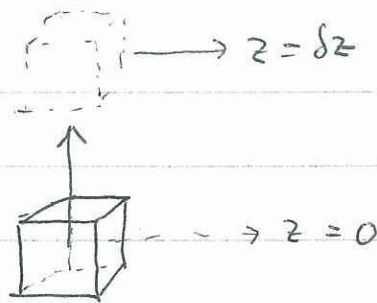
Valid for any shape

Convective stability:

$$M \frac{dw}{dt} = \rho_i \delta x \delta y \delta z \frac{dw}{dt} = (\rho_e - \rho_i) g \delta x \delta y \delta z$$

$$\frac{dw}{dt} = g \frac{\rho_e - \rho_i}{\rho_i} = g \frac{\alpha_i - \alpha_e}{\alpha_e}$$

Test sample:



Reversible, adiabatic displacement: Will  $\alpha_i > \alpha_e$ ?

§ Let  $\alpha = \alpha(s, p)$

$$\left[ (\delta\alpha)_p = \left( \frac{\partial\alpha}{\partial s} \right)_p \delta s \right]$$

Division: Derive one of Maxwell's relations:

First define a quantity called enthalpy:

$$H \equiv u + p\alpha$$

$$\left[ \delta H = \delta u + p \delta\alpha + \alpha dp \right]$$

Definition of entropy,  $s$ :

$$T \delta s = \delta u + p \delta\alpha = \delta H - \alpha dp$$

Now:  $\left( \frac{\partial H}{\partial s} \right)_p = T$

$$\left( \frac{\partial H}{\partial p} \right)_s = \alpha$$

Cross differentiation: :

$$\left[ \left( \frac{\partial \alpha}{\partial s} \right)_p = \left( \frac{\partial T}{\partial p} \right)_s \right] \quad \text{Maxwell}$$

$$\text{Thus } (\partial \alpha)_p = \left( \frac{\partial T}{\partial p} \right)_s ds$$

$$\text{Thus } \frac{dw}{dt} = \frac{g}{\alpha_e} (\alpha_i - \alpha_e) = \frac{g}{\alpha_e} \left( \frac{\partial T}{\partial p} \right)_s (s_i - s_e)$$

~~Given~~ Assuming that  $\left( \frac{\partial T}{\partial p} \right)_s > 0$ ,

acceleration will be upward if  $s_i > s_e$ , i.e. if  $s$  decreases upward.

Hydrostatic:  $\alpha dp = -g dz$

$$\frac{dw}{dt} = - \left( \frac{\partial T}{\partial z} \right)_s (s_i - s_e) \equiv \Gamma (s_i - s_e)$$

$$\Gamma \equiv - \left( \frac{\partial T}{\partial z} \right)_s = \text{adiabatic lapse rate}$$

$$s_i - s_e = - \left( \frac{\partial s}{\partial z} \right) dz$$

$$\frac{dw}{dt} = \frac{d^2 dz}{dt^2} = - \Gamma \frac{\partial s}{\partial z} dz$$



If  $\frac{\partial s}{\partial z} > 0$ , oscillator:

$$\left[ \omega^2 = \rho \frac{\partial s}{\partial z} \right]$$

Example: Ideal gas:

$$c_p dT - \alpha dp = 0 \quad \text{adiabatic}$$

$$c_p dT + g dz = 0 \quad \text{hydrostatic}$$

$$-\left( \frac{\partial T}{\partial z} \right)_s = \Gamma = g/c_p$$

$$\left[ \omega^2 = \frac{g}{c_p} \frac{\partial s}{\partial z} \right]$$

If  $\frac{\partial s}{\partial z} < 0$  unstable:  $\delta z = A e^{\sigma t}$

$$\sigma^2 = \cancel{\frac{g}{c_p} \left( \frac{\partial s}{\partial z} \right)} \quad \Gamma \left( -\frac{\partial s}{\partial z} \right)$$

Self-organized Criticality:  $\frac{\partial s}{\partial z} \approx 0$  isentropic state.

# Stellar Physics:

Approximations:

Self-gravitating sphere

neglect rotation

Hydrostatic equilibrium

Mass in spherical shell of thickness  $dr$ :

$$dm = 4\pi r^2 \rho dr$$

$$\Rightarrow \frac{dr}{dm} = \frac{1}{4\pi \rho r^2}$$

$$\left[ \frac{d}{dm} = \frac{1}{4\pi \rho r^2} \frac{d}{dr} \right]$$

Gravitational acceleration:

$$g = \frac{Gm}{r^2}$$

$G$  = gravitational constant

$m$  = mass within radius  $r$

Gravitational potential  $\phi$  satisfies:

$$\nabla^2 \phi = 4\pi G \rho$$

$$\Rightarrow \frac{1}{r^2} \frac{d}{dr} \left( r^2 \frac{d\phi}{dr} \right) = 4\pi G \rho$$

$$g = \frac{d\phi}{dr} = \frac{Gm}{r^2}$$

$$\phi = \int_0^r \frac{Gm}{r^2} dr + \text{constant}$$

Hydrostatic Equilibrium:

$$\frac{\partial p}{\partial r} = -\rho g = -\rho \frac{Gm}{r^2} = -\rho \frac{\partial \phi}{\partial r}$$

In mass coordinates:

$$\left[ \frac{\partial p}{\partial m} = -\frac{Gm}{4\pi r^2} \right]$$

Can't solve unless we ~~for~~ have  $\rho = \rho(p)$

or Unless we know  $\rho = \rho(p, \tau)$   $T = T(p \text{ or } \rho)$

Hydrostatic equilibrium in general relativity:

$$\frac{\partial p}{\partial r} = -\rho \frac{Gm}{r^2} \left(1 + \frac{p}{\rho c^2}\right) \left(1 + \frac{4\pi r^3 \rho}{mc^2}\right) \left(1 - \frac{2Gm}{rc^2}\right)^{-1} \quad \text{Ideal gas}$$

Polytropic spheres:

$$\left[ \rho = K_1 p^{1/\gamma} \right] \quad \text{or} \quad p = K \rho^\gamma$$

Example:  $\gamma \rightarrow \infty$ :  $p = K$

Hydrostatic:

$$\frac{\partial \phi}{\partial r} = -\frac{1}{\rho} \frac{\partial p}{\partial r} = -\gamma K \rho^{\gamma-2} \frac{\partial \rho}{\partial r} = -\frac{\gamma}{\gamma-1} K \frac{\partial}{\partial r} [\rho^{\gamma-1}]$$

Unless  $\gamma=1$ , can be integrated, to get (w.m.  $p=0$  at  $\phi=0$ )

$$\rho = \left[ \frac{-\phi}{K(\gamma-1)} \right]^n \quad \text{w.m. } n \equiv \frac{1}{\gamma-1} \quad (\phi < 0)$$

Now we have Poisson equation:

$$\left[ \frac{1}{r^2} \frac{d}{dr} \left( r^2 \frac{d\phi}{dr} \right) = 4\pi \rho G = 4\pi G \left[ \frac{-\phi}{K(1+n)} \right]^n \right]$$

Scale Variables:

$$r \rightarrow z/A$$

$$\phi \rightarrow \phi_c w$$

$$A^2 \equiv \frac{4\pi G}{(n+1)^n K^n} (-\phi_c)^{n-1}$$

$\phi_c \equiv \phi$  at center of star

$$\rightarrow \left[ \frac{1}{z^2} \frac{d}{dz} \left( z^2 \frac{dw}{dz} \right) + w^n = 0 \right]$$

Lane-Emden equation

Interested in solution that are finite at  $z=0$

Also means that  $\frac{dw}{dz} = 0$  at  $z=0$

By definition  $w=1$  at  $z=0$

Also remember that  $\rho = \left[ \frac{-\phi}{K(1+n)} \right]^n$

$$w \equiv \frac{\phi}{\phi_c} = (\rho/\rho_c)^{1/n}$$

So  $\rho = \rho_c w^n$

$$\rho_c = \left[ \frac{-\phi_c}{(n+1)K} \right]^n$$



Also,  $p = k p^8 = k p^{\frac{1+n}{n}}$

$$\left[ p = p_c w^{n+1} \right] \quad p_c = k p_c^8$$

Lane-Emden equation has regular singularity at  $z=0$

Express  $w$  as a power series:

$$w(z) = 1 + a_1 z + a_2 z^2 + a_3 z^3 + \dots$$

$$\left( \frac{dw}{dz} \right)_0 = 0: \quad a_1 = 0$$

Insert power series in Lane-Emden equation, match coefficients for like powers:

$$w(z) = 1 - \frac{1}{6} z^2 + \frac{n}{120} z^4 + \dots$$

Has maximum value at  $z=0$

~~Analytic~~  $S$

Analytic solutions only for certain values of  $n$ !

$$n=0: \quad w(z) = 1 - \frac{1}{6} z^2$$

$$n=1 \quad w(z) = \frac{\sin z}{z}$$

$$n=5: \quad w(z) = \frac{1}{(1 + z^2/3)^{1/2}}$$

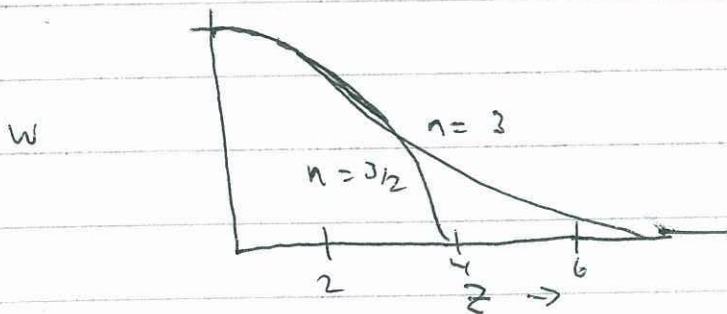
Surface of polytrope of index  $n$  defined by value of  $z, z_n$ , for which  $p = w = 0$ .

For  $n = 0, 1$   $z_n$  finite

$n = 5$   $z_n$  infinite

Can be shown that  $z_n$  finite for  $n < 5$

General (numerical) solutions look like



| $n$ | $z_n$    | $p_c / \bar{p}$ |
|-----|----------|-----------------|
| 0   | 2.45     | 1               |
| 1   | $\pi$    | 3.29            |
| 2   | 4.35     | 11.40           |
| 3   | 6.90     | 54.18           |
| 4.5 | 31.84    | 6189            |
| 5   | $\infty$ | $\infty$        |

Total Mass of Star:

$$M = \int_0^{r_n} 4\pi \rho r^2 dr$$

But  $\rho = \rho_c w^n$

$$r \rightarrow z/A$$

$$r_n \equiv R = z_n/A$$

$$M = \frac{4\pi \rho_c}{A^3} \int_0^{z_n} w^n z^2 dz$$

But from Lane-Emden equation,  $w^n z^2 = - \frac{d}{dz} \left( z^2 \frac{dw}{dz} \right)$

$$M = \frac{4\pi}{A^3} \rho_c \left[ -z^2 \frac{dw}{dz} \right]_{z_n} \quad A = z_n/R$$

$$M = 4\pi R^3 \rho_c \left( \frac{1}{z_n} \frac{dw}{dz} \right) \Big|_{z_n}$$

Determines  $\rho_c$  given  $M$  and  $R$

But  $A^2 = \frac{z_n^2}{R^2} = \frac{4\pi G}{(n+1)K} (-\rho_c)^{n-1}$  and  $\rho_c = \left[ \frac{-\rho_c}{K(n+1)} \right]^n$

So  $\frac{z_n^2}{R^2} = \frac{4\pi G}{(n+1)K} \rho_c^{\frac{n-1}{n}}$

Then

$$M = 4\pi R^{\frac{n-3}{n-1}} z_n^{\frac{n+1}{n-1}} \left[ \frac{(n+1)K}{4\pi b} \right]^{\frac{n}{n-1}} \frac{dw}{dz} \Big|_{z_n}$$

This determines  $M$  as function of  $R$  or

$$R \sim M^{\frac{n-1}{n-3}}$$

$$p_c = \left\{ z_n^4 \left[ \frac{n(n+1)}{4\pi b} \right]^3 \left[ \frac{m}{4\pi} \frac{dw}{dz} \Big|_{z_n} \right]^{-2} \right\}^{\frac{n}{n-3}}$$

Some interesting cases:

White dwarfs of small mass: Degenerate electron gas

$$\underline{n = 3/2}$$

$$R \sim M^{-1/3}$$

$$p_c \sim m^2$$

Oddly enough,  $R$  decreases with increasing  $M$ .

Also, density increases. At some point, interior becomes relativistic

Here  $n = 3$  in limit of extreme relativity

We can try to solve for relativistic interior at match to non-relativistic exterior



Simplest model: Fit interior with  $n=3$  to exterior with  $n=3/2$

Chandrasekhar's first model

In extreme case, all of star is relativistic

Then  $M$  independent of  $R$ :

$$M = 4\pi Z_3^2 \left[ \frac{K}{\pi G} \right]^{3/2} \frac{dw}{dz} \Big|_{3/2}$$

Chandrasekhar mass

Limiting mass of star composed of  
degenerate, relativistic electron gas

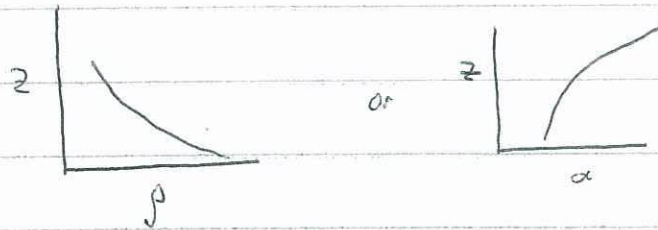
As central density increases without limit, mass approaches this  
finite limit as  $R \rightarrow 0$ . But before this limit is reached,  
properties of gas change again

Some stars have  $M > M_{ch}$ , but these are not composed of degenerate  
electron gas

Whole class of ~~star~~ degenerate, relativistic degenerate electron stars with  $M = M_{ch}$

## Stratified Flows:

Consider a fluid at rest in a gravitational ~~field~~ <sup>field</sup>, where density varies with altitude:



Linearize equations of motion about this state:

$$\frac{du}{dt} = \frac{\partial u}{\partial t} + \underline{V} \cdot \nabla u = -\alpha \nabla_{\underline{p}} \frac{\partial p}{\partial x}$$

$$\frac{\partial u}{\partial t} + \underline{V} \cdot \nabla u = -\alpha \frac{\partial p}{\partial x}$$

$$\frac{1}{\alpha} \left[ \frac{\partial \alpha}{\partial t} + \underline{V} \cdot \nabla \alpha \right] = \nabla \cdot \underline{V}$$

$$\frac{\partial w}{\partial t} + \underline{V} \cdot \nabla w = -\alpha \frac{\partial p}{\partial z} - g$$

$$\frac{\partial s}{\partial t} + \underline{V} \cdot \nabla s = 0$$

$$\Rightarrow \frac{\partial u'}{\partial t} = -\bar{\alpha}(z) \frac{\partial p'}{\partial x}$$

$$\frac{\partial v'}{\partial t} = -\bar{\alpha}(z) \frac{\partial p'}{\partial y}$$

$$\frac{\partial w'}{\partial t} = -\bar{\alpha}(z) \frac{\partial p'}{\partial z} - \frac{\partial \bar{p}}{\partial z} \alpha'$$

$$= -\bar{\alpha}(z) \frac{\partial p'}{\partial z} + g \frac{\alpha'}{\bar{\alpha}}$$

$$\frac{\partial \alpha'}{\partial t} = \bar{\alpha} \nabla \cdot \underline{V}'$$

$$\frac{\partial s'}{\partial t} + w' \frac{\partial \bar{s}}{\partial z} = 0$$

~~$\frac{dw}{dz}$~~  Remember that  $g \frac{\alpha'}{\alpha} = - \int \frac{d\bar{s}}{dz} dz$

$$\left[ \frac{dw'}{dt} = -\bar{\alpha}(z) \frac{dp'}{dz} - \int \frac{d\bar{s}}{dz} dz \right]$$

Also,  $\frac{1}{\alpha} \frac{d\alpha'}{dt} = \nabla \cdot \mathbf{V}'$

$$\frac{g}{\alpha} \frac{d\alpha'}{dt} = g \nabla \cdot \mathbf{V}'$$

$$\left[ - \int \frac{d\bar{s}}{dz} w' = g \nabla \cdot \mathbf{V}' \right]$$

Two further approximations to the equations:

Continuity:  $-\frac{1}{g} \int \frac{d\bar{s}}{dz} w' = \frac{dw'}{dx} + \frac{dw'}{dy} + \frac{dw'}{dz}$

Scale equation:  $\frac{dw'}{dz} \sim \frac{w'}{H}$  compare to  $-\frac{1}{g} \int \frac{d\bar{s}}{dz} w'$

Latter negligible if  $\frac{1}{H} \gg \frac{1}{g} \int \frac{d\bar{s}}{dz}$

$$\left[ H \ll H_0 \equiv \frac{1}{\int \frac{d\bar{s}}{dz}} \right]$$

For example, in earth's troposphere  $H_0 \approx 100 \text{ km}$

Second approximation: If  $H \ll H_T \equiv \frac{RT}{g}$  then  $\alpha(z) \approx \alpha_0$  in equations

$$\left\{ \begin{array}{l} \frac{\partial u'}{\partial t} = -\alpha_0 \frac{\partial p'}{\partial x} \quad (1) \\ \frac{\partial v'}{\partial t} = -\alpha_0 \frac{\partial p'}{\partial y} \quad (2) \\ \frac{\partial w'}{\partial t} = -\alpha_0 \frac{\partial p'}{\partial z} - \int \frac{\partial \bar{s}}{\partial z} \delta z' \quad (3) \\ \frac{\partial}{\partial t} \delta z' = w' \quad (4) \\ \nabla \cdot V' = 0 \quad (5) \end{array} \right.$$

Eliminate variables in favor of  $w'$ :

$$-\frac{\partial}{\partial x} (1) + \frac{\partial}{\partial y} (2): \quad = -\frac{\partial}{\partial t} \left( \frac{\partial u'}{\partial x} + \frac{\partial v'}{\partial y} \right) = \frac{\partial}{\partial t} \frac{\partial w'}{\partial z} = \alpha_0 \left( \frac{\partial^2 p'}{\partial x^2} + \frac{\partial^2 p'}{\partial y^2} \right)$$

$$\equiv \alpha_0 \nabla_2^2 p'$$

$$\left[ \frac{\partial}{\partial t} \frac{\partial^2 w'}{\partial z^2} = \alpha_0 \nabla_2^2 \frac{\partial p'}{\partial z} \right]$$

$$\text{Take } \nabla_2^2 (3): \quad \left[ \nabla_2^2 \frac{\partial w'}{\partial t} = -\alpha_0 \nabla_2^2 \frac{\partial p'}{\partial z} - \int \frac{\partial \bar{s}}{\partial z} \nabla_2^2 \delta z' \right]$$

$$\text{Eliminate } p': \quad \nabla^2 \frac{\partial w'}{\partial t} = - \int \frac{\partial \bar{s}}{\partial z} \nabla_2^2 \delta z'$$

$$\text{Take } \frac{\partial}{\partial t}: \quad \left[ \nabla^2 \frac{\partial^2 w'}{\partial t^2} + \underbrace{\int \frac{\partial \bar{s}}{\partial z}}_{\equiv N^2} \nabla_2^2 w' = 0 \right]$$

3-D wave equation



Constant coefficients. Look for solutions of form

$$W' \sim W_0 e^{iK_x x + iK_y y + i r z - i \omega t}$$

$$K^2 \equiv K_x^2 + K_y^2$$

$$\left[ \frac{\omega^2}{N^2} = K^2 \right]$$

$$\left[ \omega = \pm K N \sqrt{1 + r^2} \right]$$

Phase of wave:  $iK_x x + iK_y y + i r z - i \omega t = \text{constant}$

$$C_x \equiv \left( \frac{\partial x}{\partial t} \right)_{\text{phase}, y, z} = \omega / K_x = \pm \frac{K}{K_x} \frac{N}{\sqrt{1 + r^2}}$$

$$C_y \equiv \left( \frac{\partial y}{\partial t} \right)_{\text{phase}, x, z} = \omega / K_y = \pm \frac{K}{K_y} \frac{N}{\sqrt{1 + r^2}}$$

~~$$K_z \equiv \left( \frac{\partial z}{\partial t} \right)_{\text{phase}, x, y} = \omega / r = \pm \frac{K}{r} \frac{N}{\sqrt{1 + r^2}}$$~~

$$C_z \equiv \left( \frac{\partial z}{\partial t} \right)_{\text{phase}, x, y} = \omega / r = \pm \frac{K}{r} \frac{N}{\sqrt{1 + r^2}}$$

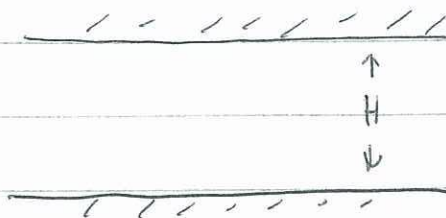
Consider plane wave ~~travelling in x direction~~ along y axis ( $K_y = 0$ )  
for which  $\lambda \gg H$  ( $K_x \ll r$ ):

$$C_x = \pm \frac{N}{r} \quad (\text{independent of } K_x)$$

$$\left[ C_z = \pm \frac{k_x}{r} \frac{N}{r} \right]$$

Non-dispersive in x direction

Example: Flow between plates:



$$r = \pi/H$$

$$\left[ C_x = \pm \frac{NH}{\pi} \right]$$

non dispersive, like shallow water waves

Group velocities:

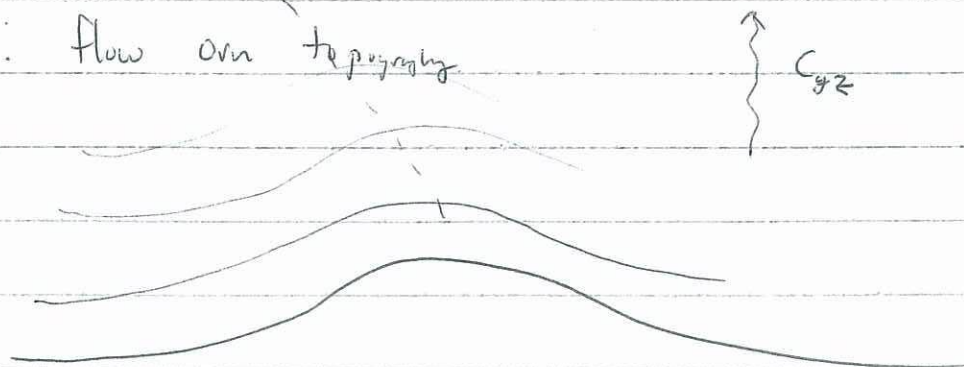
Hydrostatic limit:  $\omega = \pm \frac{KN}{r}$

$$C_{gx} = \frac{\partial \omega}{\partial k_x} = \pm \frac{N}{r} = C_x \quad \text{wave traveling in x direction}$$

$$C_{gz} = \frac{\partial \omega}{\partial r} = \mp \frac{KN}{r^2} = -C_z$$

Upward propagating waves have negative vertical <sup>phase</sup> group velocity:

Example: flow over topography



$$\frac{N}{r} = -\bar{u}$$

$$r = -N/\bar{u}$$

$$\text{But } -\frac{KN}{r^2} > 0$$

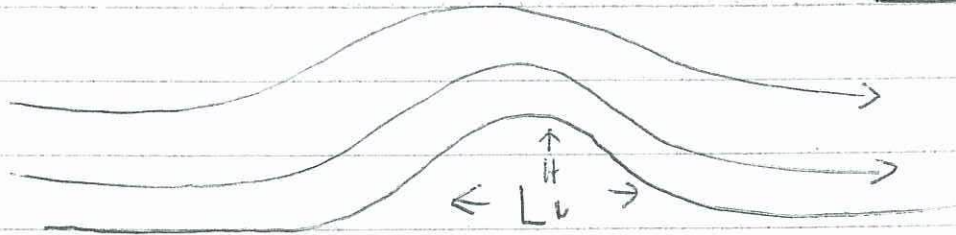
$$\left( \frac{\partial \omega}{\partial k} \right)_{\text{phase}} = -\frac{\bar{u}}{K} < 0$$

Downward phase speed

$K \neq 0$  so

Reel 5.1-5.11

## 2-D Hydrostatic Flow over a hill:



Valid if  $H/L \ll 1$

$$\frac{dy}{dt} = -\alpha \frac{\partial p}{\partial x}$$

$$-\alpha \frac{\partial p}{\partial z} = -g$$

First transform into system with  $p$  as the independent coordinate:

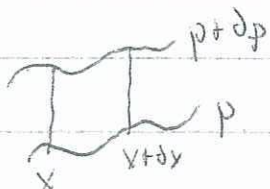
$$\text{Slope of surface of constant } p = \left( \frac{\partial z}{\partial x} \right)_p = - \frac{(\partial p / \partial x)_z}{(\partial p / \partial z)_x}$$

$$= \frac{(\partial p / \partial x)_z}{-pg}$$

$$\left[ \frac{dy}{dt} = -g \left( \frac{\partial z}{\partial x} \right)_p \right]$$

$$\left[ g \left( \frac{\partial z}{\partial p} \right)_x = -\alpha \right]$$

Hydrostatic form of mass continuity equation:



$$M = \rho dx dz = \rho dx \frac{\partial z}{\partial p} dp = -\frac{1}{g} dx dp$$

$$\begin{aligned}
 0 = \frac{dm}{dt} &= \rho u \frac{dz}{dx} \Big|_x - \rho u \frac{dz}{dx} \Big|_{x+\delta x} + \delta x \rho \frac{dp}{dt} \frac{\partial z}{\partial p} \Big|_p - \delta x \rho \frac{dp}{dt} \frac{\partial z}{\partial p} \Big|_{p+\delta p} \\
 &= \frac{1}{g} \left[ \delta p \left( u|_{x+\delta x} - u|_x \right) + \delta x \left( \omega|_{p+\delta p} - \omega|_p \right) \right] \quad \omega \equiv \frac{dp}{dt} \\
 &\quad \left[ \frac{\partial u}{\partial x} + \frac{\partial \omega}{\partial p} = 0 \right]
 \end{aligned}$$

Define a mass stream function:  $u = -\frac{\partial \psi}{\partial p}$   $\omega = \frac{\partial \psi}{\partial x}$

$$\frac{du}{dt} = -g \left( \frac{\partial z}{\partial x} \right)_p$$

$$g \frac{\partial z}{\partial p} = -\alpha$$

$$\Rightarrow \frac{\partial}{\partial p} \frac{du}{dt} = \left( \frac{\partial u}{\partial x} \right)_p = \left( \frac{\partial u}{\partial s} \right)_p \frac{\partial s}{\partial x} = \left( \frac{\partial T}{\partial p} \right)_s \frac{ds}{dT} \frac{\partial \psi}{\partial x}$$

$$= \left( \frac{\partial T}{\partial p} \right)_s \frac{ds}{dT} \frac{dp}{dt} = \frac{ds}{dT} \left( \frac{dT}{dt} \right)_s$$

$$\frac{\partial}{\partial p} \frac{du}{dt} = \frac{\partial}{\partial p} \left( u \frac{du}{dx} + \omega \frac{du}{dp} \right) = \frac{\partial u}{\partial p} + \frac{du}{\partial p} \frac{du}{dx} + \frac{\partial \omega}{\partial p} \frac{du}{dp}$$

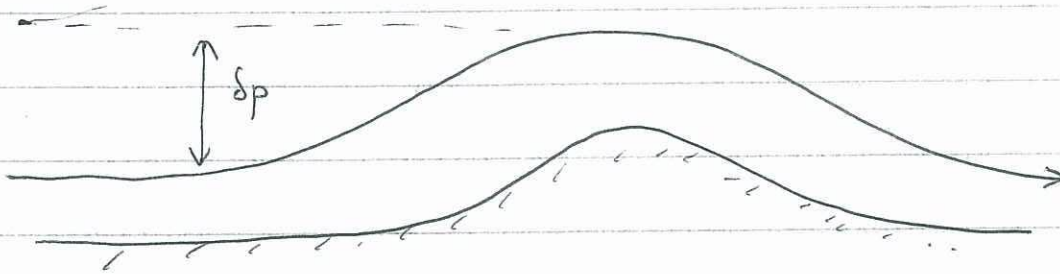
$$= \frac{d}{dt} \frac{du}{dp} + \frac{du}{dp} \left( \frac{du}{dx} + \frac{d\omega}{dp} \right) \rightarrow 0$$

$$\frac{d}{dt} \left[ \frac{\partial^2 \psi}{\partial p^2} + T \frac{ds}{dT} \right] = 0$$



First Consider simple case where upstream flow is constant

$$-\left(\frac{\partial k}{\partial p}\right)_0 = u_0 = \text{constant}$$



At  $a_0 \times p$ :  $\psi' = -\left(\frac{\partial k}{\partial p}\right)_0 \delta p = +u_0 \delta p$

Also,  $T' = \int \rho g \delta z = \int \rho g \delta p$

$$\left[ \frac{\partial^2}{\partial p^2} (+u_0 \delta p) + \frac{\rho g}{\rho g} \delta p \frac{ds}{dk} = 0 \right]$$

Upstream:  $\frac{ds}{dk} = \frac{ds/dp}{dk/dp} = \frac{ds/dp}{-u_0}$

Remember that  $N^2 \equiv \rho \frac{\partial s}{\partial z} = -\rho \frac{\partial s}{\partial p} \rho g$

$$\frac{ds}{dk} = \frac{N^2}{\rho g u_0}$$

$$\left[ \frac{\partial^2}{\partial p^2} \delta p + \frac{N^2}{\rho^2 g^2 u_0^2} \delta p = 0 \right]$$

Linear equation!

Consider case where  $\left[ \frac{N^2}{\rho g^2 u_0^2} = \text{Constant} = \frac{1}{l^2} \right]$

$$\left[ \frac{\partial^2}{\partial p^2} \delta p + l^2 \delta p = 0 \right]$$

$$\delta p = \delta p_0 \iint (\delta p_0)_{r,n} e^{i(r p - k x)} dr dk$$

$$r^2 = l^2$$

$$\left[ \delta p = \iint (\delta p_0)_{r,n} e^{i(l p - k x)} dr dk \right]$$

$$\psi' = + u_0 \delta p$$

$$\frac{\partial \psi'}{\partial p} = + u_0 \frac{\partial}{\partial p} \delta p = + u_0 \int (\delta p_0)_n l i e^{i(l p - k x)} dk$$

Sinusoidal topography:

$$u' = -u_0 l \delta p_0 \sin(l p - k x)$$

$$u' = -u_0 \left[ 1 + \frac{l \delta p_0 u_0}{\rho g u_0} \right] \sin(l p - k x)$$

$$= -u_0 \left[ 1 + \delta p_0 \frac{N u_0}{\rho g u_0} \right] \sin(l p - k x)$$

$$\delta p_0 = -\rho g \delta z_0$$

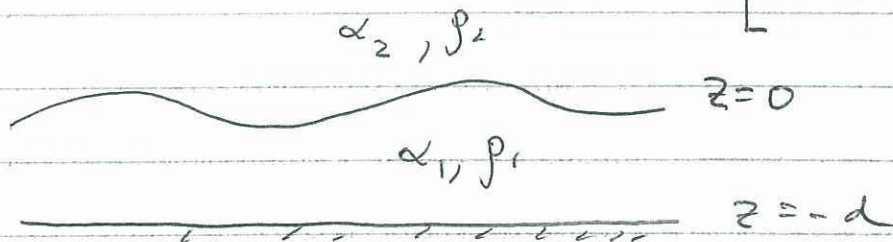
$$u' = -u_0 \left[ 1 - \frac{N \delta z_0}{g} \right] \sin(l p - k x)$$

Flow reversal at certain altitudes when  $\left[ \delta z_0 \geq a_0/N \right]$

Solution not viable then

Show solution

~~Stagger~~ Surface waves incompressible fluid  $\left[ \text{Read Chapter 5 in Faber} \right]$



Infinite fluid ! Consider 2-D problem first:

$$\rho \frac{du}{dt} = - \frac{\partial p}{\partial x}$$

$$\rho \frac{dw}{dt} = - \frac{\partial p}{\partial z} - \rho g$$

$$\Rightarrow \frac{\partial}{\partial z} \left[ \rho \frac{du}{dt} \right] - \frac{\partial}{\partial x} \left[ \rho \frac{dw}{dt} \right] = g \frac{\partial p}{\partial x}$$

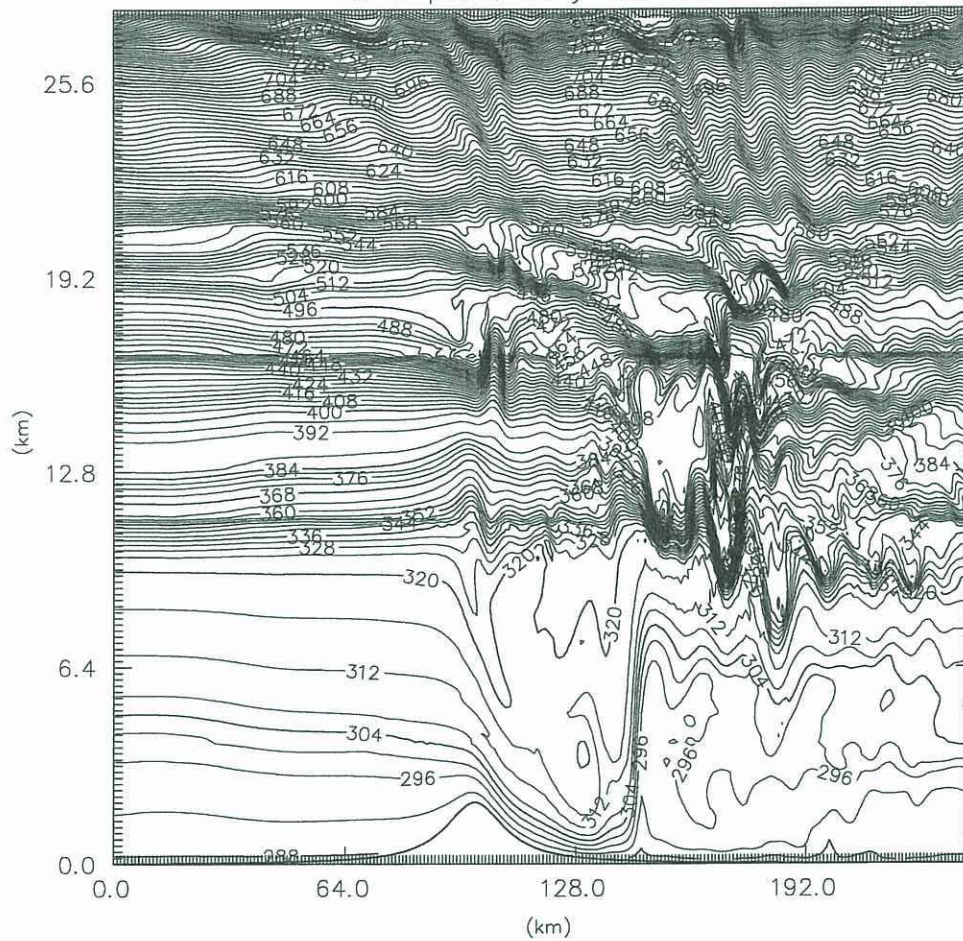
Linearize about resting state:

~~$$\frac{\partial}{\partial t} \left( \frac{\partial u}{\partial z} - \frac{\partial w}{\partial x} \right) = g \frac{\partial p}{\partial x} \quad \frac{\partial}{\partial z} \left( \frac{\partial u}{\partial t} \right) - \frac{\partial}{\partial x} \left( \frac{\partial w}{\partial t} \right) = g \frac{\partial p}{\partial x}$$~~

Now:  $\frac{dp}{dt} = 0$   $\frac{\partial p'}{\partial t} + w' \frac{dp}{dz} = 0$



19:00Z Thu 20 Jan 1977      t=14400.0 s (4:00:00)  
X-Z plane at y=0.5 km



pt (K, contour)

Min=285. Max=826. Inc=4.00

ARPS/ZXPLOT    wind1    Plot: 2002/01/22 12:56CST6CDT



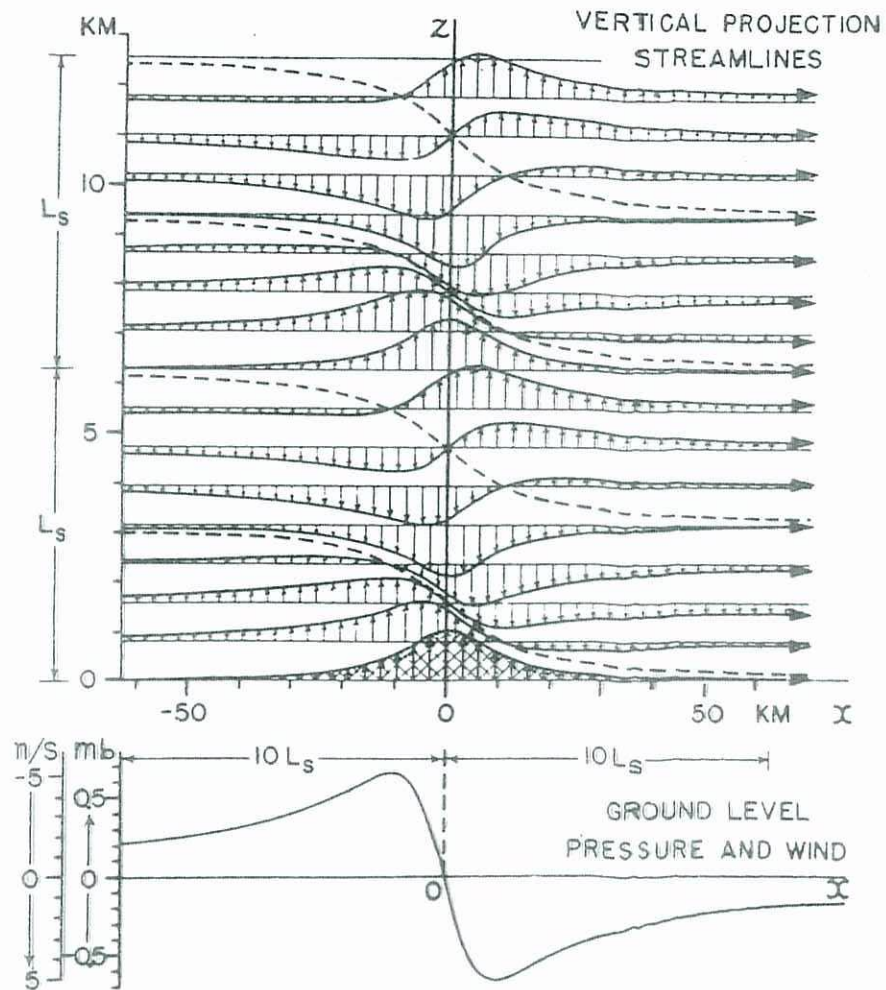


FIG. 3. Buoyancy-dominated hydrostatic flow over an isolated two-dimensional ridge, given by (2.56). The disturbance is composed of vertically propagating internal gravity waves of the sort shown in Fig. 1b. The evident upstream tilt of the phase lines indicates that disturbance energy is propagating upward away from the mountain. The maximum wind speed and minimum pressure occur on the lee slope of the ridge. The mountain height  $h_m = 1$  km, the half-width  $a = 10$  km, the mean wind speed  $U = 10$  m/sec, the Brunt-Vaisala frequency  $N = 0.01 \text{ sec}^{-1}$ , and the vertical wavelength  $L_s = 2\pi U/N = 6.28$  km. (From Queney, 1948.)

~~Oscillation frequency~~

Criticality

Stratified flows

Brigance oscillations: dispersion relation

Flow over topography: Linear and nonlinear

Surface waves

Rotating flows

Euler flows

Taylor - Proudman theorem

Point vortices

Vorticity equation: entropy gradients

potential vorticity

Rossby waves

Shear instability

Rayleigh Theorem

Parametric instability of shear flows

K-H instability

Equations in rotating reference frame: Coriolis

Geostrophy

Dynamics of "slow" motions: quasi-geostrophic flow

Rossby waves

Baroclinic Instability

Stellar physics

Marine physics

~~4~~ Eliminate  $p'$ :

$$\frac{\partial}{\partial z} \left( \bar{p} \frac{\partial^2 u'}{\partial t^2} \right) - \frac{\partial}{\partial x} \left( \bar{p} \frac{\partial^2 w'}{\partial t^2} \right) = -g \frac{\partial w'}{\partial x} \frac{d\bar{p}}{dz} \quad (1)$$

Within each layer  $p' = 0$   $\bar{p} = \text{constant}$

$$\nabla \cdot \underline{V} = 0: \quad u' = -\frac{\partial \psi}{\partial z} \quad w' = \frac{\partial \psi}{\partial x}$$

Then (1) becomes  $\frac{\partial^2}{\partial t^2} \nabla^2 \psi = 0$   $\nabla^2 \psi = 0$

$z > 0$ :  $\psi = A e^{iKx - i\omega t} e^{-Kz}$  decaying to zero at  $\infty$

Boundary at  ~~$z = -d$~~   $z = -d$ ,  $\frac{\partial \psi}{\partial x} = 0 \Rightarrow \psi = 0$

$$\psi = B e^{iKx - i\omega t} \sinh K(d+z)$$

Matching conditions at interface: 1.  $w^+ = w^- \Rightarrow \frac{\partial \psi^+}{\partial x} = \frac{\partial \psi^-}{\partial x} \Rightarrow \psi^+ = \psi^-$

$$A = B \sinh(Kd)$$

Second B.C.: Integrate (1) across interface

$$p_1 \frac{\partial^2 u_1}{\partial t^2} - p_2 \frac{\partial^2 u_2}{\partial t^2} = -g \frac{\partial w'}{\partial x} \Big|_+ (p_1 - p_2)$$

$$-\omega^2 p_1 A K - \omega^2 p_2 B K \cosh Kd = g K^2 A (p_1 - p_2)$$

$$\rho \frac{du}{dt} = \frac{\partial}{\partial t} \rho u + \nabla \cdot \rho u \underline{\underline{V}}$$

$$\rho \frac{dw}{dt} = \frac{\partial}{\partial t} (\rho w) + \nabla \cdot \rho w \underline{\underline{V}}$$

$$\rho \frac{du}{dt} = \rho \left( \frac{\partial u}{\partial t} + \underline{\underline{V}} \cdot \nabla u \right)$$

$$\frac{\partial \rho}{\partial t} = - \rho \nabla \cdot \underline{\underline{V}}$$

$$\rho u \equiv u'$$

$$\rho w \equiv w'$$

$$\frac{\partial}{\partial z} \frac{du}{dt} = \frac{\partial}{\partial x} \frac{dw}{dt} = 0$$

$$\nabla \cdot \underline{\underline{V}} = 0$$

$$\nabla \times \frac{d\underline{\underline{V}}}{dt} = 0$$

$$\underline{\underline{V}} \Rightarrow u = \frac{\partial \psi}{\partial x}$$

$$w = \frac{\partial \psi}{\partial z}$$

$$\nabla^2 \psi = 0$$

$$\frac{dz}{dt} = w_i = \frac{\partial z_i}{\partial t} + \underline{\underline{V}} \cdot \nabla z_i$$

$$\frac{\partial}{\partial z} \left[ \rho \frac{du}{dt} \right] = \frac{\partial}{\partial x} \left[ \rho \frac{dw}{dt} \right] = g \frac{\partial \rho}{\partial x}$$

$$\frac{d\rho}{dt} = 0$$



$$\nabla^2 \psi = 0$$

$$\rho \frac{\partial}{\partial t} \frac{\partial \psi}{\partial z} = \frac{\partial p}{\partial x}$$

$$\psi_{-}^+ = 0$$

$$\rho \frac{\partial}{\partial t} \frac{\partial \psi}{\partial x} = - \frac{\partial p}{\partial z} - \rho g$$

$$\frac{\partial}{\partial z} \left[ \rho \frac{\partial}{\partial t} \frac{\partial \psi}{\partial z} \right] = - g \frac{\partial p}{\partial x} - \frac{\partial}{\partial x} \rho \frac{\partial}{\partial t} \frac{\partial \psi}{\partial x}$$

$$\frac{\partial}{\partial t} \frac{\partial}{\partial z} \left[ \rho \frac{\partial}{\partial t} \frac{\partial \psi}{\partial z} \right] = + g \frac{\partial^2 \psi}{\partial x \partial z} \frac{\partial p}{\partial z} - \frac{\partial}{\partial x} \rho \frac{\partial^2 \psi}{\partial t^2} \frac{\partial \psi}{\partial x} + \frac{\partial}{\partial t} \frac{\partial \psi}{\partial x} \frac{\partial}{\partial x} \frac{\partial \psi}{\partial x} \frac{\partial p}{\partial z}$$

$$\frac{\partial}{\partial z} \left[ \rho \frac{\partial^2 \psi}{\partial t^2} \frac{\partial \psi}{\partial z} - \frac{\partial}{\partial t} \frac{\partial \psi}{\partial z} \frac{\partial \psi}{\partial x} \frac{\partial p}{\partial z} \right] = g \frac{\partial}{\partial x} \left[ \frac{\partial \psi}{\partial x} \frac{\partial p}{\partial z} \right] - \frac{\partial}{\partial x} \left[ \rho \frac{\partial^2 \psi}{\partial t^2} \frac{\partial \psi}{\partial x} \right] + \frac{\partial}{\partial t} \frac{\partial \psi}{\partial x} \frac{\partial}{\partial x} \left( \frac{\partial \psi}{\partial x} \frac{\partial p}{\partial z} \right)$$

$$\frac{\partial \psi}{\partial x} = \frac{\partial \psi}{\partial t} + u \frac{\partial \psi}{\partial x} = \frac{\partial \psi}{\partial z}$$

$$\frac{\partial \psi}{\partial t} + \frac{1}{2} u^2 + g \psi = 0$$

$$g =$$

$$\omega^2 (-\cosh(kd) p_2 - p_1) = -gk (p_2 - p_1)$$

$$\left[ \omega^2 = gk \frac{p_2 - p_1}{p_2 \cosh(kd) + p_1} \right]$$

Free surface (w/o surface tension):  $p_i \rightarrow 0$

$$\left[ \omega^2 = gk \tanh(kd) \right]$$

Deep water limit:  $d \rightarrow \infty, \tanh kd \rightarrow 1$

$$\omega^2 \rightarrow gk$$

$$\omega = \pm \sqrt{gk}$$

$$c = \pm \sqrt{\frac{g}{k}} = \pm \sqrt{\frac{1}{2\pi} g \lambda}$$

$$\underline{c_g} = \frac{\partial \omega}{\partial k} = \pm \frac{1}{2} \sqrt{\frac{g}{k}} = \underline{\underline{\frac{1}{2} c}}$$

Group velocity =  $\frac{1}{2}$  phase speed.

Boats moving through water generate wave with  $\lambda \approx L$ , when  $v=c$ ,

Resonance

$$v = \sqrt{\frac{gL}{2\pi}} \quad \text{"hull speed"}$$

$$10 \text{ metres} \approx 4 \text{ m s}^{-1} \text{ or } \sim 8 \text{ knots}$$

Shallow water limit:

$$d \rightarrow 0$$

$$\tanh kd \rightarrow kd$$

$$\omega^2 \Rightarrow g d k^2$$

$$\underline{c^2 \rightarrow g d}$$

Already, explore the shallow water equations

## Rotating Flows

Euler Fluid:

$$\frac{d\mathbf{v}}{dt} = -\alpha_0 \nabla p - \nabla g z$$

$$\nabla \times \frac{d\mathbf{v}}{dt} = 0$$

$$= \nabla \times \left[ \frac{\partial \mathbf{v}}{\partial t} + (\mathbf{v} \cdot \nabla) \mathbf{v} \right]$$

$$= \nabla \times \left( \frac{\partial \mathbf{v}}{\partial t} \right) + (\mathbf{v} \cdot \nabla) (\nabla \times \mathbf{v}) - \mathbf{v} (\nabla \times \mathbf{v} \cdot \nabla)$$

$$\boldsymbol{\omega} \equiv \nabla \times \mathbf{v}$$

$$\left[ \frac{d\boldsymbol{\omega}}{dt} = (\boldsymbol{\omega} \cdot \nabla) \mathbf{v} \right]$$

Stretching: illustrate

$$\text{also: } \frac{d\boldsymbol{\omega}}{dt} = -(\mathbf{v} \cdot \nabla) \boldsymbol{\omega} + (\boldsymbol{\omega} \cdot \nabla) \mathbf{v} = \nabla \times (\mathbf{v} \times \boldsymbol{\omega})$$

In a steady flow of an Euler Fluid

$$\nabla \times (\mathbf{v} \times \boldsymbol{\omega}) = 0$$

If inhomogeneity small,  $(\boldsymbol{\omega} \cdot \nabla) \mathbf{v} = 0$

no variations along rotation axis!

Taylor columns



2-D Euler flows: Flow in x-y plane

$$\omega_i = \omega_j = 0$$

$$\frac{d\omega_n}{dt} = 0$$

Conservation of Vorticity

$$\nabla \times \underline{V} = 0$$

$$\underline{V} = \hat{k} \times \nabla \psi$$

$$\nabla \cdot \underline{V} = 0$$

$$\nabla \times \underline{V} = \nabla^2 \psi$$

Point Vortex:

$$\nabla^2 \psi = 0$$

everywhere except at points

$$\frac{1}{r} \frac{\partial}{\partial r} r \frac{\partial \psi}{\partial r} = 0$$

$$r \frac{\partial \psi}{\partial r} = C$$

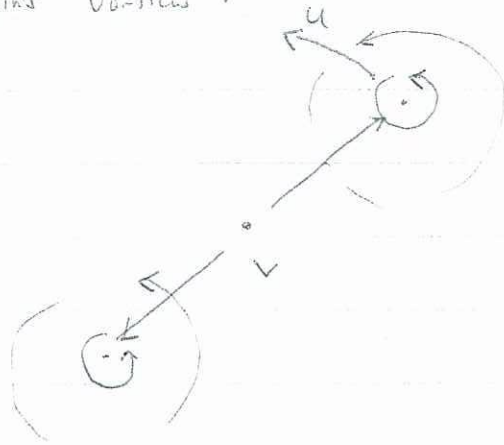
$$\frac{\partial \psi}{\partial r} = \frac{C}{r} \rightarrow \psi = C \ln r$$

$$C = \frac{\kappa}{2\pi} \quad \kappa = \text{circulation strength}$$

$$v = \frac{\kappa}{2\pi r}$$

$$\oint_{\text{closed}} v \, dl = \kappa$$

Consider two-point vertices:



$$\left[ u = \frac{k}{2\pi L} \right]$$

rotate around each other

~~So~~ Opposite Sign vertices:



$$u = \frac{k}{2\pi L}$$

-842 -64-

General Compressible Fluid:

$$\frac{dV}{dt} = -\alpha \nabla p - \nabla g z$$

$$\nabla \times \frac{dV}{dt} = \frac{d\omega}{dt} - (\omega \cdot \nabla) \underline{V} = -\nabla \alpha \times \nabla p$$

$$\alpha = \alpha(p, s)$$

$$\left[ \frac{d\omega}{dt} = (\omega \cdot \nabla) \underline{V} - \left( \frac{\partial \alpha}{\partial s} \right)_p \nabla s \times \nabla p \right]$$

Example: 2-D flow in  $x$ - $z$  plane:

$$\underline{\omega} = \left( \frac{\partial w}{\partial x} - \frac{\partial u}{\partial z} \right) \hat{j} = \omega_0$$

$$(\underline{\omega} \cdot \nabla) \underline{V} = 0$$

$$\frac{d\omega_0}{dt} = - \left( \frac{\partial \alpha}{\partial s} \right)_p \left( \frac{\partial s}{\partial x} \frac{\partial p}{\partial z} - \frac{\partial s}{\partial z} \frac{\partial p}{\partial x} \right) = - \left( \frac{\partial \alpha}{\partial s} \right)_p \left( \frac{\partial s}{\partial x} \right)_p \frac{\partial p}{\partial z}$$

$$\approx - \left( \frac{\partial \alpha}{\partial s} \right)_p \frac{\partial s}{\partial x} \frac{\partial p}{\partial z} = - \left( \frac{\partial T}{\partial p} \right)_s \frac{\partial p}{\partial z} \left( \frac{\partial s}{\partial x} \right)_p = \rho \left( \frac{\partial s}{\partial x} \right)_p$$

Contours of entropy on isobaric surfaces imply velocity changes

# Potential Vorticity

$$\frac{1}{\rho} \nabla s \cdot \frac{d\omega}{dt} = \frac{1}{\rho} \nabla s \cdot ((\omega \cdot \nabla) \underline{v}) - \frac{1}{\rho} \left( \frac{\partial \omega}{\partial s} \right)_p \nabla s \cdot (\nabla s \times \nabla \phi) \quad \rightarrow 0$$

$$= \frac{1}{\rho} \nabla s \cdot \left( \frac{\partial \omega}{\partial t} + (\underline{v} \cdot \nabla) \underline{v} \right)$$

$$= \frac{1}{\rho} \frac{\partial}{\partial t} \left( \omega \cdot \nabla s \right) - \frac{1}{\rho} \frac{(\omega \cdot \nabla) \frac{\partial s}{\partial t}}{\partial t} + \frac{1}{\rho} (\underline{v} \cdot \nabla) (\omega \cdot \nabla s) - \frac{1}{\rho} \omega \cdot (\underline{v} \cdot \nabla) \nabla s$$

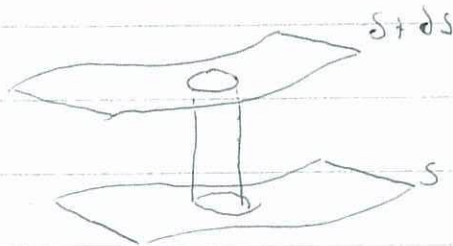
$$\Rightarrow \frac{1}{\rho} \frac{d}{dt} (\omega \cdot \nabla s) = \frac{1}{\rho} \omega \cdot \left[ \nabla \frac{\partial s}{\partial t} + (\underline{v} \cdot \nabla) \nabla s \right] + \frac{1}{\rho} \nabla s \cdot [(\omega \cdot \nabla) \underline{v}]$$

$$= \frac{1}{\rho} \omega \cdot \nabla \left( \frac{\partial s}{\partial t} + (\underline{v} \cdot \nabla) s \right) - \frac{1}{\rho} \nabla s \cdot ((\omega \cdot \nabla) \underline{v}) + \frac{1}{\rho} \nabla s \cdot [(\omega \cdot \nabla) \underline{v}] - \frac{1}{\rho} (\omega \cdot \nabla s) (\nabla \cdot \underline{v})$$

$$\frac{d}{dt} \left( \frac{\omega \cdot \nabla s}{\rho} \right) = 0 \quad = \quad \frac{1}{\rho^2} \frac{d}{dt} \left( \omega \cdot \nabla s \right) + \frac{1}{\rho^2} (\omega \cdot \nabla s)'$$

Conservation of potential vorticity

Alternative derivation:



$$\frac{d\underline{v}}{dt} = -\alpha \nabla p$$

$$\oint \frac{d\underline{v}}{dt} \cdot d\underline{l} = - \oint \alpha \nabla p \cdot d\underline{l} = - \oint \alpha(s, p) \cdot \nabla p \cdot d\underline{l} = 0$$

$$= \frac{d}{dt} \left( \oint \nabla \times \underline{v} \cdot d\underline{A} \right) = \frac{d}{dt} \left[ (\nabla \times \underline{v}) \cdot \frac{\nabla s}{|\nabla s|} \delta A \right] = 0$$



$$\delta m = \rho \delta A \delta h = \rho \delta A \frac{\partial h}{\partial s} \delta s = \frac{\rho \delta A \delta s}{|\nabla s|}$$

$$\frac{d}{dt} \left[ (\nabla \times \underline{v}) \cdot \nabla s \delta s \delta m / \rho \right] = 0$$

But  $\delta s, \delta m$  constant

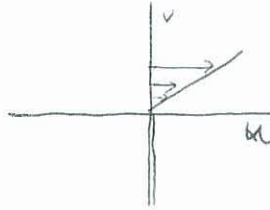
$$\frac{d}{dt} \left[ \frac{1}{\rho} \nabla \times \underline{v} \cdot \nabla s \right] = 0$$

Conservation in rotating flows!

## Rossby Waves:

Ubiquitous in rotational flows. Requires gradient in vorticity.

Consider the following flow: of an Euler fluid:



$$u = \begin{cases} \beta y & y > 0 \\ 0 & y \leq 0 \end{cases}$$

$$\vec{\omega} = \nabla \times \vec{V} = - \frac{\partial u}{\partial y} \hat{n} = \begin{cases} -\beta \hat{n} & y > 0 \\ 0 & y < 0 \end{cases}$$

$$\frac{\partial u}{\partial t} + \vec{V} \cdot \nabla u = -\alpha_0 \frac{\partial p}{\partial x}$$

$$\frac{\partial v}{\partial t} + \vec{V} \cdot \nabla v = -\alpha_0 \frac{\partial p}{\partial y}$$

$$\frac{\partial u}{\partial x} + \frac{\partial v}{\partial y} = 0$$

Linearize about base state:

$$\frac{\partial \bar{u}}{\partial t} + \bar{u} \frac{\partial \bar{u}}{\partial x} + \bar{v} \frac{\partial \bar{u}}{\partial y} = -\alpha_0 \frac{\partial \bar{p}}{\partial x}$$

$$\frac{\partial \bar{v}}{\partial t} + \bar{u} \frac{\partial \bar{v}}{\partial y} = -\alpha_0 \frac{\partial \bar{p}}{\partial y}$$

$$\frac{\partial \bar{u}}{\partial x} + \frac{\partial \bar{v}}{\partial y} = 0$$

Look for disturbances of form  $F(y)e^{ik(x-ct)}$

$$ik(\bar{u}-c)u + v\bar{u}_y = -ik\alpha_0 p \quad (1)$$

$$ik(\bar{u}-c)v = -\alpha_0 \frac{dp}{dy} \quad (2)$$

$$iku = -\frac{dv}{dy} \quad (3)$$

Eliminate  $p$  using (1), (2):

$$ik \frac{d}{dy} \{ (\bar{u}-c)u \} + \frac{dv}{dy} \bar{u}_y + v\bar{u}_{yy} = -k^2(\bar{u}-c)v$$

Eliminate  $u$  using (3):

$$-\frac{d}{dy} \left( (\bar{u}-c) \frac{dv}{dy} \right) + \bar{u}_y \frac{dv}{dy} + v\bar{u}_{yy} = -k^2(\bar{u}-c)v$$

$$-(\bar{u}-c) \frac{d^2 v}{dy^2} + k^2(\bar{u}-c)v + v\bar{u}_{yy} = 0$$

$$\left[ \frac{d^2 v}{dy^2} - \left( k^2 + \frac{\bar{u}_{yy}}{\bar{u}-c} \right) v \right] = 0 \quad (4)$$

(Genent; haven't made assumptions about  $\bar{u}(y)$  yet)

In particular problem,  $\bar{u}_{yy} = 0$  in each region:

$$\frac{d^2 \bar{v}}{dy^2} - k^2 v = 0$$

$$y > 0 \quad v = A e^{-ky}$$

$$y < 0 \quad v = B e^{ky}$$

matching:  $V|_+ = V|_- \quad A = B$

Integrate (4) across  $y=0$ :

$$\left[ \frac{dv}{dy} \right]_+^- = \left[ \frac{\bar{u}_{yy}}{\bar{u}-c} \right]_+^- V_0$$

$$-kA e^{-k \cdot 0} - kA = \frac{A}{-c} B$$

$$\left[ 2k = \frac{B}{c} \right]$$

$\Rightarrow$

$$c = \frac{B}{2k}$$

Rossby wave

$$B > 0 \quad c > 0, \quad B < 0, \quad c < 0$$

Dispersive in odd way

$$\omega = ck = \frac{B}{2}$$

$$\frac{d\omega}{dk} = 0!$$

Group velocity = 0!

More about Rossby waves later.

(Go on to rotating reference frames)

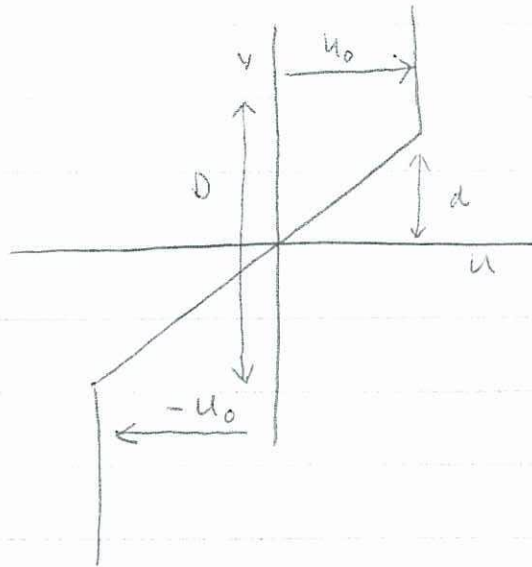
Hemlock model



## Shear Instability:

Rossby wave trains whose intrinsic phase speeds are of opposite sign may occasionally interact unstably through phase locking.

Example:



Each interface, acting independently, supports Rossby waves whose phase speed is given by

$$C = \frac{\beta}{2\kappa} - U_0 = \frac{U_0}{\frac{D\kappa}{Dx} - U_0} = U_0 \left[ \frac{1}{\frac{D\kappa}{Dx} - 1} \right]$$

lower interface

$$C = U_0 - \frac{U_0}{2\kappa} = U_0 \left[ 1 - \frac{1}{\frac{D\kappa}{Dx}} \right] \quad \text{upper interface}$$

Sometimes waves can phase lock and make each other unstable. Very different from convective instability.

$$\frac{d^2 \tilde{v}}{dy^2} - \kappa^2 \tilde{v} - \tilde{v} \frac{d^2 \bar{u}/dy^2}{\bar{u} - c} = 0$$

$$v = \tilde{v} e^{i\kappa(x-ct)}$$

$$y > d \quad \tilde{v} = A e^{-\kappa y}$$

$$\tilde{v}, c \text{ complex}$$

$$y < -d \quad \tilde{v} = F e^{\kappa y}$$

$$-d \leq y \leq d \quad \tilde{v} = B e^{-\kappa y} + C e^{\kappa y}$$

Matching at interfaces:

$v$  Continuous

$$i\kappa(\bar{u}-c)u + v \frac{d\bar{u}}{dy} = -i\kappa d \phi$$

$$i\kappa u = -\frac{dv}{dy}$$



$$(\bar{u}-c) \frac{dv}{dy} = v \frac{d\bar{u}}{dy}$$

Continuous

$$\text{Four conditions for 4 constants} \Rightarrow \left( \frac{D\kappa c}{u_0} \right)^2 = (D\kappa - 1)^2 - e^{-2\kappa D}$$

$$\kappa D \rightarrow \infty: \quad c^2 \rightarrow u_0^2 \left[ 1 - \frac{2}{\kappa D} \right]$$

$$c = \pm u_0 \left[ 1 - \frac{1}{\kappa D} \right] \quad \text{Independent Rossby waves}$$

$$\kappa D \rightarrow 0: \quad \frac{D^2 \kappa^2 c^2}{u_0^2} \approx D^2 \kappa^2 - 2D\kappa + 1 = 1 + 2\kappa D - 2D^2 \kappa^2 = -D^2 \kappa^2$$

$$c^2 \rightarrow -u_0^2$$

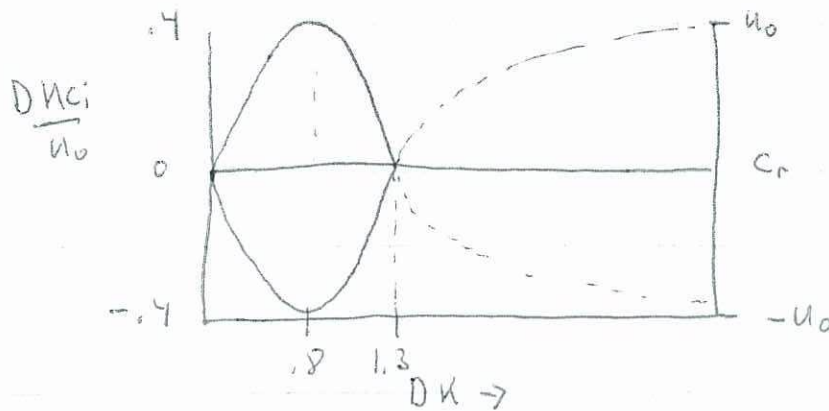
$$e^{i\kappa(x-ct)} \sim e^{i\kappa(x-ct) + \kappa c t}$$

$$\kappa c_i = \pm \kappa u_0$$

Growth rate linear in  $\kappa$ :

Other values:  $C^2 = 0$  at  $DK = 1.778$  ( $L = \frac{2\pi}{k} = 4.92 D$ )

$HC$ : maximum of  $.402 u_0/D$  at  $L = 7.884 D$



Explanation in terms of phase-locked waves, vortex displacements

Necessary conditions for instability in general flows

Start with 
$$\frac{d^2 \tilde{v}}{dy^2} - u^2 \tilde{v} - \tilde{v} \frac{d^2 \bar{u}/dy^2}{\bar{u} - c} = 0$$

Multiply through by  $\tilde{v}^*$ , integrate over domain

$$\int_{-\infty}^{\infty} \left[ \tilde{v}^* \frac{d^2 \tilde{v}}{dy^2} - u^2 |\tilde{v}|^2 - |\tilde{v}|^2 \frac{d^2 \bar{u}/dy^2}{\bar{u} - c} \right] dy = 0$$

Integrate first term by parts:

$$- \int_{-\infty}^{\infty} \left[ \left| \frac{d \tilde{v}}{dy} \right|^2 + |\tilde{v}|^2 \left( u^2 + \frac{d^2 \bar{u}/dy^2}{\bar{u} - c} \right) \right] dy = 0$$

Real and imaginary parts must vanish separately:

$$\text{Im: } \int_{-\infty}^{\infty} |\tilde{v}|^2 \frac{d^2 \bar{u}}{dy^2} \frac{\bar{u} - c_r}{(\bar{u} - c_r)^2 + c_i^2} dy = 0$$

For  $c_i > 0$  (growing waves)  $\frac{d^2 \bar{u}}{dy^2}$  must change sign in domain

$\Rightarrow \frac{d\bar{u}}{dy}$  must have an extreme value

$$\text{Re: } \int_{-\infty}^{\infty} \left[ \left| \frac{d\tilde{v}}{dy} \right|^2 + |\tilde{v}|^2 \left( k^2 + \frac{d^2 \bar{u}}{dy^2} \frac{\bar{u} - c_r}{(\bar{u} - c_r)^2 + c_i^2} \right) \right] dy = 0$$

$$\Rightarrow \int_{-\infty}^{\infty} |\tilde{v}|^2 \frac{d^2 \bar{u}}{dy^2} \frac{\bar{u} - c_r}{(\bar{u} - c_r)^2 + c_i^2} dy < 0$$

$$\text{But } \int_{-\infty}^{\infty} |\tilde{v}|^2 \frac{d^2 \bar{u}}{dy^2} \frac{1}{(\bar{u} - c_r)^2 + c_i^2} dy = 0$$

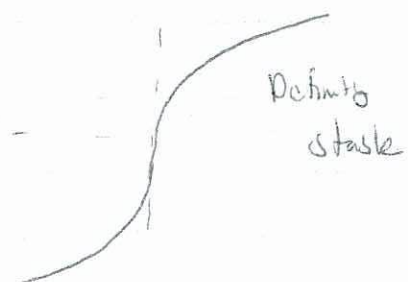
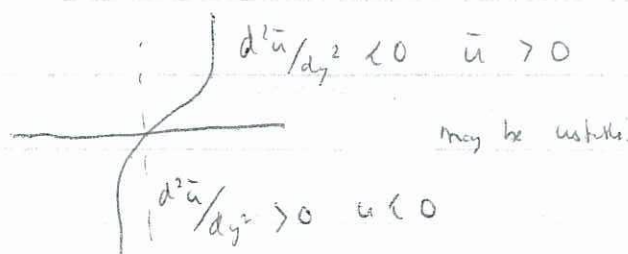
$$\Rightarrow \int_{-\infty}^{\infty} |\tilde{v}|^2 \frac{\bar{u} \frac{d^2 \bar{u}}{dy^2}}{(\bar{u} - c_r)^2 + c_i^2} dy < 0$$

$(\bar{u})$  ~~is~~ negatively correlated with  $d^2 \bar{u}/dy^2$

For  $c_i$  additive constant  $\bar{u}$

Example: Symmetric profile ~~for which~~  $c_r = 0$ :

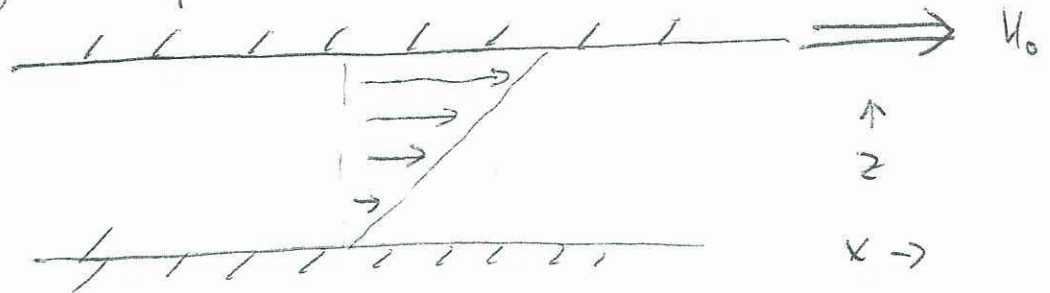
Example:





# Parametric Instability of Shear Flows:

Labratory Set-up:



Vorticity constant: Rayleigh Theorem: no instability.

Equation of vorticity about  $y$  axis:

$$\frac{d\eta}{dt} = 0 \quad \eta = \frac{\partial w}{\partial x} - \frac{\partial u}{\partial z} = \nabla^2 \psi$$

Linearize:

$$\left( \frac{\partial}{\partial t} + \bar{u} \frac{\partial}{\partial x} \right) \eta' = 0$$

$$\left( \frac{\partial}{\partial t} + \lambda z \frac{\partial}{\partial x} \right) \eta' = 0$$

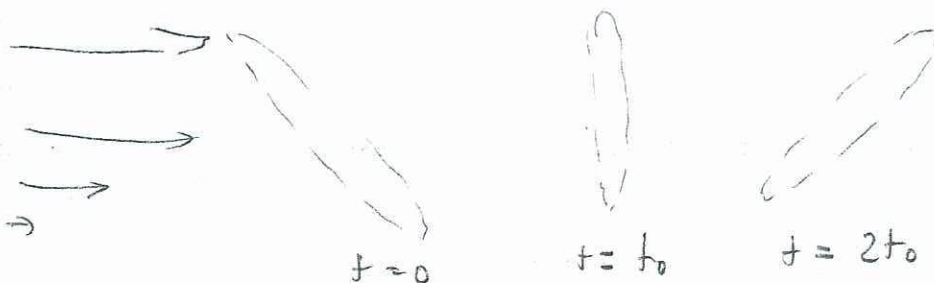
Unbounded case: Satisfied by  $\eta = F(x - \lambda z(t - t_0))$

Non-modal

(not separable in time)

Example:

$$\eta = \eta_0 \sin k [x - \lambda z(t - t_0)]$$



Now solve for  $\psi$ :

$$\nabla^2 \psi = A_0 \sin k [x - \lambda z (t - t_0)]$$

$$\Rightarrow \psi = \frac{-\eta_0}{k^2 (1 + \lambda^2 (t - t_0)^2)} \sin k [x - \lambda z (t - t_0)]$$

$$u' = -\frac{\partial \psi}{\partial z} = \frac{-\eta_0 \lambda (t - t_0)}{k^2 [1 + \lambda^2 (t - t_0)^2]} \cos k [x - \lambda z (t - t_0)]$$

$$w' = \frac{\partial \psi}{\partial x} = \frac{-\eta_0}{k^2 [1 + \lambda^2 (t - t_0)^2]} \cos k [x - \lambda z (t - t_0)]$$

$$KE' = \frac{1}{2} (u'^2 + w'^2) = \frac{\frac{1}{2} \eta_0^2 \cos^2 k [x - \lambda z (t - t_0)]}{k^2 [1 + \lambda^2 (t - t_0)^2]}$$

max at  $t = t_0$

# Kelvin-Helmholtz instability

Shear in the presence of stable stratification:

$$\frac{du}{dt} = -\alpha_0 \frac{\partial p}{\partial x}$$

$$\frac{dw}{dt} = -\alpha_0 \frac{\partial p}{\partial z} - \int \frac{ds}{dz} dz$$

$$\frac{\partial u}{\partial x} + \frac{\partial w}{\partial z} = 0$$

$$\bar{u} = \bar{u}(z) + u'$$

$$\frac{\partial u'}{\partial t} + \bar{u} \frac{\partial u'}{\partial x} + w \frac{d\bar{u}}{dz} = -\alpha_0 \frac{\partial p'}{\partial x}$$

$$\frac{\partial w'}{\partial t} + \bar{u} \frac{\partial w'}{\partial x} = -\alpha_0 \frac{\partial p'}{\partial z} - \int \frac{ds}{dz} dz'$$

$$\left( \frac{\partial}{\partial t} + \bar{u} \frac{\partial}{\partial x} \right) dz' = w'$$

$$\frac{\partial u'}{\partial x} + \frac{\partial w'}{\partial z} = 0$$

$$u', w', z', p' \sim F(z) e^{ik(x-ct)}$$

$$-(\bar{u}-c) \frac{dw}{dz} + w \frac{d\bar{u}}{dz} = -\alpha_0 i k p$$

$$i k (\bar{u}-c) w = -\alpha_0 \frac{dp}{dz} - \int \frac{d\bar{s}}{dz} dz$$

$$i k (\bar{u}-c) \delta z = w$$

$$\Rightarrow \frac{d}{dz} \left[ (\bar{u}-c) \frac{dw}{dz} - w \frac{d\bar{u}}{dz} \right] + w \left[ \frac{\int \frac{d\bar{s}}{dz}}{\bar{u}-c} - k^2 (\bar{u}-c) \right] = 0 \quad (1)$$

$$\Rightarrow \frac{1}{2} u \quad S_0 + \frac{1}{2} \Delta S$$

$$\leftarrow \frac{1}{2} u \quad S_0 - \frac{1}{2} \Delta S$$

In each layer:  $\frac{d^2 w}{dz^2} - k^2 w = 0$

$$w^+ = A e^{-kz}$$

$$w^- = B e^{kz}$$

Matching: Displacement, not  $w$ !

$$\delta z = \frac{w}{i k (\bar{u}-c)}$$

$$\left[ \frac{A}{\frac{1}{2} u - c} = \frac{B}{-\frac{1}{2} u - c} \right]$$

Integrate 1:

$$(\bar{u}-c) \frac{dw}{dz} \Big|_{-}^{+} + \frac{w}{\bar{u}-c} \int \Delta \bar{s} = 0$$

$$\Rightarrow -\left(\frac{1}{2} u - c\right) k A - \left(-\frac{1}{2} u - c\right) k B + \frac{A \int \Delta \bar{s}}{\frac{1}{2} u - c} = 0$$

$$\Rightarrow \left[ c^2 = \frac{1}{2} \frac{\int \Delta \bar{s}}{k} - \frac{1}{4} u^2 \right] \quad \text{unstable for } k \text{ sufficiently large}$$



General criterion for instability:

Define  $v \equiv \frac{w}{(\bar{u}-c)^{1/2}}$

Substitute into (1):

$$\frac{d}{dz} \left[ (\bar{u}-c) \frac{dv}{dz} \right] + v \left[ \frac{\int \frac{d\bar{u}}{dz} - \frac{1}{4} \left( \frac{d\bar{u}}{dz} \right)^2}{(\bar{u}-c)} - k^2(\bar{u}-c) \right] = 0$$

Multiply through by  $v^*$  and integrate over all  $z$ :

$$\int_{-\infty}^{\infty} \left\{ -(\bar{u}-c) \left| \frac{dv}{dz} \right|^2 + |v|^2 \left[ \frac{\int \frac{d\bar{u}}{dz} - \frac{1}{4} \left( \frac{d\bar{u}}{dz} \right)^2}{\bar{u}-c} - k^2(\bar{u}-c) \right] \right\} dz = 0$$

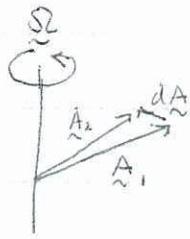
Imaginary part:

$$c_i \int_{-\infty}^{\infty} \left\{ 2 \left| \frac{dv}{dz} \right|^2 + |v|^2 \left[ \frac{\int \frac{d\bar{u}}{dz} - \frac{1}{4} \left( \frac{d\bar{u}}{dz} \right)^2}{(\bar{u}-c_r)^2 + c_i^2} - k^2 \right] \right\} dz = 0$$

$$\frac{\int \frac{d\bar{u}}{dz}}{\left( \frac{d\bar{u}}{dz} \right)^2} < \frac{1}{4} \quad \text{necessary condition}$$

# Equations in a rotating reference frame

Consider a fixed <sup>position</sup> vector  $\vec{A}$  on a rotating reference frame at two different times.



Clearly,  $d\vec{A} = (\Omega \times \vec{A})dt$

or  $\frac{d\vec{A}}{dt} = \Omega \times \vec{A}$

One such vector is the unit vector  $\hat{i}$ . In absolute coordinates, any vector  $\vec{B}$  can be written

$$\vec{B} = B_x \hat{i} + B_y \hat{j} + B_z \hat{k} = \vec{B} \cdot \vec{i}$$

or

$$\begin{aligned} \frac{d\vec{B}}{dt} &= \frac{dB_x}{dt} \hat{i} + \frac{dB_y}{dt} \hat{j} + \frac{dB_z}{dt} \hat{k} + B_x \frac{d\hat{i}}{dt} + B_y \frac{d\hat{j}}{dt} + B_z \frac{d\hat{k}}{dt} \\ &= \left( \frac{d\vec{B}}{dt} \right)_r + \Omega \times \vec{B} \end{aligned}$$

$\Rightarrow \frac{d\hat{i}}{dt} = \Omega \times \hat{i}, \frac{d\hat{j}}{dt} = \Omega \times \hat{j}$  etc.

So  $\left( \frac{d}{dt} \right)_a = \left( \frac{d}{dt} \right)_r + \Omega \times$

Position: Velocity:

$$\left(\frac{d\vec{r}}{dt}\right)_a = \left(\frac{d\vec{r}}{dt}\right)_r + \vec{\Omega} \times \vec{r}$$

Acceleration:

$$\left(\frac{d}{dt}\right)_a \left(\frac{d\vec{r}}{dt}\right)_a = \left(\left(\frac{d}{dt}\right)_r + \vec{\Omega} \times\right) \left(\left(\frac{d\vec{r}}{dt}\right)_r + \vec{\Omega} \times \vec{r}\right)$$

$$= \left(\frac{d\vec{v}_r}{dt}\right)_r + 2\vec{\Omega} \times \vec{v}_r + \vec{\Omega} \times \vec{\Omega} \times \vec{r}$$

$$= \left(\frac{d\vec{v}_r}{dt}\right)_r + 2\vec{\Omega} \times \vec{v}_r - |\vec{\Omega}|^2 \vec{R}$$

$\vec{R} \equiv$  vector perpendicular to axis

$$\frac{d\vec{v}_r}{dt} = \frac{d\vec{v}_a}{dt} - \underbrace{2\vec{\Omega} \times \vec{v}_r}_{\text{Coriolis}} + \underbrace{|\vec{\Omega}|^2 \vec{R}}_{\text{Centrifugal}}$$

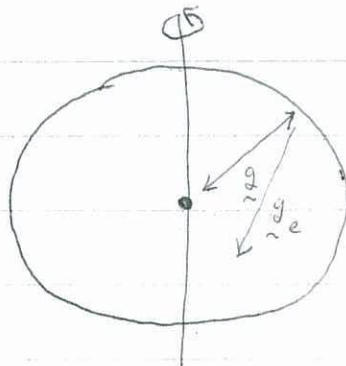
Inviscid equations:

$$\frac{d\vec{v}_a}{dt} = -\alpha \nabla p + \vec{g}$$

$$\Rightarrow \frac{d\vec{v}_r}{dt} = -\alpha \nabla p - 2\vec{\Omega} \times \vec{v}_r + \underbrace{(\vec{g} + |\vec{\Omega}|^2 \vec{R})}_{\text{effective gravity}}$$

$$= -\nabla \phi_e$$

"effective gravity"



oblate spheroid

Oblateness small for earth; we shall use strictly spherical coordinates. Here another coordinate transformation is necessary to convert from Cartesian to Spherical Coordinates. Definitions:

$$u \equiv r \cos \theta \frac{d\lambda}{dt}$$

$$v \equiv r \frac{d\theta}{dt}$$

$$w \equiv \frac{dz}{dt}$$

$$\frac{dV_r}{dt} = \hat{i} \left( \frac{du}{dt} - \frac{uv \tan \theta}{r} + \frac{uw}{r} \right) + \hat{j} \left( \frac{dv}{dt} + \frac{u^2 \tan \theta}{r} + \frac{vw}{r} \right) + \hat{k} \left( \frac{dw}{dt} - \frac{u^2 + v^2}{r} \right)$$

$$\Rightarrow \begin{cases} \frac{du}{dt} = \frac{uv \tan \theta}{r} - \frac{uw}{r} + 2|\underline{\Omega}| \sin \theta v - 2|\underline{\Omega}| \cos \theta w - \alpha \frac{\partial p}{\partial x} \\ \frac{dv}{dt} = -\frac{u^2 \tan \theta}{r} - \frac{vw}{r} - 2|\underline{\Omega}| \sin \theta u - \alpha \frac{\partial p}{\partial y} \\ \frac{dw}{dt} = \frac{u^2 + v^2}{r} + 2|\underline{\Omega}| \cos \theta u - \alpha \frac{\partial p}{\partial z} - g_{eff} \end{cases}$$

In sufficiently slow or shallow motions,



$$\frac{du}{dt} \approx -\alpha \frac{\partial p}{\partial x} + f v$$

$$\frac{dv}{dt} = -\alpha \frac{\partial p}{\partial y} - f u$$

$$f \equiv 2|\underline{\Omega}| \sin \theta$$

$$\frac{dw}{dt} = -\alpha \frac{\partial p}{\partial z} - g_e$$

or

$$\frac{d\underline{V}}{dt} = -\alpha \nabla p - \nabla \phi_e + f \hat{n} \times \underline{V}$$

Sufficiently slow motions:

$$\alpha \frac{\partial p}{\partial x} \approx f v \quad \equiv f v_y$$

$$-\alpha \frac{\partial p}{\partial y} = f u \quad \equiv f u_x$$

What is "sufficiently slow"?

$$V = V_y + \frac{1}{f} \frac{dV_y}{dt}$$

$$u = u_x - \frac{1}{f} \frac{du_x}{dt}$$

$$\frac{dV}{dt} = \frac{dV_y}{dt} + \frac{1}{f} \frac{d^2 V_y}{dt^2} = \frac{dV_y}{dt} + \frac{1}{f} \frac{d^2 u_x}{dt^2} - \frac{1}{f^2} \frac{d^3 V_y}{dt^3} \dots$$

Define acceleration time scale  $\tau$ :  $t \rightarrow \tau t$

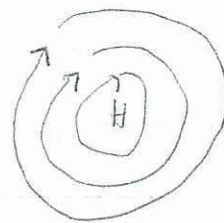
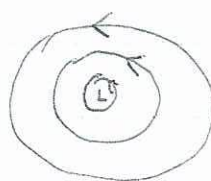
~~$$\frac{dV}{dt}$$~~

$$V = V_y + \frac{1}{f\tau} \frac{dV_y}{dt} - \frac{1}{f^2 \tau^2} \frac{d^2 V_y}{dt^2} + \dots$$

$$R_0 \equiv \frac{1}{f\tau}$$

$$V = V_y + R_0 \frac{dV_y}{dt} - R_0^2 \frac{d^2 V_y}{dt^2} + \dots$$

If  $R_0 \ll 1$ ,  $\underline{V} \approx \underline{V}_\theta$   
 $u \approx u_\theta$



Some of consequences of geostrophic balance:

$$fu \approx -\alpha \frac{\partial p}{\partial y}$$

$$fv \approx +\alpha \frac{\partial p}{\partial x}$$

$$g \approx -\alpha \frac{\partial p}{\partial z}$$

Pressure Co-ordinates:

$$dp = \left(\frac{\partial p}{\partial x}\right)_{y,z} dx + \left(\frac{\partial p}{\partial y}\right)_{x,z} dy + \left(\frac{\partial p}{\partial z}\right)_{x,y} dz = 0$$

$$\left(\frac{\partial p}{\partial y}\right)_{x,z} = - \left(\frac{\partial p}{\partial z}\right)_{x,y} \left(\frac{\partial z}{\partial x}\right)_p = \rho g \left(\frac{\partial z}{\partial x}\right)_p$$

$$\left(\frac{\partial p}{\partial x}\right)_{y,z} = \rho g \left(\frac{\partial z}{\partial y}\right)_p$$

$$fu = -g \left(\frac{\partial z}{\partial y}\right)_p = -\frac{\partial \phi}{\partial x}$$

$$fv = \frac{\partial \phi}{\partial x}$$

$$\frac{\partial \phi}{\partial p} = -\alpha$$

$$\text{Now } f \frac{\partial u}{\partial p} = - \frac{\partial}{\partial p} \left( \frac{\partial \psi}{\partial y} \right)_p = - \left( \frac{\partial}{\partial y} \right)_p \frac{\partial \psi}{\partial p} = \left( \frac{\partial \alpha}{\partial y} \right)_p$$

$$\text{Similarly, } \left[ f \frac{\partial v}{\partial p} = - \left( \frac{\partial \alpha}{\partial x} \right)_p \right]$$

$$\text{Thermal wind relation: } f \frac{\partial \underline{v}}{\partial p} = \hat{n} \times \nabla \alpha$$

$$= - \hat{n} \times \left( \nabla_p \right) \frac{RT}{p}$$

$$\left[ f \frac{\partial \underline{v}}{\partial p} = - \frac{R}{p} \hat{n} \times \nabla_p T \right]$$

$$\text{Also, } (\partial \alpha)_p = \left( \frac{\partial \alpha}{\partial s} \right)_p (\delta s)_p = \left( \frac{\partial T}{\partial p} \right)_s (\delta s)_p$$

$$f \frac{\partial \underline{v}}{\partial p} = - \left( \frac{\partial T}{\partial p} \right)_s \hat{n} \times \nabla_p s$$

$$\left[ f \frac{\partial \underline{v}}{\partial p} = R \left( \hat{n} \times \nabla_p s \right) \right]$$

Geostrophic relations are degenerate in that they provide no information about the evolution of atmospheric flows

# Dynamics of "slow" motions in rotating, stratified fluid

Potential vorticity:

$$PV \equiv \frac{1}{\rho} [\tilde{\omega}_a \cdot \nabla s]$$

$$\tilde{\omega}_a = 2\tilde{\Omega} + \nabla \times \tilde{V}_r$$

Rossby number small:  $R_o \equiv \frac{1}{f\tau} \quad \tau \sim L/u_o \Rightarrow R_o \sim \frac{u_o}{fL}$

$$\tilde{\omega}_a = f \left[ \underbrace{\frac{2\tilde{\Omega}}{f}}_{O(1)} + \underbrace{\frac{\nabla \times \tilde{V}_r}{f}}_{O(R_o)} \right]$$

Assumption:  $\nabla s$  dominated by mean stratification:

$$s = \bar{s}(z) + s' \quad \nabla s = \frac{d\bar{s}}{dz} \hat{n} + \nabla s'$$

$$= \frac{d\bar{s}}{dz} \left[ \hat{n} + \frac{\nabla s'}{d\bar{s}/dz} \right]$$

Assumption:  $\frac{\nabla s'}{d\bar{s}/dz} = O(R_o)$

~~PV~~ also:  $f = 2\Omega \sin \theta = 2\Omega \sin \theta_o \frac{\sin \theta}{\sin \theta_o} \approx f_o \frac{\sin \theta_o + \cos \theta_o (\theta - \theta_o)}{\sin \theta_o}$

$$\approx f_o + \beta y$$

$$\beta \equiv \frac{2\Omega \cos \theta_o}{R}$$

Another assumption:  $\frac{\beta_y}{f_0} \sim O(R_0)$

$$PV = \frac{1}{\rho} \left[ (2\Omega + \nabla \times \underline{V}_r) \cdot \frac{d\underline{s}}{dz} \left( \hat{n} + \frac{\nabla s'}{d\underline{s}/dz} \right) \right]$$

$$= \frac{d\underline{s}/dz}{\rho} \left[ \left[ (f_0 + \beta_y) \hat{n} + \nabla \times \underline{V}_r \right] \cdot \left[ \hat{n} + \frac{\nabla s'}{d\underline{s}/dz} \right] \right]$$

$$= f_0 \frac{d\underline{s}/dz}{\rho} \left[ \left( 1 + \frac{\beta_y}{f_0} \right) \hat{n} + \frac{\nabla \times \underline{V}_r}{f_0} \right] \cdot \left[ \hat{n} + \frac{\nabla s'}{d\underline{s}/dz} \right]$$

Retain terms of  $O(1) + O(R_0)$

$$PV = f_0 \frac{d\underline{s}/dz}{\rho} \left[ 1 + \frac{\beta_y}{f_0} + \frac{d\underline{s}'/dz}{d\underline{s}/dz} + \frac{1}{f_0} \left( \frac{\partial v}{\partial x} - \frac{\partial u}{\partial y} \right) \right]$$

Note:  $V = V_g + O(R_0)$

$U = U_g + O(R_0)$

$f_0$  be constant,

$$PV = f_0 \frac{d\underline{s}/dz}{\rho} \left[ 1 + \frac{\beta_y}{f_0} + \frac{1}{f_0} \left( \frac{\partial v_g}{\partial x} - \frac{\partial u_g}{\partial y} \right) + \frac{1}{d\underline{s}/dz} \frac{ds'}{dz} \right]$$

Let  $\phi = \bar{\phi}(z) + \phi'$

$$v_g = \frac{1}{f} \frac{\partial \phi'}{\partial x}$$

$$u_g = -\frac{1}{f} \frac{\partial \phi'}{\partial y}$$



$$\frac{\partial \phi}{\partial p} = -\alpha$$

$$\frac{\partial \phi'}{\partial p} = -\alpha' = -\left(\frac{\partial \alpha}{\partial p}\right)_s s'$$

$$\left[ \frac{\partial \phi'}{\partial z} = \rho s' \right]$$

$$PV \approx f_0 \frac{ds/dz}{\rho} \left[ 1 + \frac{\beta y}{f_0} + \frac{1}{f_0^2} \left( \frac{\partial^2 \phi'}{\partial x^2} + \frac{\partial^2 \phi'}{\partial y^2} \right) + \frac{1}{\rho ds/dz} \frac{\partial^2 \phi'}{\partial z^2} \right]$$

$$\uparrow \frac{ds}{dz} = N^2$$

No friction or heating:  $\frac{d PV}{dt} = 0$

$$\frac{\beta}{f_0} \frac{dy}{dt} + \frac{d}{dt} \left[ \frac{1}{f_0^2} \left( \frac{\partial^2 \phi'}{\partial x^2} + \frac{\partial^2 \phi'}{\partial y^2} \right) + \frac{1}{N^2} \frac{\partial^2 \phi'}{\partial z^2} \right] = 0$$

$$\frac{d}{dt} \left[ \frac{1}{f_0} \left( \frac{\partial^2 \phi'}{\partial x^2} + \frac{\partial^2 \phi'}{\partial y^2} \right) + \frac{f_0}{N^2} \frac{\partial^2 \phi'}{\partial z^2} \right] = -\beta V \approx -\beta V_y$$

also:  $\frac{d}{dt} = \frac{\partial}{\partial t} + u \frac{\partial}{\partial x} + v \frac{\partial}{\partial y} + w \frac{\partial}{\partial z}$

Note:  $\delta z \sim \frac{f_0}{N} \delta x$

$$\frac{\partial w}{\partial z} \sim -\frac{\partial u}{\partial x} - \frac{\partial v}{\partial y} = -\frac{\partial u_a}{\partial x} - \frac{\partial v_a}{\partial y}$$

$$w \sim \frac{\delta z}{\delta x} O(u_a)$$

$$\frac{d}{dt} \approx \frac{\partial}{\partial t} + u_g \frac{\partial}{\partial x} + v_g \frac{\partial}{\partial y}$$

$$PV \approx \frac{1}{f_0} \left( \frac{\partial^2 \phi'}{\partial x^2} + \frac{\partial^2 \phi'}{\partial y^2} \right) + \frac{f_0}{N^2} \frac{\partial^2 \phi'}{\partial z^2}$$

$$\left[ \left( \frac{\partial}{\partial t} + \underline{V_g} \cdot \nabla \right) PV = - B V_g \right]$$

Everything a function of  $\phi'$ !

Same idea as  $\frac{dS}{dt} = 0 \quad S = \nabla^2 \phi$

$$\left( \frac{\partial}{\partial t} - \frac{\partial \phi}{\partial y} \frac{\partial}{\partial x} + \frac{\partial \phi}{\partial x} \frac{\partial}{\partial y} \right) \nabla^2 \phi = 0$$

Scheme: Advection PV around, invert PV to find  $\phi'$

Just like barotropic case, but elliptic operator is 3-D

Slow dynamics of rotating, stratified flows  
Quasi-geostrophy

$$PV = \frac{1}{f_0} \left( \frac{\partial^2 \psi'}{\partial x^2} + \frac{\partial^2 \psi'}{\partial y^2} \right) + \frac{f_0}{N^2} \frac{\partial^2 \psi'}{\partial z^2}$$

$$\left( \frac{\partial}{\partial t} + \underline{V}_g \cdot \nabla \right) PV' = -\beta V_g$$

Longwave about  $\underline{V}_g = \bar{u}_g \hat{i} = \text{constant}$

2-D:

$$\left( \frac{\partial}{\partial t} + \bar{u}_g \frac{\partial}{\partial x} \right) \left( \frac{1}{f_0} \right) \left( \frac{\partial^2 \psi'}{\partial x^2} + \frac{\partial^2 \psi'}{\partial y^2} \right) = -\frac{\beta}{f_0} \frac{\partial \psi'}{\partial x}$$

$i\omega(x-t) + iky$

$$(\bar{u}_g - c)(k^2 + l^2) = \frac{\beta k}{k^2 + l^2}$$

$$\left[ c = \bar{u}_g - \frac{\beta k}{k^2 + l^2} \right] \quad \text{Rossby waves in 2-D}$$

The Early Model of the atmosphere

$$\psi = u_g k - \frac{\beta k^2}{k^2 + l^2}$$

$$\frac{\partial \psi}{\partial k} = u_g - \frac{2\beta k}{k^2 + l^2} + \frac{2\beta k^3}{(k^2 + l^2)^2} = u_g + \frac{2\beta k l^2}{(k^2 + l^2)^2}$$

$\text{For } l=0$

$$\omega = u_y u - \frac{\beta u}{u^2 + \lambda^2}$$

$$c_{gx} = \frac{\partial \omega}{\partial u} = u_y - \frac{\beta}{u^2 + \lambda^2} + \frac{2\beta u^2}{(u^2 + \lambda^2)^2}$$

$$= u_y + \beta \frac{u^2 - \lambda^2}{(u^2 + \lambda^2)^2}$$

$$\lambda = 0: c_{gx} = u_y + \beta/u^2 > u_y$$

$$c_{gy} = \frac{2\beta u \lambda}{(u^2 + \lambda^2)^2} \quad \text{depends on slope of phase lines in } x-y \text{ plane}$$

The Eady model



$$pV = \text{constant}$$

$$(pV' = 0)$$



$$\left( \frac{\partial}{\partial t} + u_y \cdot \nabla \right) \left[ \frac{1}{f_0} \left( \frac{\partial \psi'}{\partial x^2} + \frac{\partial \psi'}{\partial y^2} \right) + \frac{f_0}{N^2} \frac{\partial^2 \psi'}{\partial z^2} \right] = 0$$

Looks trivial, but need to apply sensible boundary conditions

$\Rightarrow$  Periodic in  $x, y$

What about in  $z$ ?

$\psi' = 0$  on  $z=0, H$  ... no pressure perturbations on boundaries

Remember that  $\frac{\partial \psi'}{\partial z} = p s'$

Setting  $\frac{\partial \psi'}{\partial z} = 0$  on boundaries gives no  $s'$  (or  $T'$ ) there

Equation for  $s'$  on boundaries:

$$\left( \frac{\partial}{\partial t} + \underline{U}_y \cdot \nabla \right) s' = 0 \quad (\text{also})$$

(Note: this is not valid in interior of flow!)

~~$$\left( \frac{\partial}{\partial t} + \underline{U}_y \cdot \nabla \right) s' = 0$$~~

Steady model:

$$\frac{\partial^2 \phi'}{\partial x^2} + \frac{\partial^2 \phi'}{\partial y^2} + \frac{f_0^2}{N^2} \frac{\partial^2 \phi'}{\partial z^2} = 0 \quad \text{everywhere}$$

$$\frac{\partial \phi'}{\partial z} = \rho s'$$

$$\left( \frac{\partial}{\partial t} + \underline{U}_y \cdot \nabla \right) s' = 0 \quad \text{on } z=0, H$$

Dynamics enters only through B.C.s!

Linearize B.C. about constant north-south  $T$  gradient:

$$\left( \frac{\partial}{\partial t} + \bar{U}_y \frac{\partial}{\partial x} \right) s' + v' \frac{d\bar{s}}{dy} = 0$$

Thermal wind:  $f_0 \frac{d\bar{u}_y}{dz} = -\rho \frac{d\bar{s}}{dy}$

$$U_{g_0} = + \frac{\rho H}{2f_0} \frac{d\bar{s}}{dy}$$

$$U_{g_H} = - \frac{\rho H}{2f_0} \frac{d\bar{s}}{dy}$$



$$\left( \frac{\partial}{\partial t} + \frac{\rho H}{2f_0} \frac{d\bar{s}}{dy} \frac{\partial}{\partial x} \right) \frac{\partial \psi'}{\partial z} + \frac{\rho}{f_0} \frac{\partial \psi'}{\partial x} \frac{d\bar{s}}{dy} = 0 \quad \text{on } z = 0$$

$$\left( \frac{\partial}{\partial t} + \frac{\rho H}{2f_0} \frac{d\bar{s}}{dy} \frac{\partial}{\partial x} \right) \frac{\partial \psi'}{\partial z} + \frac{\rho}{f_0} \frac{\partial \psi'}{\partial x} \frac{d\bar{s}}{dy} = 0 \quad \text{on } z = H$$

Non-dimensionalization :

$$z \rightarrow H z \quad t \rightarrow - \frac{\rho H}{N} \frac{d\bar{s}}{dy} t$$

$$x, y \rightarrow \frac{N}{f_0} H (x, y)$$

~~At~~

$$\nabla_3^2 \psi' = 0 \quad \text{interior}$$

$$\left( \frac{\partial}{\partial t} + \frac{1}{2} \frac{\partial}{\partial x} \right) \frac{\partial \psi'}{\partial z} - \frac{\partial \psi'}{\partial x} = 0 \quad \text{on } z = 0$$

$$\left( \frac{\partial}{\partial t} + \frac{1}{2} \frac{\partial}{\partial x} \right) \frac{\partial \psi'}{\partial z} - \frac{\partial \psi'}{\partial x} = 0 \quad \text{on } z = 1$$

$$e^{iK(x-ct) + i\ell y}$$

$$\frac{d^2 \psi'}{dz^2} - (K^2 + \ell^2) \psi' = 0$$

$$\left( \frac{1}{2} + c \right) \frac{d\psi'}{dz} + \psi' = 0 \quad \text{on } z = 0$$

$$\left( \frac{1}{2} + c \right) \frac{d\psi'}{dz} - \psi' = 0 \quad \text{on } z = 1$$

$$\psi' = A \cosh r z + B \sinh r z \quad r^2 \equiv K^2 + \ell^2$$

$$B r \left( \frac{1}{2} + c \right) + A = 0$$

$$r \left( \frac{1}{2} - c \right) (A \sinh r + B \cosh r) - A \cosh r - B \sinh r = 0$$

$$r \left( \frac{1}{2} - c \right) [-r \left( \frac{1}{2} + c \right) \sinh r + \cosh r] - \sinh r + \cosh r \left( r \left( \frac{1}{2} + c \right) \right) = 0$$

~~$$r^2 \sinh r \left( \frac{1}{4} - c^2 \right) + r \cosh r \left[ \frac{1}{2} - c \right]$$~~

$$-r^2 \sinh r \left( \frac{1}{4} - c^2 \right) + r \cosh r \left[ \frac{1}{2} + c + \frac{1}{2} - c \right] - \sinh r = 0$$

$$r^2 \left( c^2 - \frac{1}{4} \right) + r \cosh r - 1 = 0$$

$$\left[ c^2 = \frac{1}{4} + \frac{1 - r \cosh r}{r^2} \right]$$

Remember that  $r^2 = k^2 + l^2$

Shortwave limit:  $k^2 + l^2 \rightarrow \infty, r \rightarrow \infty$

$$c^2 \rightarrow \frac{1}{4}$$

$$\left[ c = \pm \frac{1}{2} \right]$$

Boundary Eady waves

$$r\left(\frac{1}{2}-c\right)B_3 - A = 0$$

$$r\left(\frac{1}{2}+c\right)(A\sinh r + B\cosh r) + A\cosh r + B\sinh r = 0$$

$$r\left(c+\frac{1}{2}\right)(B\cosh r + r\left(\frac{1}{2}-c\right)\sinh r) + \sinh r + r\cosh r\left(\frac{1}{2}-c\right) = 0$$

$$\left[ r\cosh r + r^2 \right]$$

$$r^2\left[\frac{1}{4}-c^2\right]\sinh r + r\cosh r\left[\cancel{c+\frac{1}{2}} + \frac{1}{2}-\cancel{c}\right] + \sinh r = 0$$

$$1 + r^2\left(\frac{1}{4}-c^2\right) + r\cosh r = 0$$

$$\left[ c^2 = \frac{1}{4} + \frac{r\cosh r + 1}{r^2} \right]$$

Longwave limit:

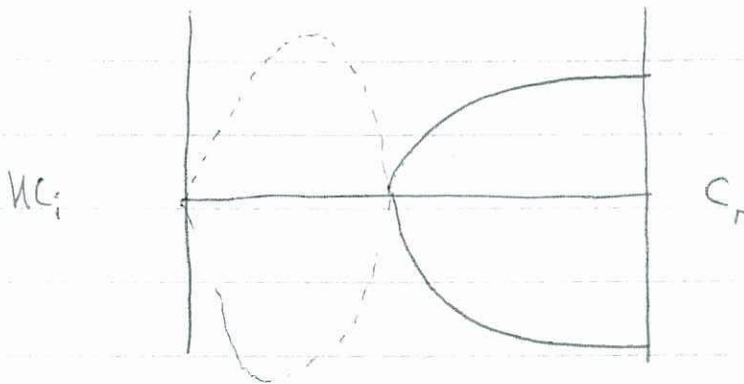
$$\omega \rightarrow \frac{1}{\sqrt{2}}$$

$$C^2 \sim \frac{1 - r \frac{\text{Cosh } r}{\text{Shch } r}}{r^2} \sim \frac{1 - r \frac{1 + \frac{1}{2}r^2}{r + \frac{1}{6}r^3}}{r^2}$$

$$= \frac{1 - (1 + \frac{1}{2}r^2)(1 - \frac{1}{6}r^2)}{r^2} = -\frac{1}{3}$$

$$C = \pm \sqrt{\frac{1}{3}} i$$

$$\left[ \kappa C_i = \pm \kappa / \sqrt{3} \right]$$



$$C^2 = 0 \quad \text{when} \quad \frac{1}{4}r^2 + 1 - r \text{ch } r = 0$$

Interaction of ~~fluid~~ boundary Eddy waves:

$$\nabla^2 \phi' = 0$$

$$\phi' \rightarrow 0 \text{ at } z \rightarrow \infty$$

$$\left(\frac{1}{2} + c\right) \frac{d\phi'}{dz} + \phi' = 0 \text{ on } z = 0$$

$$\phi' = A e^{-rz}$$

$$-r\left(\frac{1}{2} + c\right) + 1 = 0$$

$$\left[ c = -\frac{1}{2} + \frac{1}{r} \right] \text{ Bottom}$$

Likewise, for top boundary only,  $\phi' = A e^{r(z-1)}$

$$r\left(\frac{1}{2} - c\right) - 1 = 0$$

$$\left[ c = \frac{1}{2} - \frac{1}{r} \right] \text{ Top}$$

Interacting Eddy edge waves



## Tropical Cyclones

## \* Observed Characteristics and Climatology

Mature storm energetics:

How intense can a hurricane get?

Heat source: Ocean surface + dissipative heating

per unit area:  $\int C_H |\underline{V}_s| (V_s^* - u) + \int C_D |\underline{V}_s|^3$   $K \equiv C_H T + L \alpha g$

Carnot theorem: maximum energy available from cycle =  $\epsilon Q$ 

$$\epsilon = \frac{T_s - T_0}{T_s}$$

$$\text{Energy input} = \frac{T_s - T_0}{T_s} 2\pi \int_0^{r_0} \rho \left[ C_H |\underline{V}_s| (V_s^* - u) + C_D |\underline{V}_s|^3 \right] r dr$$

$$= \text{dissipation} = 2\pi \int_0^{r_0} \rho C_D |\underline{V}_s|^3 r dr$$

$$\Rightarrow 2\pi \int_0^{r_0} \rho \left[ \frac{T_s - T_0}{T_s} C_H |\underline{V}_s| (V_s^* - u) - \frac{T_0}{T_s} C_D |\underline{V}_s|^3 \right] r dr = 0$$

If dominated by conduction near radius of maximum winds:

$$\left[ |\underline{V}_s|^2 \approx \frac{T_s - T_0}{T_0} \frac{C_H}{C_D} (V_s^* - u) \right]$$

$$k_s^* - k \approx L_v (q_s^* - q) \quad q = e^{e^*} / p$$

$$\approx e L_v \frac{e^*(T)}{p} (1-H)$$

$$|V_s|^2 \approx \frac{T_s - T_0}{T_0} \frac{C_H}{C_D} e L_v \frac{e^*(T_s)}{p} (1-H)$$

$$\left[ p |V_s|^2 \approx \frac{T_s - T_0}{T_0} \frac{C_H}{C_D} \frac{L_v}{R_v T_s} e^*(T_s) (1-H) \right]$$

Disturbance here increases greatly with  $e^*(T_s)$

## Read Chapter 6

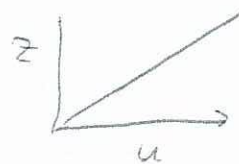
### Shear Stresses

Random motions of molecules transport momentum within fluid and between fluids at rigid surfaces.

Newton: Shear stress linearly related to velocity

For simple 1-D shear flows,

$$\tau = \eta \frac{\partial u}{\partial z}$$



$\eta$  a function of fluid.

For general flows,  $\tau$  is a tensor

Example: Flow between parallel plates

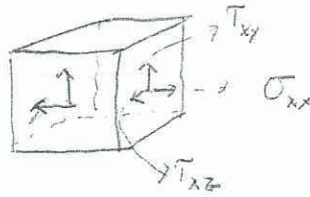
$$\frac{\partial^2 u}{\partial z^2} = \frac{\partial p}{\partial x}$$

$$u = 0 \text{ on } z = 0, H$$

$$u = \frac{1}{2\eta} \left( z^2 - \frac{1}{4} \right) \frac{\partial p}{\partial x}$$

$$u = \frac{1}{2\eta} z(z-H) \frac{\partial p}{\partial x}$$

Stress Terms in Equations:



$$\text{Force in } x \text{ direction} = \frac{\partial}{\partial x}(\sigma_{xx}) dx dy dz + \frac{\partial}{\partial y}(\tau_{yx}) dx dz + \frac{\partial}{\partial z}(\tau_{zx}) dx dy$$

$$\rho \frac{du}{dt} = \rho f_x + \frac{\partial}{\partial x}(\sigma_{xx}) + \frac{\partial}{\partial y}(\tau_{yx}) + \frac{\partial}{\partial z}(\tau_{zx})$$

etc.

Stresses related linearly to velocity, (~~isotropic~~) are isotropic

$$\tau_{xy} = \tau_{yx} \quad \text{etc.}$$

Turns out that

$$\sigma_{xx} = -p - \frac{2}{3}\mu(\nabla \cdot \underline{v}) + 2\mu \frac{\partial u}{\partial x}$$

$$\sigma_{yy} = -p - \frac{2}{3}\mu(\nabla \cdot \underline{v}) + 2\mu \frac{\partial v}{\partial y}$$

$$\sigma_{zz} = -p - \frac{2}{3}\mu(\nabla \cdot \underline{v}) + 2\mu \frac{\partial w}{\partial z}$$

$$\tau_{xy} = \tau_{yx} = \mu \left( \frac{\partial u}{\partial y} + \frac{\partial v}{\partial x} \right)$$

$$\tau_{xz} = \tau_{zx} = \mu \left( \frac{\partial u}{\partial z} + \frac{\partial w}{\partial x} \right)$$

$$\tau_{yz} = \tau_{zy} = \mu \left( \frac{\partial v}{\partial z} + \frac{\partial w}{\partial y} \right)$$

Example: Unidirectional Flow

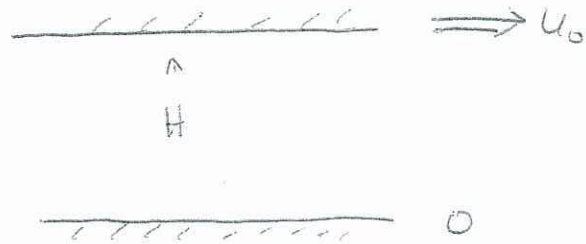
$$\rho \frac{du}{dt} = - \frac{\partial p}{\partial x} + \frac{\partial}{\partial x} \left( \mu \frac{\partial u}{\partial x} \right) + \frac{\partial}{\partial y} \left( \mu \frac{\partial u}{\partial y} \right) + \frac{\partial}{\partial z} \left( \mu \frac{\partial u}{\partial z} \right)$$

If  $u = u(z)$   $\rho \frac{du}{dt} = - \frac{\partial p}{\partial x} + \frac{\partial}{\partial z} \left( \mu \frac{\partial u}{\partial z} \right)$

Constant  $\mu$ :

$$\frac{du}{dt} = - \frac{1}{\rho} \frac{\partial p}{\partial x} + \mu \nabla^2 u + \frac{1}{3} \mu (\nabla \cdot \underline{v})$$





$$u = u_0 \frac{z}{H}$$

Oscillating Plate:



$$u_{z=0} = u_0 \cos(\omega t)$$

$$\frac{\partial u}{\partial t} = \nu \frac{\partial^2 u}{\partial z^2}$$

Free solution:  $u = e^{irz + i\omega t}$

$$+ \nu r^2 = i\omega$$

$$r = \pm (1+i) \sqrt{\frac{\omega}{2\nu}} \equiv \pm \frac{(1+i)}{\delta}$$

$$\delta \equiv \sqrt{\frac{2\nu}{\omega}}$$

$$\left[ u \sim u_0 \cos\left(\frac{z}{\delta} - \omega t\right) e^{-z/\delta} \right]$$

EKman Layer:

$$0 = \frac{\partial u}{\partial t} = -\frac{1}{f} \frac{\partial p}{\partial x} + f v + \nu \frac{\partial^2 u}{\partial z^2} = f(v - V_g) + \nu \frac{\partial^2 u}{\partial z^2}$$

$$0 = \frac{\partial v}{\partial t} = -f u + \nu \frac{\partial^2 v}{\partial z^2} \quad V, u = 0 \text{ on } z = l$$

$$\nu^2 \frac{\partial^4 v}{\partial z^4} = f^2 (V_g - v)$$

$$\left[ \nu^2 \frac{\partial^4 v}{\partial z^4} + f^2 v = f^2 V_g \right]$$

Inhomogeneous:  $v = V_g$

$$\text{Homogeneous: } \nu^2 \frac{\partial^4 v}{\partial z^4} + f^2 v = 0$$

$$v = A e^{i r z}$$

$$\left[ r^4 = -f^2 / \nu^2 \right]$$

$$r^4 = \frac{f^2}{\nu^2} e^{i\pi + 2i n \pi}$$

$$r = \sqrt[4]{\frac{f}{\nu}} e^{i\pi/4 + i n \pi/2}$$

$$n = 0, 3$$

$$r_4 = \sqrt[4]{\frac{f}{\nu}} [1+i, 1-i, -1+i, -1-i]$$

$$\delta \equiv \sqrt[4]{2\nu/f}$$

$$u = e^{-z/\delta} \left[ \cancel{A \cos z/\delta} + B \sin z/\delta \right]$$

$$u = V_g + e^{-z/\delta} \left[ A \cos z/\delta + B \sin z/\delta \right]$$

$$A = -V_g$$

$$V = V_g + e^{-z/\delta} \left[ B \sin z/\delta - V_g \cos z/\delta \right]$$

$$u = \frac{1}{f} \frac{\partial^2 V}{\partial z^2} = \frac{1}{\delta} \frac{1}{f} \frac{\partial}{\partial z} \left[ -B \sin z/\delta + V_g \cos z/\delta + B \cos z/\delta + V_g \sin z/\delta \right] e^{-z/\delta}$$

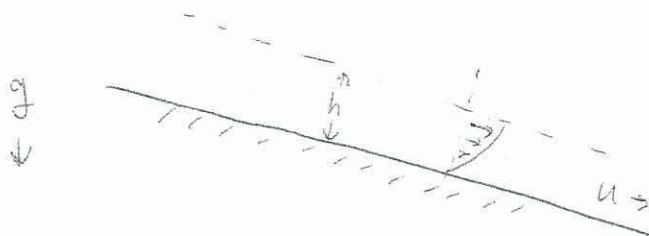
$$= \frac{1}{\delta^2} \frac{1}{f} \left[ \cancel{B \sin z/\delta} - \cancel{V_g \cos z/\delta} - 2B \cos z/\delta - 2V_g \sin z/\delta - \cancel{B \cos z/\delta} - \cancel{V_g \sin z/\delta} - \cancel{B \sin z/\delta} + \cancel{V_g \cos z/\delta} \right] e^{-z/\delta}$$

$$= - \left[ B \cos z/\delta + V_g \sin z/\delta \right] e^{-z/\delta}$$

$$\Rightarrow B = 0$$

$$u = -V_g \sin z/\delta e^{-z/\delta}$$

One more Viscous flow problem:



Incompressible flow on a slope

$$\frac{du}{dt} = 0 = -g \frac{\partial h}{\partial x} + \nu \frac{\partial^2 u}{\partial z^2}$$

$$u = -\frac{g}{\nu} \frac{\partial h}{\partial x} z \left( h - \frac{1}{2} z \right)$$

$$u = 0 \text{ on } z = 0$$

$$\nu \frac{\partial u}{\partial z} = 0 \text{ on } z = h$$

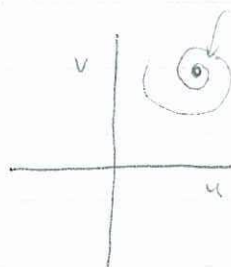
## \* Turbulence (Real Fluids Chapter 9)

Last great frontier in classical physics

Often triggered by instability

General Character of Nonlinear Systems:

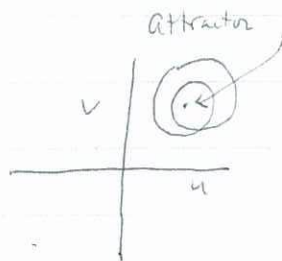
1. Steady, Stable



Steady Phase Point

All arbitrary initial states decay to one or more steady phase points

2. Periodic

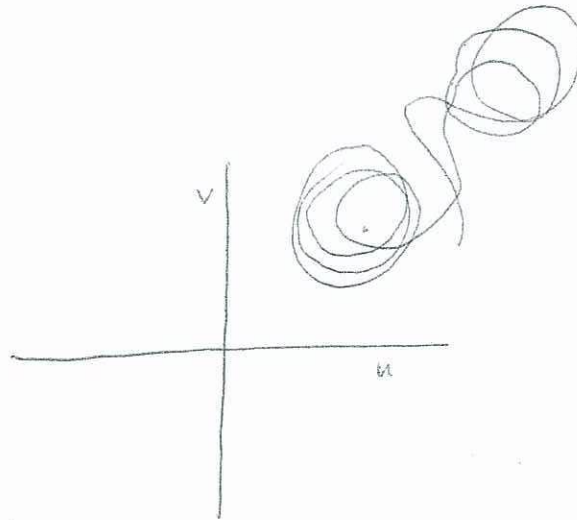


Attractor

Period  $2T$  in this case

Small perturbations decay onto orbit or form adjacent orbit

### 3. Chaotic



- Characteristics:
1. Never repeats
  2. Two arbitrary points may be separated by infinite length along phase line but infinitesimal distance in phase space
  3. Distance between two adjacent points grows exponentially in time: extreme sensitivity to initial conditions

Routes to Chaos: Period doubling  
Intermittency

### Characterization of Turbulence

Ex: Pipe Flow:



Laminar: away from inlet,  $u = \bar{u}(z)$

Turbulent: Define ~~the~~ <sup>ensemble</sup> average  $\bar{u}$ . away from inlet,  $\bar{u} = \bar{u}(z)$

Characterization: Auto correlation function

$$\langle u_i(\underline{x}, t) u_i(\underline{x}', t) \rangle$$

or

$$\langle u_i(\underline{x}, t) u_i(\underline{x}', t') \rangle$$



Also we may look at the Fourier transform of  $\underline{u}(\underline{x}, t)$ :  $\underline{u}(\underline{k}, \omega)$

Structure of turbulence in this case constrained by pipe, so statistical theories applied usually to homogeneous turbulence: turbulence does not vary statistically in space. If moving with  $\bar{u}$ , may be isotropic as well.

Nature of problem: Let  $\underline{u} = \bar{\underline{u}}(\underline{x}, y, z, t) + \underline{u}'(\underline{x}, y, z, t)$

$$\frac{\partial \underline{u}}{\partial t} + \underline{V} \cdot \nabla \underline{u} = -\frac{1}{\rho} \frac{\partial p}{\partial x}$$

$$\begin{aligned} \frac{\partial \bar{\underline{u}}}{\partial t} + \frac{\partial \underline{u}'}{\partial t} + \bar{\underline{V}} \cdot \nabla \bar{\underline{u}} + \underline{V}' \cdot \nabla \bar{\underline{u}} + \bar{\underline{V}} \cdot \nabla \underline{u}' + \underline{V}' \cdot \nabla \underline{u}' \\ = -\frac{1}{\rho} \frac{\partial \bar{p}}{\partial x} - \frac{1}{\rho} \frac{\partial p'}{\partial x} \end{aligned}$$

Ensemble Average: ~~the time~~

$$\frac{\partial \bar{\underline{u}}}{\partial t} + \bar{\underline{V}} \cdot \nabla \bar{\underline{u}} + \overline{\underline{V}' \cdot \nabla \underline{u}'} = -\frac{1}{\rho} \frac{\partial \bar{p}}{\partial x}$$

To predict evolution of  $\bar{\underline{u}}$  need to know correlation  $\overline{\underline{V}' \cdot \nabla \underline{u}'}$

$$\left( \text{Can also write } \frac{\partial \bar{\underline{u}}}{\partial t} + \bar{\underline{V}} \cdot \nabla \bar{\underline{u}} + \nabla \cdot \bar{\underline{p}} \underline{\underline{V}}' \underline{\underline{u}'} = -\frac{1}{\rho} \frac{\partial \bar{p}}{\partial x} \right)$$

Can form equations for second order correlations, e.g.

$$\frac{\partial \underline{u}'}{\partial t} + \underline{V}' \cdot \nabla \bar{\underline{u}} + \bar{\underline{V}} \cdot \nabla \underline{u}' = -\frac{1}{\rho} \frac{\partial p'}{\partial x}$$

$$\frac{\partial}{\partial t} \overline{\underline{V}' \cdot \nabla \underline{u}'} + \overline{\underline{V}' \underline{V}'} \cdot \nabla \underline{u}' + \overline{\underline{V}' \bar{\underline{V}} \cdot \nabla \underline{u}'} = -\frac{1}{\rho} \overline{\underline{V}' \cdot \nabla p'}$$

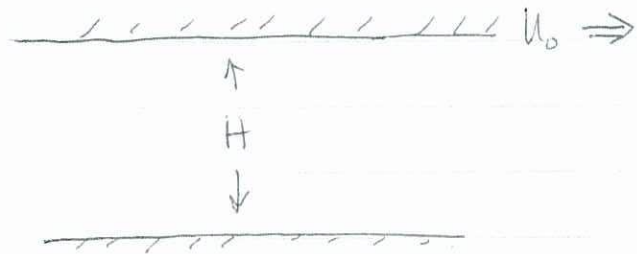
Now we need third order correlations, like  $\overline{v'v'\cdot \nabla u'}$   
etc.

### Closure Problem

Several approaches: Empirical representations of higher order correlations  
Dimensional analysis.

Some examples of turbulent flow:

Concette Flow



Control parameters:  $U_0, \nu, h$

Buckingham Pi Theorem:

$M \equiv$  number of <sup>independent</sup> dimensional control parameters

$N =$  number of fundamental dimensions (length, time, mass etc)

$P =$  number of independent dimensionless control parameters

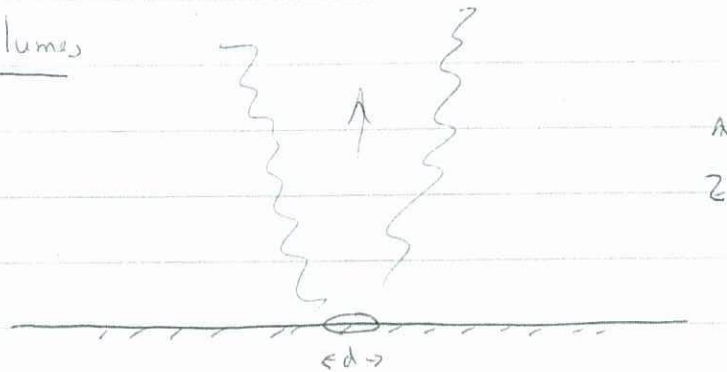
$$P = M - N$$

In this case,  $P=1$

$$R_e \equiv \frac{U_0 H}{\nu}$$

Observations: Though flow becomes unstable for sufficiently large  $R_e$

Convective Plumes



Parameters:  $F \equiv \text{Buoyancy flux} \sim wBA \sim L^4 t^{-3}$   
 $d$

$\nu$

$K$

$$M = 4$$

$$N = 2$$

$$P = 2$$

$$\sigma \equiv \nu/K$$

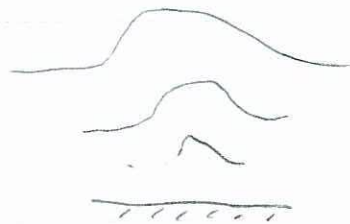
Let  $d \rightarrow 0$  ("Point Source")

$$R \equiv Fd^2/\nu^3$$

Hypothesis: If  $d, \nu$  small enough

If  $R$  large enough, properties become independent of  $R$ :

No nondimensional parameters:



Geometric similarity implied:

$$\bar{r} = \alpha z$$

$$W \sim L t^{-1} \sim F^{1/3} z^{-1/3}$$

$$M \approx \pi R^2 W \sim F^{1/3} z^{5/3}$$

implies entrainment

$$\bar{u} = \gamma W$$

Note:  $\frac{dM}{dz} = 2\pi R \bar{u} = 2\pi R \gamma W$

$$\frac{d}{dz} (\pi R^2 W) = 2\pi R \gamma W$$

$$W = A F^{1/3} z^{-1/3}$$

$$R = \alpha z$$

$$\frac{d}{dz} (A F^{1/3} \alpha^2 z^{5/3}) = 2\alpha \gamma A F^{1/3} z^{5/3}$$

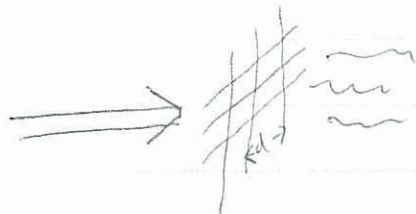
$$\frac{5}{3} \alpha = 2\gamma$$

$$\left[ \gamma = \frac{5}{6} \alpha \right]$$

Substantiated Results from Dimensional Analysis alone

## Energy Cascade in Homogeneous Turbulence

Consider what happens when uniform flow passes through gridded mesh.



Vortices shed from mesh

Vortices intertwine  $\Rightarrow$  stretching

local spin-up of vortices

Large shears eventually develop at small enough scales that molecular dissipation can be effective

Energy cascades from wavenumber  $\sim 1/d$  to higher values

Idealization: Nearly homogeneous, isotropic turbulence with energy input to largest scales,  $U_0$ :

Energy cascade to very small scales where  $\mu$  is effective

Kinetic energy per unit mass in range  $k, k+dk \equiv E(k)dk$

Control parameter: Rate of input of energy per unit mass:  
 $\epsilon_0$  (= rate of dissipation)

$$\sim L^{2+3}$$

IF  $U_0$ ,  $\nu$  irrelevant

While  $E(k) \sim L^{3+2}$

Dimensionally,  $\left[ E(k) = \cancel{\epsilon_0^{3/2}} \epsilon_0^{2/3} k^{-5/3} \right]$



\* Experimentally, this seems to be valid even near  $U_0$ , but must fail at large enough  $U$ :

$$\frac{dU}{dt} = \dots + \nu \nabla^2 U \sim -\nu U^2$$

$$E \sim \frac{d}{dt} \frac{1}{2} \langle U^2 \rangle \sim -\nu \langle U^2 \rangle^2 \sim -\nu \int U^3 E(U) dU$$

$$\frac{dE}{dU} \sim -\nu U^2 E = \frac{-\nu E_0^{2/3} U^{1/3}}{\text{if } E \sim U^{-5/3} \text{ holds}}$$

Owing to molecular dissipation

Integrate:  $E \sim E_0 - \frac{3}{4} \nu E_0^{2/3} U^{4/3}$

Second term becomes significant when  $U \sim E_0^{3/4} \nu^{-3/4} \equiv U_c$

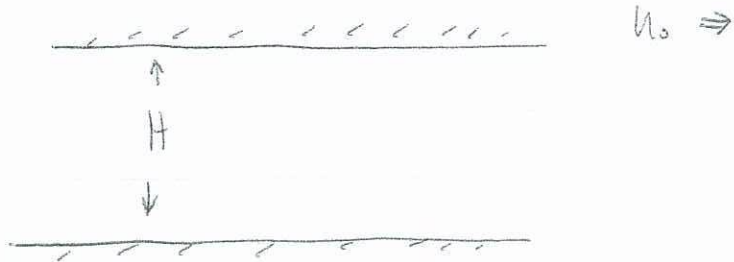


Inertial Subrange: Nonlinear cascade to small scales

Dissipation range: Molecular processes become important

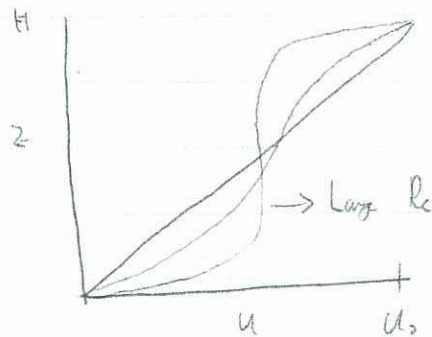
# Boundary Layers.

First consider Couette Flow:



$U_0, \nu, H$

$$Re \equiv \frac{U_0 H}{\nu}$$



At large  $Re$ , flow confined to thin boundary layers.

Near walls,  $H$  irrelevant:  $u \sim u(z, \nu, U_0)$

$$\delta \sim \sqrt{\nu/U_0} \quad \text{dimensionally}$$

$$u \sim U_0 F(z/\delta)$$

Near top of boundary layer, we expect flow to ~~become~~ shear to be independent of  $\nu, H$ .

$$\frac{du}{dz} \sim \frac{U_0}{\delta} \quad u = \frac{U_0}{K} \ln z/\delta$$

$K = \text{Von Karman's constant}$

Experiments,  $K \approx .35$

$$u = \frac{u_0}{K} \ln \frac{u_0 z}{u}$$

logarithmic "law of the wall"

In real, physical systems, boundary is usually "rough"



Characterize roughness of boundary by length scale  $z_0$ :

If  $z_0 \gg \delta$ , assume  $\delta \rightarrow$  irrelevant

$$\frac{du}{dz} = \frac{u_0}{z} \quad \left[ u = \frac{u_0}{K} \ln z/z_0 \right]$$

Applied pressure gradient:  $\frac{du}{dz} \sim \sqrt{\frac{1}{\rho} \frac{\partial p}{\partial x}} \frac{1}{z^{1/2}}$

$$u \sim \sqrt{\frac{1}{\rho} \frac{\partial p}{\partial x}} (z^{1/2} - z_0^{1/2})$$

Convective Boundary Layer: No mean wind:

Applied buoyancy flux / unit area:  $wB \sim L^2 t^{-3} \equiv F$

$$w' \sim (Fz)^{1/3}$$

$$B' \sim F^{2/3} z^{-1/3}$$

Wealth near wind: Treat as passive scalar:

Given source at surface:  $\overline{w'u'} \approx M \sim L^2 t^{-2}$

$M$  independent of height:

Dimensionally,  $\frac{d\bar{u}}{dz} =$

Flow governed by a single dimensionless parameter:

Pipe Flow:  $\nu, \frac{1}{\rho} \frac{\partial p}{\partial x}, R$

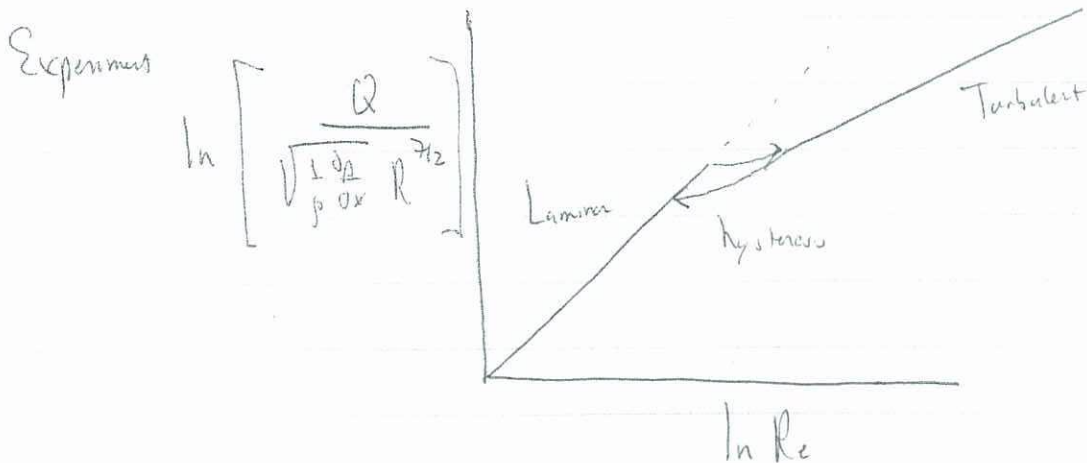
$$Re = \frac{\sqrt{\frac{1}{\rho} \frac{\partial p}{\partial x}} R^{3/2}}{\nu}$$

Flow rate should be function of  $Re$ :

$$Q \sim \int u r R^2 \sim \rho L^3 +^{-1} \sim \sqrt{\frac{1}{\rho} \frac{\partial p}{\partial x}} R^{7/2}$$

$$Q \sim \sqrt{\frac{1}{\rho} \frac{\partial p}{\partial x}} R^{7/2} F(Re)$$

Lamina Flow: Exact solution:  $F \sim Re$



Transition also depends on smoothness of pipe



From Gortstein

Behaviour of passive tracer  $\theta$  with source  $F$  in momentum BL:

$$\frac{d\bar{\theta}}{dz} = -C$$

Rough Boundary

Control parameter:  $\int_0^\infty \frac{1}{\rho} \frac{dp}{dz} dz \equiv u^*{}^2$

$$\frac{d\bar{u}}{dz} = \frac{1}{K} \frac{u^*}{z}$$

$$u = \frac{1}{K} u^* \ln^2 z/z_0$$

$$\frac{d\bar{\theta}}{dz} = -C \frac{F}{u^* z}$$

$$\bar{\theta} = \bar{\theta}_0 - C \frac{F}{u^*} \ln^2 z/z_0$$

$$\bar{\theta}_0 \equiv \bar{\theta} \text{ at } z = z_0$$

$$\text{also: } F = \frac{1}{C} u^* (\bar{\theta}_0 - \bar{\theta}_a) \left( \ln^2 z_0/z_0 \right)$$

$$\text{But } u^* = K \bar{u}_a \left( \ln^2 z_0/z_0 \right)^{-1}$$

$$\text{so } F = \frac{K}{C} \bar{u}_a (\bar{\theta}_0 - \bar{\theta}_a) \left( \ln^2 \frac{z_0}{z_0} \right)^{-1}$$

Convective boundary layer:

External control parameter = buoyancy flux  $Q_0 = \overline{w'b'}$

$$w' \sim (Q_0 z)^{1/3}$$

$$B' = g \beta T' \sim Q_0^{2/3} z^{-1/3}$$

$$q_0 \left( \equiv |w'| \text{ at } z = z_0^T \right) \sim \left( z_0^T Q_0 \right)^{1/3}$$

$$\frac{d\bar{T}}{dz} = g\beta \frac{d\bar{T}}{dz} = -C_1 Q_0^{2/3} z^{-4/3}$$

$$\bar{T} = \bar{T}_0 - \frac{3C_1}{g\beta} Q_0^{2/3} \left[ (z_0^T)^{-1/3} - z^{-1/3} \right]$$

$$\bar{T} = \bar{T}_0 \quad \text{at} \quad z = z_0^T$$

$$T_A = \bar{T}(z=\infty) = \bar{T}_0 - \frac{3C_1}{g\beta} Q_0^{2/3} z_0^T^{-1/3}$$

Also,  $\overline{w'b'} = Q_0 = (3C_1)^{-3/2} (z_0^T)^{1/2} \left[ g\beta (\bar{T}_0 - \bar{T}_A) \right]^{3/2}$

Passive tracer  $\theta$  with source  $F$ :

$$\frac{d\bar{\theta}}{dz} = -C_2 F Q_0^{-1/3} z^{-4/3}$$

$$\bar{\theta} = \bar{\theta}_A + 3C_2 F Q_0^{-1/3} z^{-1/3}$$

$$\bar{\theta} = \bar{\theta}_A \quad \text{when} \quad z \rightarrow \infty$$

What if surface is both momentum sink at buoyancy source?

Now there is a definite length scale in the problem!

Momentum-obstacle length  $L \equiv \frac{u_*^3}{\overline{w'b'}} Q_0$  • *buoyancy sign*  
*is from buoyancy!*

In general, quantities will be a function of  $z/L$

Limiting cases:  $Q_0 \rightarrow 0, \quad z/L \rightarrow 0$

$u^* \rightarrow 0 \quad z/L \rightarrow \infty$

$$-g\beta \frac{d\bar{T}}{dz} = \begin{cases} +C_1 Q_0^{2/3} z^{-1/3} & z/L \rightarrow \infty \\ C \frac{Q_0}{u^* z} & z/L \rightarrow 0 \end{cases} \quad \text{passive}$$

$$+ \frac{d\bar{u}}{dz} = \begin{cases} C_2 u^{*2} Q_0^{-1/3} z^{-4/3} & z/L \rightarrow \infty \\ \frac{1}{K} u^* z^{-1} & z/L \rightarrow 0 \end{cases} \quad \text{passive}$$

$$\frac{d\bar{\theta}}{dz} = \begin{cases} -C_2 F Q_0^{-1/3} z^{-4/3} & z/L \rightarrow \infty \\ -C F / u^* z & z/L \rightarrow 0 \end{cases}$$

Candidate Functions to fit field experiment data:

$$g\beta \frac{d\bar{T}}{dz} = -C \frac{Q_0}{u^* z} \left[ 1 + \left( \frac{C}{C_1} \right)^3 \frac{z}{L} \right]^{-1/3}$$

$$\frac{d\bar{u}}{dz} = \frac{1}{K} \frac{u^*}{z} \left[ 1 + \left( \frac{1}{K C_2} \right)^3 \frac{z}{L} \right]^{-1/3}$$

$$\frac{d\bar{\theta}}{dz} = -C \frac{F}{u^* z} \left[ 1 + \left( \frac{C}{C_2} \right)^3 \frac{z}{L} \right]^{-1/3}$$

Surface Fluxes:

$$Q_0 = g \beta (\bar{T}_0 - \bar{T}_A) * \begin{cases} \frac{1}{Kc} u^* & z/L \rightarrow 0 \\ w^* & z/L \rightarrow \infty \end{cases}$$

$$M \overset{\text{friction}}{\cancel{w^*}^2} = \bar{u}_a * \begin{cases} u^* & z/L \rightarrow 0 \\ \frac{c_1}{c_2} \frac{w^*}{\theta_1} & z/L \rightarrow \infty \end{cases}$$

$$F = (\bar{\theta}_0 - \bar{\theta}_a) * \begin{cases} \frac{1}{Kc} u^* & z/L \rightarrow 0 \\ \frac{c_1}{c_2} w^* & z/L \rightarrow \infty \end{cases}$$

$$u^* \equiv K^2 \left( \ln \frac{z_a}{z_0} \right)^{-2} \bar{u}_a$$

$$w^* \equiv (3C_1)^{-3/2} \left[ g \beta (\bar{T}_0 - \bar{T}_A) z_0 \right]^{1/2}$$

"friction" and "convective" velocities

Composite functions:

$$Q_0 = g \beta (\bar{T}_0 - \bar{T}_A) \sqrt{\left( \frac{1}{Kc} u^* \right)^2 + w^{*2}}$$

$$\overset{\text{friction}}{\cancel{w^*}^2} \Rightarrow M = \bar{u}_a \sqrt{u^{*2} + \left( \frac{c_1}{c_2} w^* \right)^2}$$

$$F = (\bar{\theta}_0 - \bar{\theta}_a) \sqrt{\left( \frac{u^*}{Kc} \right)^2 + \left( \frac{c_1}{c_2} w^* \right)^2}$$

Spring 2002

# Atmospheric Convection

1. Review:  $(\partial \alpha)_p = \left( \frac{\partial \alpha}{\partial s} \right)_p ds = \left( \frac{\partial T}{\partial p} \right)_s ds$

$$B = g \frac{\delta \alpha}{\alpha} = \Gamma \delta s \quad \Gamma \equiv - \left( \frac{\partial T}{\partial z} \right)_p$$

Moist air. In general,  $\alpha = \alpha(p, s, q_+)$

$q_+$  = specific concentration of water

Saturation Law:  $p_v^* = p_v^*(T) = p_{v0} \exp \left( \frac{17.67 T_c}{243.5 + T_c} \right)$

Doubles for every  $10^\circ \text{C}$  rise in  $T$

$$q = m_v / m = p_v / p = \frac{p_v / R_v T}{p / \bar{R} T} = \frac{\bar{R}}{R_v} p_v / p = \cancel{F(T)}$$

$$q^* \equiv \frac{\bar{R}}{R_v} p_v^* / p = F(T, p)$$



Adiabatic ascent eventually leads to saturation

## Homogeneous, Heterogeneous nucleation



## Two approaches:

### Thermodynamic Diagrams

- \* Pseudo-adiabatic process
- \* Reversible Adiabatic process
- \* Conditional Instability

### Tropical Convection

Quasi & photos

Quasi-equilibrium

\* stability diagrams

Interaction of tropical convection with large-scale waves

### Extratropical Convection

- \* Air Mass Storms
- \* Squall lines
- \* Supercell Storms

# Fluid Physics

Review for Final Exam:

Spring, 2002

Euler Fluid: Definition

Hydrostatic equilibrium

incompressible fluid

ideal gas

Mass continuity

general expression

incompressible case

Total versus partial derivatives in time

Shallow water equations

conditions for validity

equations

shallow water waves: dispersion, group velocity

hydraulic jumps: wave steepening

Bernoulli equation

Irrotational Flow

Circulation Theorem

Velocity potential

point sources/sinks of mass

Compressible Fluids

Ideal gas

First Law of Thermodynamics

Sound waves

Shocks

Bernoulli equation

Convection

Buoyancy

Stability criterion